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Regional and seasonal variations in snow-cover density in the U.S.S.R.



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Michael A. Bilello

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20. Abstract (cont'd).

included in this report are 1) a compilation of pertinent passages in the Soviet literature on SCD, 2) a map showing the location of SCD measurements, and 3) an average winter wind speed chart for the U.S.S.R.

PREFACE

This report was prepared by Michael A. Bilello, Meteorologist, formerly of the Geophysical Sciences Branch, Research Division, U.S. Army Cold Regions Research and Engineering Laboratory. The work was performed under DA Project 4A762730AT42, *Design, Construction and Operations Technology for Cold Regions*, Technical Effort E1, *Environmental Control Methods*, Work Unit 005, *Winter Battlefield Climatology*.

Assistance in the effort of conducting the comprehensive survey of Soviet literature required for this study was provided by the library personnel at CRREL. The collection of climatic information for the U.S.S.R. was accomplished through the cooperation of the U.S. Air Force Environmental Technical Application Center (ETAC). The investigation was performed in conjunction with the SNOW Experiments Project under George Aitken, Program Manager.

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REGIONAL AND SEASONAL VARIATIONS IN SNOW-COVER DENSITY IN THE U.S.S.R.

Michael A. Bilello

INTRODUCTION

In a previous study (Bilello 1967), snow-cover density (SCD) observations made from the beginning of November through March were used to develop an average seasonal SCD map for northern North America. The study investigated the relationship between climatic variables and regional variations in SCD and showed that the data could be grouped into four categories of snow density. The objective of the study described here was to adopt the approach used previously and to develop a similar map for the U.S.S.R. This report also includes excerpts of pertinent SCD data from the Soviet Union, and other associated maps obtained from a survey of available publications.

Some significant areas in which the results of these SCD studies would be applicable to winter environmental technology are 1) the thermal effects of snow density on the depth of frost penetration in the ground and on the growth of ice on water bodies, 2) the role of snow density in the ability for wheeled or tracked vehicles to move over snow-covered terrain, 3) the problems associated with snow removal on roads or buildings and 4) the importance of snow density data in various hydrologic and agricultural activities that contribute to the national economy.

With recent advances in aerial photography and satellite imagery, knowledge of specific features of the earth's surface in winter, such as the snow-cover density, is becoming particularly useful in a number of interpretative analyses. Information on the physical characteristics of the snow cover, especially its density, is essential in the analysis of photographs or satellite images of a snow-covered terrain.

SOVIET SNOW-COVER DENSITY DATA

Soviet sources

An important source of Soviet data sources resulted from a United States-Soviet Scientific Exchange Visit to the Kolyma Water Balance Station in the Magadan Oblast, U.S.S.R., by the author and a representative of the U.S. Forest Service (Slaughter and Bilello 1977). During this trip, the watershed hydrology, instrumentation, field practices, and Soviet snow research objectives were reviewed. In addition to the useful discussions regarding snow measuring equipment and snow-cover conditions observed in that area of the country, the author later received a copy of a book by Kopanev (1978) entitled *Snow Cover in the U.S.S.R.* Average monthly SCD measurements made at hydrometeorological stations across the U.S.S.R. were tabulated in this book.

Additional SCD data were provided to the author in 1981 by a member of the Soviet Academy of Sciences. Observed densities for 105 locations in the Soviet Union were used in the development of a map on the regional distribution of SCD. The application, adjustments, and analysis of these data are discussed in detail later in the report.

Literature sources

The 36 volumes of the CRREL *Bibliography on Cold Regions Science and Technology* (U.S. Army Cold Regions Research and Engineering Laboratory 1951-1982) were used as the primary source for a comprehensive search for information on SCD in the Soviet Union. A search through the 36 volumes provided a substantial list of articles on the subject under study. Further examination of the abstracts of the selected references revealed 24 Russian articles that contained information pertinent to this study and many other papers that presented snow-cover research not directly applicable to this study. The following procedures for incorporating the pertinent SCD material given in these Soviet articles were used in this report.

Pertinent SCD articles

The 24 Russian articles that provided information on SCD in the Soviet Union were translated (when necessary) and pertinent subject material was extracted and included as an appendix to this report (App. A). A review of information given in Appendix A indicated that it would be worthwhile to further extract all the quoted *numerical* values on SCD and to tabulate the data with respect to the area in which they were recorded. These areas were then located on a map (Fig. A1) of the U.S.S.R. so that the reported densities could be easily associated with the areas to which they apply. Statements of interest regarding snow cover conditions and further details on the data summaries are given in Appendix A. The 24 references that were used for this survey are also listed in the appendix.

Other articles of interest

The numerous articles that appear to provide interesting and possibly useful information on snow-cover characteristics across northern Eurasia are listed separately as a CRREL Internal Report 840. This list of over 150 items includes articles prepared in languages besides Russian. However, except for their titles (and some abstracts) the foreign papers were not translated. The bulk of these papers provided results of re-

search on essential aspects of the snow cover such as 1) instrumentation and field measurement evaluations, 2) snow conditions in specific regions of the U.S.S.R. and in other European countries, 3) areal variability in snow properties due to elevation and exposure, and 4) interrelationships between the snow cover and hydrology and/or meteorology.

UTILIZATION OF SOVIET SNOW-COVER DENSITY DATA

Station selection

Although average monthly SCD values for 137 field locations were provided by the two Soviet sources, the data for only 105 of those stations were used in this study. The other 32 Soviet stations were excluded for the following reasons:

1. Since no information was provided regarding precise locations of any of the reported stations, it was necessary to determine and verify the exact coordinates for each site. Three sources were used to obtain this information: 1) an official standard name gazetteer for the U.S.S.R. (U.S. Army 1970), 2) a National Geographic Society Map of the Soviet Union (National Geographic Society 1976), and 3) an Army Map Service, Asian Series Map (U.S. Army 1962 and 1964). Latitude and longitude values, to the degree and nearest minute, based on at least two of the above three sources were obtained and verified for the 105 stations. Confirmation of the position of many of the other sites could not be determined because a number of different coordinates were given but the correct one could not be ascertained. Other excluded locations were not listed in any of the above sources.

2. Density readings for only three winter months were provided in the tabulations for some of the stations. As noted below, the period from November through March was used to compute an average "seasonal" SCD. This portion of the winter was also used in the North American study because it generally includes months with below-freezing average air temperatures. It was decided, therefore, that a 3-month "winter" would be too brief for inclusion in this investigation.

3. A few stations were found to appear in both the Soviet data sources. In these cases, the reported densities were compared for verification purposes and the station was listed only once.

Data discrepancies

The author's initial optimism regarding the

SCD values received from the Soviet sources for this study gradually diminished after further evaluation of the data. The first indication of a problem arose when the reported SCD values for the Soviet Union were compared with those observed across northern North America. Stations that exhibited very similar winter climatic regimes (for example, arctic stations that recorded very cold temperatures and strong winds or inland stations with cold air temperatures but very light winds) were used in the comparison. Results from representative stations within these regimes showed that average SCD's for the Soviet stations ranged from 18 to 27% lower than those recorded in North America.

These disagreements could perhaps be explained by differences in the method of observation, the length of records, and the procedures for computing the average density values. Since no specific information is available on any of these points regarding the Soviet data, the degree of their influence on the issue could not be determined. Nevertheless, the discrepancies were large enough to prompt further investigation of possible causes.

Soviet observational errors

Subsequent scrutiny of available translated Soviet literature revealed the following significant quotation in a paper by Kolesov (1979):

A paper by Zmieva and Subbotin has appeared recently in which the results are reported of special measurements of the water equivalent of snow performed in the Medvenka River Basin for the purpose of determining the magnitude and reasons for errors in snow density determinations. It was found that systematic errors reaching 20-30 and even 50% occur when snow density is measured with standard density meters.

Since this disclosure was critical toward reaching a possible solution to the dilemma, a translated copy of the paper by Zmieva and Subbotin (1977) was retrieved. The results of their investigation on the accuracy of snow water equivalent with the standard Soviet densitometer are summarized in the following (quoted) four points:

1) Substantial systematic errors of the same sign (on the low side) reaching 20-30 and even 50%, rather than the 7-12% indicated in the Handbook, may arise in measurement of snow densities with the standard densitometer (especially new ones). The largest measurement errors are obtained at low air temperatures when the snow cover has a fine-grained structure with ice interlayers;

2) The instrumental error of determination of the maximum water equivalents when there is a wide air-temperature range on a given day can be reduced somewhat by making the snow survey in the latter half of the day instead of in the early morning after an overnight frost;

3) To lower the possibility of measurement errors due to clogging of densitometer by the column of snow, it is necessary to take the samples layer by layer, no more than 25-30 cm at a time;

4) In some cases, new tin plated densitometers may have the largest errors in snow-density determination because the inner walls are rougher than those of old, polished instruments. Attention should be given to the development of new methods and instruments for measurement of the density and water equivalents of snow cover.

Further evaluation of the information given in the text of the Zmieva-Subbotin paper disclosed the following:

1. When a single-step procedure was used to determine SCD with an earlier type standard densitometer, the measured values were 10 to 15% too low; if a layer by layer procedure was used with the earlier equipment the values were 20 to 30% too low.

2. Single step procedures with a new type of densitometer (one with a rougher inner surface) produced low readings that often exceeded 30%, especially on frosty days.

3. Early morning observations (i.e. the colder part of the day) resulted in readings that were 10 to 15% lower than the evening observations.

4. When measuring snow covers with crusty or icy layers, a controlled test consistently gave higher density values than those obtained from the routine densitometer method.

Data adjustment

It became obvious from these revelations that an adjustment to the SCD data obtained from the Soviet sources would be both necessary and justified. After considerable discussion and review of the facts and figures concerning the problem, it was decided that a correction factor of plus 20% on the data for all the stations would be warranted. Of course, several other adjustment schemes could have been adopted, such as a range in correction amount (e.g. from +15 to +25%). Unfortunately, the application of such an approach would require additional insight as to where and why one station would justify a higher or lower correction factor over another. Since no details on the observational techniques are available for any of the reporting stations, attempts to further refine the adjustment factor would be very subjective.

It should be noted that further inspection of the Soviet monthly SCD values (especially those given in Kopanev) revealed a few obvious unrepresentative values. For example, the station Chishmy reported values of 0.16, 0.21, 0.22, 0.23 and 0.18

Table 1. Average seasonal snow-cover density (g/cm³) for stations in the Soviet Union. Data based on snow surveys made during the last 10 days of the month for November through March in 1) open field areas, 2) forest clearings, and 3) under a forest canopy.

Station	Coordinates		Areas 1, 2 or 3 (as noted above)	Avg. seasonal snow-cover density* (g/cm ³)
	Lat(N)	Long(E)		
Aleksandrov	56°24'	38°43'	1	0.274
			2	0.258
Aleksandrovskoye	60 26	77 52	1	0.269
Amazar	53 52	120 53	1	0.180
			2	0.180
			3	0.168
Amga	60 53	132 00	1	0.206
Anuchino†	53 45	44 57	1	0.274
			2	0.247
Arkagala	63 09	146 47	1	0.204
Bakchar	57 01	82 05	1	0.257
Baunt	55 16	133 08	1	0.168
			2	0.154
Berezovo	63 56	65 03	1	0.245
			3	0.209
Birbidzhan	48 48	132 57	1	0.211
			3	0.194
Blagoveshchensk	50 16	127 32	1	0.218
			3	0.166
Bogandinskoye	56 54	65 53	1	0.242
			3	0.228
Bogotol	56 10	89 35	1	0.298
			3	0.266
Buchevaya	47 46	135 38	1	0.206
Bugulma	54 33	52 48	1	0.283
			3	0.223
Chelyabinsk	55 09	61 24	1	0.252
Chernigov	51 30	31 18	1	0.302
Chishmy	54 54	54 40	1	0.252
Danilov	58 12	40 10	1	0.266
			2	0.266
			3	0.245
Debal'tsevo	48 20	38 24	1	0.302
Dem'yanskoye	59 36	69 18	1	0.242
			2	0.223
Dolgiy Most	56 45	96 48	1	0.250
			3	0.226
Dudinka	69 24	86 15	1	0.283
Dzerzhinskoye	56 50	95 13	1	0.230
			2	0.221
			3	0.222
Dzhanky	69 46	135 04	1	0.182
			3	0.170
Gigant	46 30	41 20	1	0.290
Gor'kiy	56 20	43 59	1	0.283
			2	0.252
			3	0.242
Indiga	67 39	49 02	1	0.355
Kaluga	54 31	36 16	1	0.290
			3	0.245
Kamenka	58 33	95 51	1	0.254
			2	0.233
Karaganda	49 48	73 08	1	0.290

* Data obtained from Soviet sources and adjusted by +20% (see text).

† Same station name given for two different locations (see Appendix C).

Table 1 (cont'd).

Station	Coordinates		Areas 1, 2 or 3 (as noted above)	Avg. seasonal snow-cover density* (g/cm ³)
	Lat(N)	Long(E)		
Saskylakh	71 58	114 05	1	0.240
Semenovka	52 10	32 35	1	0.306
Serpukhov	54 55	37 25	1	0.283
Shadrinsk	56 05	63 38	1	0.298
			2	0.252
			3	0.240
Shugozero	59 55	34 12	1	0.254
			2	0.238
Solikamsk	59 39	56 47	1	0.293
			2	0.247
Sretensk	52 15	117 43	1	0.211
Staritsa	58 11	80 40	1	0.238
			3	0.209
Sterlitimak	53 37	55 58	1	0.278
Sukhobuzimskoye	56 30	93 16	1	0.266
Surgut	61 15	73 30	1	0.259
			2	0.233
			3	0.209
Sytomino	61 17	71 18	1	0.259
			3	0.211
Tobol'sk	58 09	68 11	1	0.259
			2	0.247
			3	0.216
Tommot	58 58	126 19	1	0.194
			3	0.192
Troitsk †	54 06	61 35	1	0.252
Troitsko-Pechorskoye	62 42	56 12	1	0.244
			3	0.242
Tugulym	57 05	64 33	1	0.254
Turukhansk	65 49	87 59	1	0.254
			2	0.245
			3	0.240
Tyumen'	57 09	65 30	1	0.250
			2	0.240
			3	0.204
Uglich	57 32	38 19	1	0.290
			2	0.242
			3	0.238
Ussuriysk	43 48	131 59	1	0.243
			2	0.209
Ust'-Kulom	61 42	53 40	1	0.259
Valday	57 59	33 16	1	0.305
Vereb'ye	58 41	32 42	1	0.264
			3	0.242
Vereshchagino †	58 05	54 40	1	0.242
			2	0.238
Verkhoyansk	67 33	133 23	1	0.173
Voznesen'ye	61 01	35 29	1	0.242
Vyazniki	56 15	42 10	1	0.295
Yakovlevka	44 25	133 29	1	0.238
			3	0.206
Yar-sale	66 50	70 50	1	0.317
Yessey	68 29	102 10	1	0.226
Zryanka	65 44	150 54	1	0.218
			3	0.211

* Data obtained from Soviet sources and adjusted by +20% (see text).

† Same station name given for two different locations (see Appendix C).

g/cm³ for November, December, January, February and March, respectively. Obviously the SCD of 0.18 g/cm³ for March was low, especially since 0.29 g/cm³ was observed in April as thawing began. The March value, consequently, was considered to be closer to 0.23 g/cm³ which increased the *unadjusted* average seasonal density at Chishmy to 0.210 instead of 0.200 g/cm³. Similar amendments were limited in number, and in almost all cases the average SCD did not change by more than ± 0.01 g/cm³.

The correction factor of plus 20%, and the few other minor adjustments, were then applied to the SCD's reported by the 105 Soviet stations and the results are shown in Table 1. The stations in the tabulation are listed alphabetically, with coordinates to the degree and nearest minute. Note also that many stations conducted density measurements in forest clearings and under forest

canopies. These SCD values were adjusted in a similar fashion to those taken in the open field and included in Table 1. A comparison study between these SCD measurements made at locations of different exposure is presented later in *Seasonal, Monthly and Local Density Variations*.

DEVELOPMENT OF THE SOVIET SNOW-COVER DENSITY CHART

Previous work

One of the main objectives of this study was to construct a map that shows probable distribution of SCD across the Soviet Union. The idea evolved following the development of such a map for northern North America (Bilello 1967).

In the 1967 investigation, a relationship between climate and the regional variations in SCD was de-

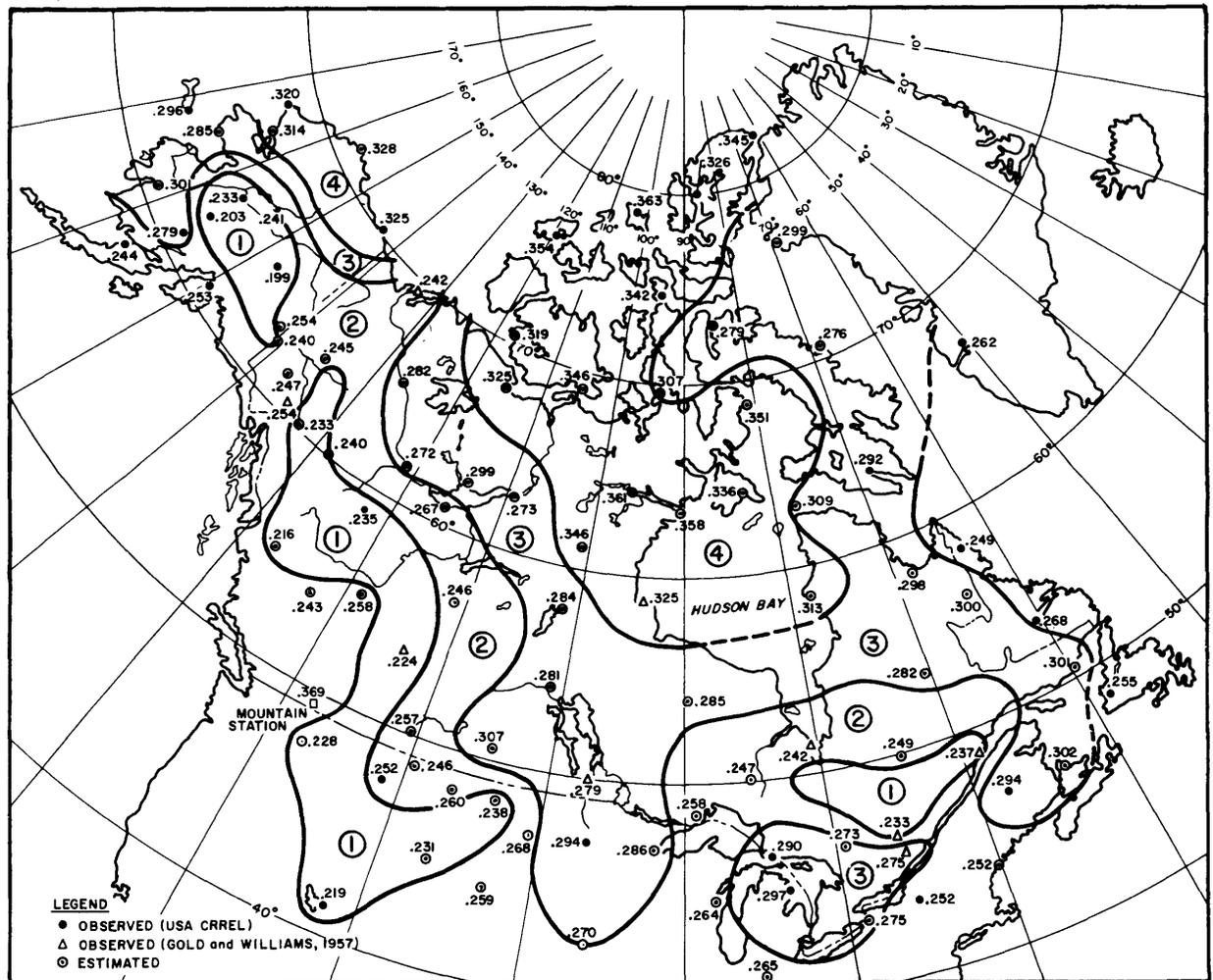


Figure 1. Average seasonal snow-cover densities for northern North America. Categories 1-4 are separated at densities of 0.24, 0.27 and 0.31 g/cm³.

veloped using data collected from a systematic observation program at stations in Canada, Alaska and the northern conterminous United States. The results of the study showed that the stations could be placed into one of four SCD categories. The density limits for these four categories were defined as $<0.24 \text{ g/cm}^3$ for category 1, ≥ 0.24 to $<0.27 \text{ g/cm}^3$ for category 2, ≥ 0.27 to <0.31 for category 3, and $\geq 0.31 \text{ g/cm}^3$ for category 4. These groupings were based on observed densities, climatic factors, and to some extent on geographic considerations. The main reason for the grouping concept is to emphasize that the study attempts to define regional variations rather than point estimates of SCD.

A multiple regression analysis in which seasonal SCD is related to concurrent air temperature and wind speed observations was also performed in the 1967 study. The resultant equation derived from 2500 density measurements obtained from 2 to 11 years of record at 27 stations was

$$\rho = 0.152 - 0.0031T + 0.019W \quad (1)$$

where ρ = average seasonal SCD (g/cm^3)

T = average seasonal air temperature
($^{\circ}\text{C}$)

W = average seasonal wind speed (m/s).

The correlation coefficient and the error of estimate in the relationship were 0.84, and 0.025 g/cm^3 respectively. Interestingly, the numerical constant (0.152) appearing in eq 1 is almost identical to that given by Dmitrieva (1950) for the average density of fresh snow (i.e. 0.15 g/cm^3).

Climatological data for 61 other stations across northern North America were then compiled and average SCD values were estimated from the equation. These values, plus the data from the 27 basic stations and from 10 other locations given in Gold and Williams (1957), were used to draw an average SCD map for North America (Fig. 1). In this figure the densities across the continent are divided into the four categories described earlier and are separated at 0.24, 0.27 and 0.31 g/cm^3 .

Relationship between Soviet climate and SCD regional variations

A test was conducted to determine the possible application of eq 1 for developing the SCD map for the Soviet Union. In order to do this, air temperature and wind speed records were required for Soviet stations where SCD data were available. A comprehensive search and retrieval of monthly weather records for stations in the U.S.S.R. was

therefore conducted. The required data consisted of long-term monthly summaries of observed air temperature and wind speed data for the period November through March. This portion of the winter was used so that the selected "seasonal" SCD period would be similar for both studies. The monthly climatic summaries were obtained from four principal sources: U.S. Air Force (1948-1971), U.S. Department of Commerce (1951-1960), U.S. Air Force (1965-1968), and Nuttinson (1950).

Climatic information for over 500 stations in the U.S.S.R. was obtained from these sources, and used for various purposes in this report, including its possible use in the relationship given in eq 1. To test the relationship, 41 Soviet stations with available SCD information and corresponding air temperature and wind data were used in the analysis. The adjusted SCD values given in Table 1 were used in these tests. Unfortunately, the results revealed that eq 1 provided estimated Soviet densities for a majority of stations that could not be realistically accepted. Of the 28 stations used in the test, five estimated densities were low (by -1% to -6%), and one value matched exactly, but 22 estimates were much too high—ranging from +2% to +33%, with an average of +10%.

Since the climatic index of combined temperature and wind speed gave unsuitable Soviet SCD estimates, separate tests on each of the climatic parameters were conducted. When average seasonal air temperatures alone were used to estimate the densities, the correlation proved to be extremely poor. The scatter of points that evolved when the data sets were plotted indicated that further statistical evaluation in the relationship would be fruitless. This incompatibility in the temperature relationship between the two continents may partially result from greater length of the winter in the Soviet Union, especially in central Siberia, than across northern North America. The average seasonal (i.e. November through March) air temperatures between the two regions would not be similar and therefore not comparable.

The test on the relationship between average wind speed (for November through March) and the adjusted seasonal SCD for the Soviet locations proved, however, to be quite successful. Weinberg and Gorlenko (1940) also note that the densification effect by wind on the snow cover is of significant importance. Wind fragments the snow crystals, causes the finer grains to re-sort and increases the density of the snow cover by packing. Data sets for the 41 test stations are presented in Table 2, and plotted in Figure 2. The stations in

Table 2. Average seasonal adjusted snow-cover density vs associated average seasonal (November through March) wind speed for 41 Soviet field sites.

<i>Station*</i>	<i>Density category</i>	<i>Snow-cover density (g/cm³)</i>	<i>Wind speed (m/s)</i>
Oymyakon	1A	0.173	0.4
Verkhoyansk	1A	0.173	0.6
Krasnyy Chikoy	1A	0.192	1.0
Kedon	1A	0.194	1.1
Zyryanka	1B	0.218	2.0
Blagoveshchensk	1B	0.218	2.8
Kirovskiy	1B	0.221	2.0
Yessey	1B	0.226	2.2
Dzerzhinskoye	1B	0.230	2.3
Saskylakh	2	0.240	3.4
Krasnoyarsk	2	0.240	3.7
Troitsko-Pechorskoye	2	0.240	3.3
Berezovo	2	0.245	3.4
Kolpashevo	2	0.245	3.7
Dolgiy Most	2	0.250	2.4
Tyumen'	2	0.250	4.7
Chelyabinsk	2	0.252	4.2
Kamenka	2	0.254	2.4
Turukhansk	2	0.254	3.9
Bakchar	2	0.257	3.8
Tobol'sk	2	0.259	4.1
Surgut	2	0.259	4.7
Kezhma	2	0.262	2.5
Kargasok	2	0.262	4.3
Leushi	2	0.264	3.8
Petropavlovsk	2	0.264	5.2
Sukhobuzimskoye	2	0.266	3.4
Aleksandrovszkoye	2	0.269	3.4
Kurgan	2	0.269	4.3
Perm	2	0.269	5.0
Kazan'	3	0.271	5.3
Khosedda-Khard	3	0.274	4.4
Dudinka	3	0.283	5.3
Gor'kiy	3	0.283	5.8
Kustanay	3	0.288	4.3
Karaganda	3	0.290	4.9
Bogotol	3	0.298	4.8
Kokchetav	3	0.298	5.8
Saratov	3	0.300	5.3
Noril'sk	3	0.300	6.1
Novyy Port	4	0.322	7.3

* Station coordinates and elevation (if available) are given in App. C.

Table 2 are listed according to increasing adjusted seasonal SCD starting with the lowest and ending with the highest value. This was done for two reasons: 1) to emphasize the marked association between increasing densities and increasing wind speeds, and 2) to show how groups of stations were placed within the density categories described earlier.

At this point, it should be noted that category 1 for the Soviet data was divided into two parts because 29 of the 105 Soviet field stations reported adjusted values of $<0.24 \text{ g/cm}^3$ and almost half of these stations (13 out of 29) reported values less than 0.21 g/cm^3 . Consequently category 1A contained values $<0.21 \text{ g/cm}^3$ and category 1B included values ≥ 0.21 to $<0.24 \text{ g/cm}^3$. As will be

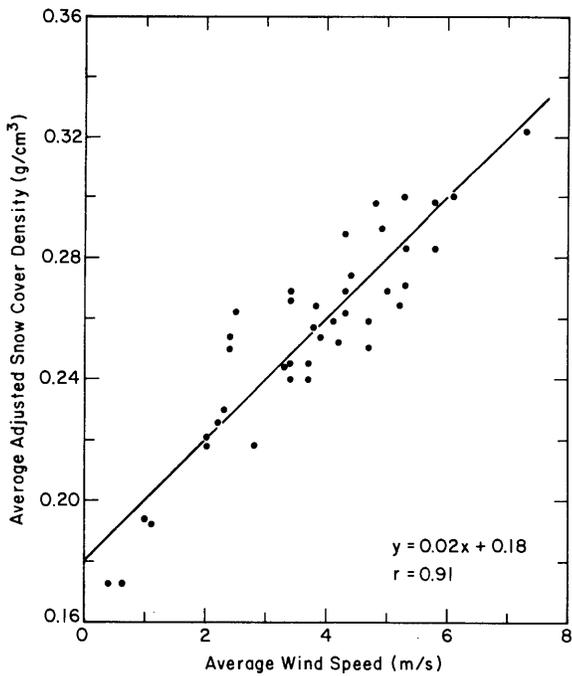


Figure 2. Average five month (Nov-Mar) snow-cover density (g/cm^3) (field site data-adjusted by +20%) vs average 5 month (Nov-Mar) observed wind speed records (m/s), for 41 field sites in the U.S.S.R.

seen later, a major portion of the Soviet Union experiencing light winter winds is included in category 1A.

A regression analysis of the data shown in Figure 2 was conducted, and the resultant equation is

$$q = 0.181 + 0.0197W$$

or

$$q = 0.181 + 0.020W \quad (2)$$

where q is the average seasonal SCD (adjusted by +20%) in grams per cubic centimeter and W is the average seasonal (November through March) wind speed in meters per second. The correlation coefficient and the standard error of estimate in the relationship are 0.91 and 0.014 g/cm^3 respectively.

PREPARATION OF THE SNOW-DENSITY CHART

Selection of base map for plotting purposes

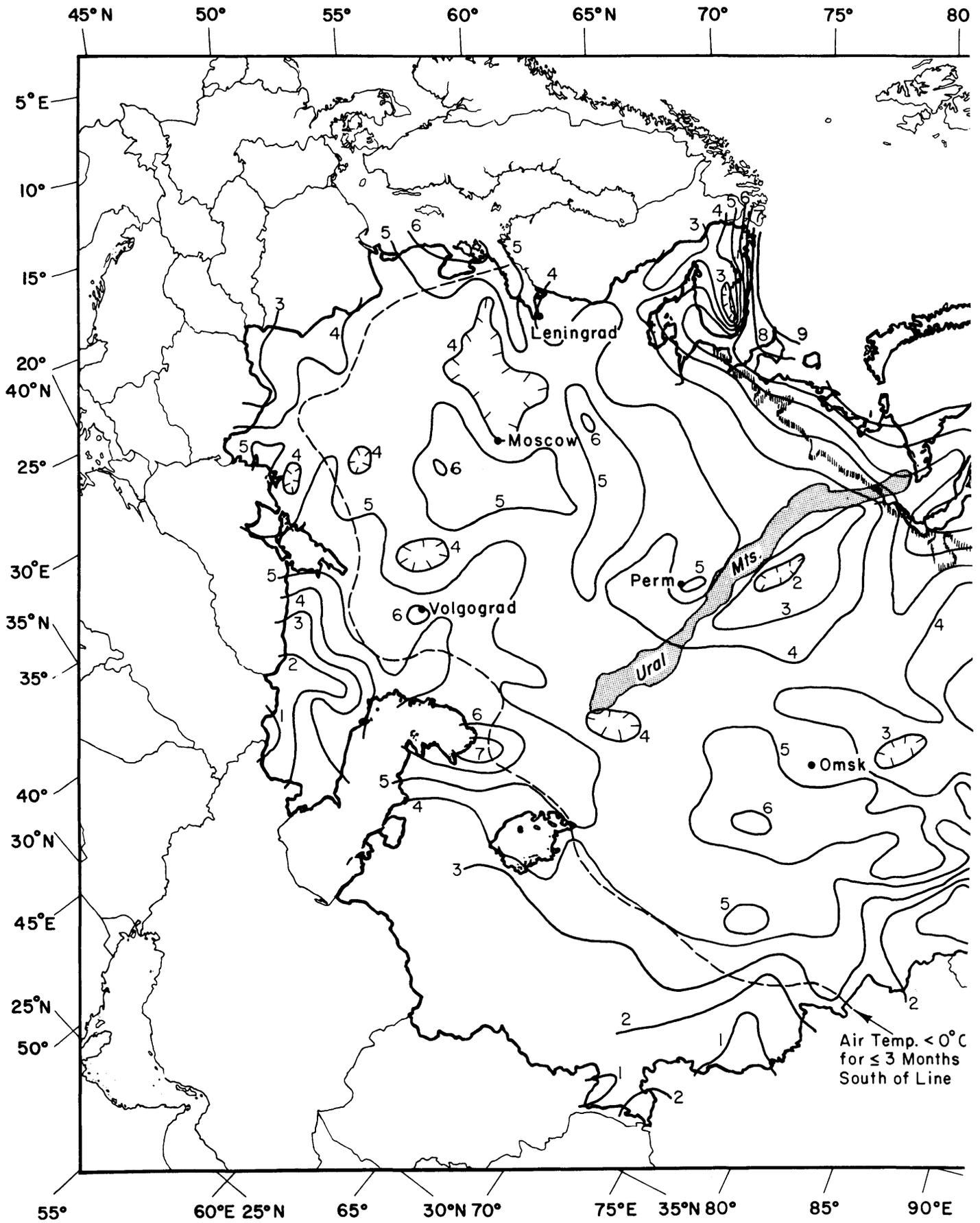
The National Geographic Map of the Soviet Union (National Geographic Society 1976) provid-

ed excellent station location coverage, and its scale was sufficiently large to accomplish all the plotting requirements. The base chart drawn for this study was constructed from this map and included 1) U.S.S.R. boundary lines, 2) coordinate identifiers, 3) the location of "the northern limits of wooded country," and 4) the location of a few major cities. Other information such as the outline of the Ural mountains, and a line demarcating the region in southwest Russia that records average freezing air temperatures during three or less months of the year were obtained from the U.S. Army (1962 and 1964), and U.S. Air Force (1948-1971) respectively.

Soviet snow-cover density chart

The initial step in the construction of a chart on the regional distribution of SCD across U.S.S.R. was to plot the adjusted densities determined for the 105 Soviet stations listed in Table 1 on the base map. Although the number of Soviet stations was almost three times as great as that available for the North American study (i.e. 37 sites), the coverage for the U.S.S.R. is still rather sparse. To supplement this primary data base, the relationship between wind speed and SCD shown in eq 2 was utilized. To accomplish this objective, average "seasonal" wind speed values for the more than 500 climatic stations (referred to earlier) across the Soviet Union were compiled (App. C). This alphabetical tabulation lists the station coordinates (to the degree and nearest minute), the average seasonal (November through March) wind speed to the nearest tenth (in meters per second), and if available, the station's elevation (meters above sea level), and the number of months per year with air temperatures of less than 0°C .

This information was then used to construct a wind chart for the Soviet Union based on average wind speed recorded from November through March (Fig. 3). Considerable confidence in the accuracy of this Soviet wind chart was gained because of the large number of data points available for the analysis. It was also noted that the average wind speed recorded during each of the months was generally quite uniform, so that the variability for this climatic parameter was minimal. Although Figure 3 was developed to provide additional guidance toward the construction of a SCD chart, it also presented some unusual Soviet wind speed conditions. For example, very strong wind speeds are observed along parts of the northern and eastern coasts of the U.S.S.R. The average mid-winter winds at some of these areas exceed 9 m/s. In contrast, very light winds of less than 2



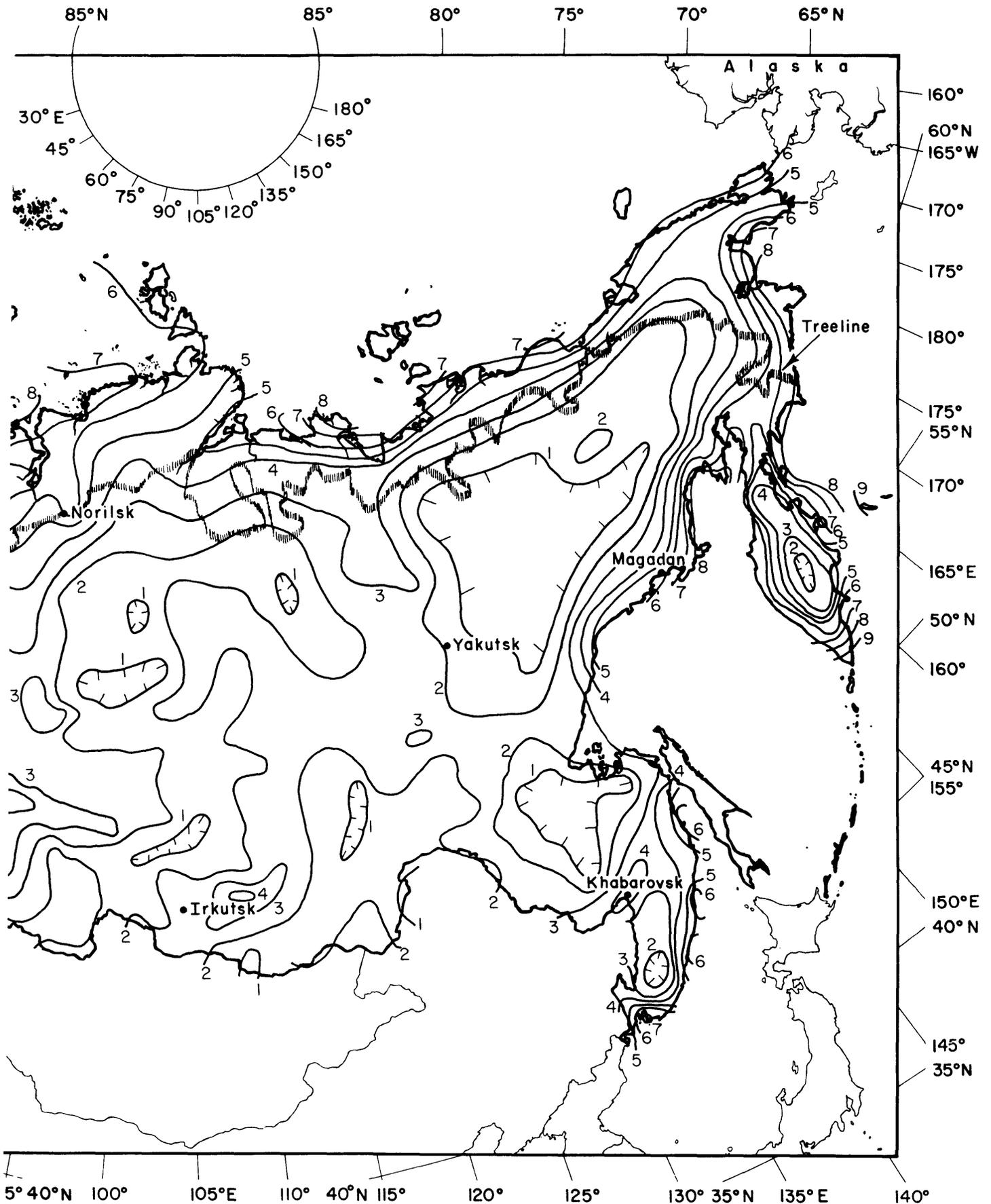
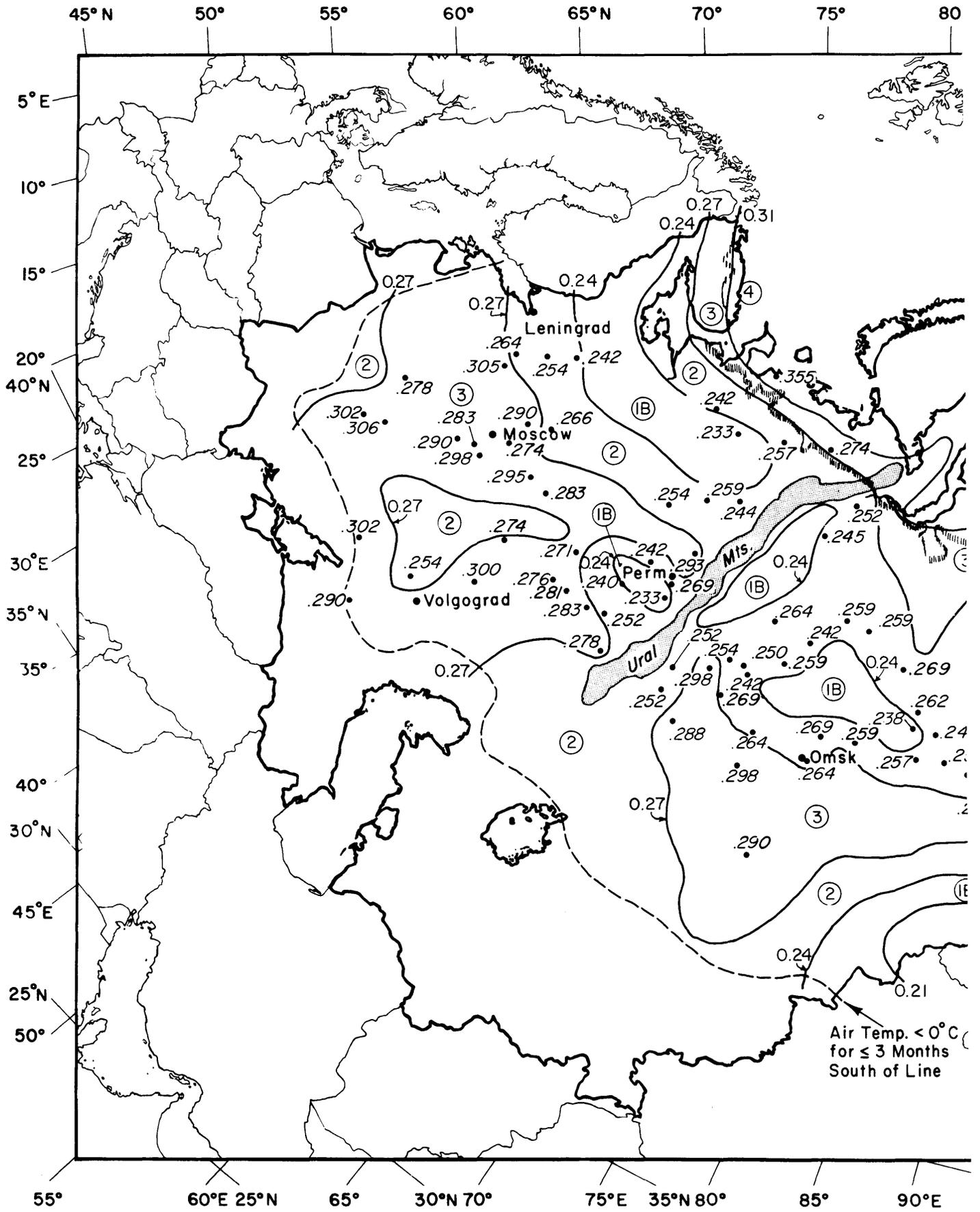


Figure 3. Isolines of average winter (November–March) wind speed (m/s) for the U.S.S.R.



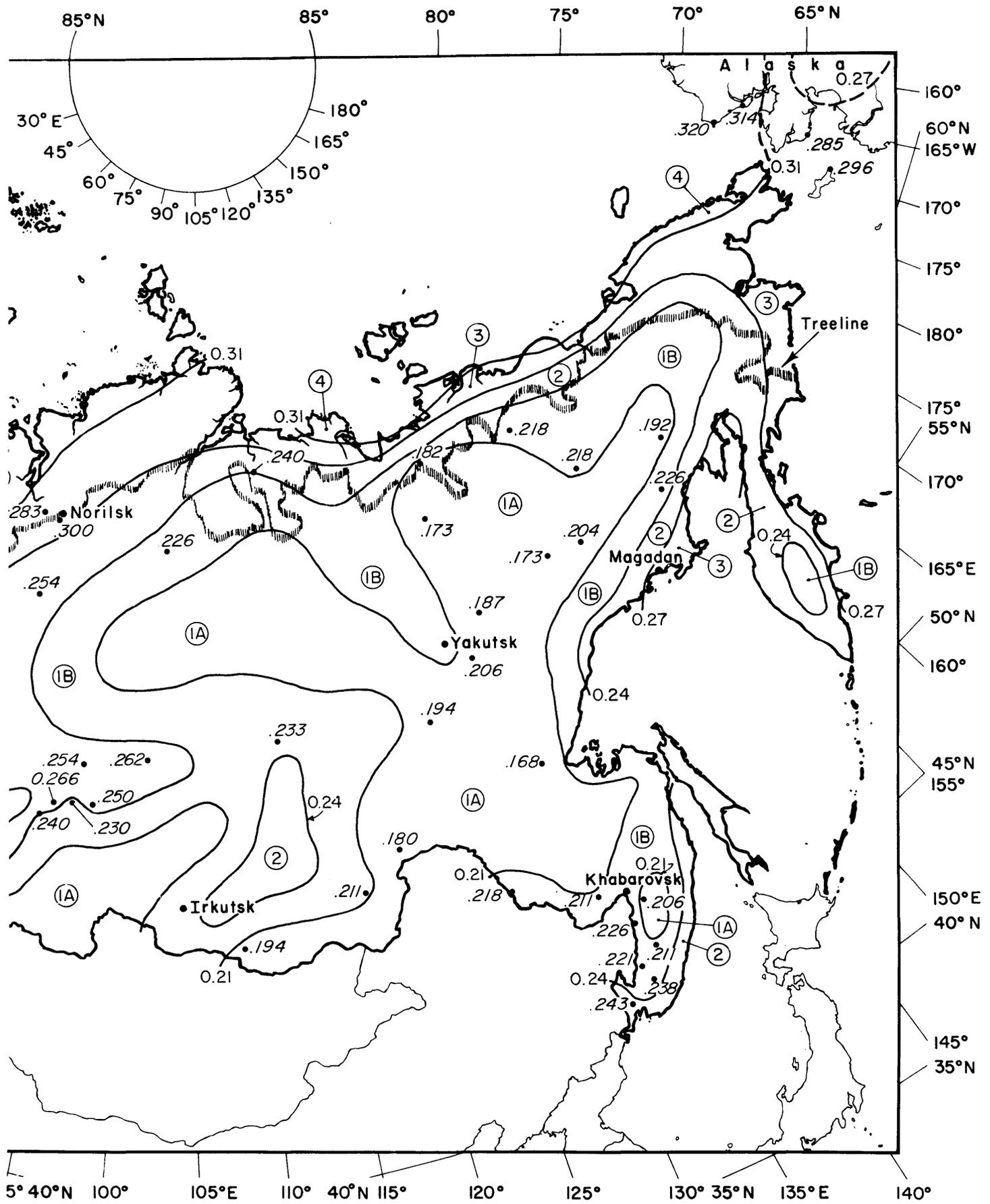


Figure 4. Average seasonal snow-cover densities for the U.S.S.R. Categories 1A, 1B and 2-4 are separated at densities of 0.21, 0.24, 0.27 and 0.31 g/cm³.

m/s on the average are recorded in winter throughout the interior region of eastern and central Siberia. Some regional wind speed differences due to land/water contrasts and mountain barriers (e.g. the Urals) were also noted.

The five SCD zones, separated at density values of 0.21, 0.24, 0.27 and 0.31 g/cm³, correspond almost exactly with seasonal wind speeds of 1.5, 3.0, 4.5 and 6.5 m/s as shown in eq 2 and Figure 2. By using the position of the contours that locate these particular wind speeds, the isoline analysis presented in Figure 3 would therefore be useful in the development of the density chart.

The application of this relationship, in combination with the plot of the 105 adjusted SCD values, produced the final Soviet Union SCD map (Fig. 4). The country was divided into the five selected categories, ranging from the very light snow cover (category 1A) in the sheltered regions of central Siberia to densely packed snow (category 4) across most of the arctic coastline. In areas void of data points, the density boundaries were determined not only by wind speed but by other physical features such as topography and the northern limit of tree growth. Incidentally, no attempt was made to evaluate the variations in SCD that occur in mountainous regions. In fact, since all but 15 of the almost 350 stations that provided elevation data were located below 800 m, estimated SCD's for locations above this height should not be considered as applicable in this study.

Lipovskaya (1966) produced a chart that shows average density of snow cover for the Soviet Union (Fig. A2). This chart was prepared in accordance with the following quoted information:

Multiyear data, as well as data calculated from annual observations based on snow surveys and constant snow-stakes, served as the basic material for plotting the chart of density occurring in the decade of deepest snow.

The results shown in Figure A2 were not in any manner included in the analysis used in Figure 4. Nevertheless, some marked similarities and one major difference between the two SCD charts were found. For example, the regions of highest densities in both studies occurred along the northern fringes and eastern edges of the U.S.S.R. Very light snow covers were found in the eastern interior (Siberian) sections, and on the sheltered (eastern) side of the Ural Mountains. Considerable regional variations in density were also found in the western sections including the European portion of the country in both studies. The major (and important) difference in the studies is, however, the actual density values shown in the figures. It is apparent that Lipovskaya used snow survey data col-

lected prior to 1966 when the densitometers in use were providing incorrect readings. Consequently, the SCD isolines shown in Figure A2 are generally about 20% lower than those described by the regional categories presented in Figure 4. Although the similarities in the general distribution of SCD throughout the nation in both (independently conducted) studies are noteworthy, the results provided in Figure 4 are believed to be more accurate.

SEASONAL, MONTHLY, AND LOCAL DENSITY VARIATIONS

The information provided in Figure 4 can be considered limited in that it only provides average regional SCD values. Therefore three questions accompany the subsequent use of such data:

1. Within each of the selected categories, what range of SCD's would be expected when all the values observed throughout the period of record are considered?
2. What variations in SCD can be expected on a month-to-month basis within each of the categories?
3. What variations in SCD were observed in areas of different exposure?

Examination of the SCD data received from the 105 Soviet stations, and results obtained from the previous SCD study for North America provided additional information that address some of the relevant questions.

Frequency distribution curves

Although a specific range in densities was selected to define each of the SCD categories, individual density measurements made from November through March would, of course, vary above and below the established limits. In order to show such a distribution of measurement points, a series of frequency curves based on all observed densities are required. Unfortunately, such curves could not be developed since only average monthly SCD values based on all the years of record are available for the Soviet data. A good indication of the probable distribution may, however, be derived from results obtained in the North American report (Bilello 1967). In that study, frequency distribution curves were developed for categories 1 through 4 (Fig. 5). Except for category 1, the SCD limits for the categories given in both studies are similar. The need to separate category 1 (Fig. 5) into 1A and 1B to conform with the Soviet classification in this case did not seem warranted. The

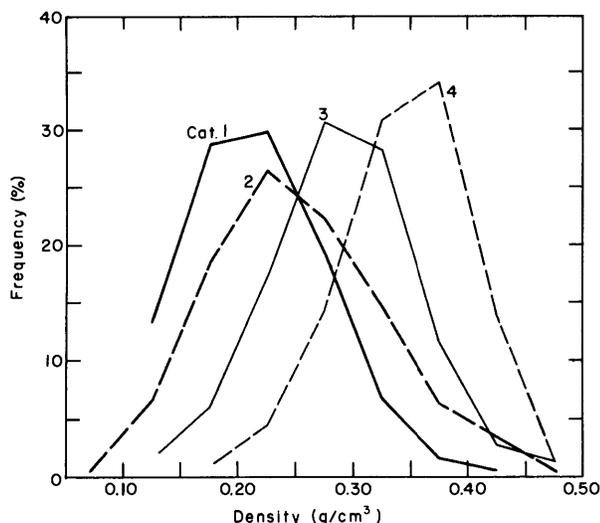


Figure 5. Frequency curves for density categories 1-4. Curves comprise line segments connecting midpoints of class intervals 0.10-0.15 g/cm³, 0.15-0.20 g/cm³, etc. Curves refer to data obtained in North America study (Bilello 1967).

distribution shown in category 1 for North America can instead be associated with the Soviet data in category 1B because the mean SCD values in these groups are quite close (i.e. 0.214 g/cm³ for North America, and 0.224 g/cm³ for the Soviet Union).

The frequency curves in Figure 5 show how each of the categories fall within a series of increasing densities. Naturally, the local topography and vegetation create differences in density from point to point within a categorically defined region. Deviations from any of the computed averages for each station can be expected from month to month and year to year. These distribution curves provide an indication of the extent of these areal, seasonal and annual variations in SCD.

Monthly increase in density

The metamorphic processes which take place within a snow cover with time have been described by a number of investigators (e.g. Bader et al. 1939, de Quervain 1945 and Kingery 1960). It has been shown (e.g., Shepelevsky 1938, Work 1948 and Gold 1958) that this aging process results in a gradual increase in the snow-cover density as the winter progresses.

This time-densification process was investigated using the average adjusted monthly density values available from the Soviet sources and compared with results obtained from the North American

Table 3. Comparison of average monthly snow-cover density (SCD) values (g/m³) between U.S.S.R and North American stations.

Category	U.S.S.R. average adjusted SCD's				
	Nov	Dec	Jan	Feb	Mar
1A	160	176	186	202	220
1B	191	206	221	234	269
2	215	238	255	268	296
3	223	252	288	314	356
4	288	308	340	356	364

Category	North American average SCD's				
	Nov	Dec	Jan	Feb	Mar
1	179	197	212	234	237
2	238	226	254	263	280
3	268	277	291	294	311
4	317	343	342	348	345

study. The monthly values for all stations within each of the categories were combined and the computed averages for both studies are presented in Table 3.

The density for the Soviet stations in categories 1A, 1B, 2 and 3 (except from February to March for category 3) increases from about 0.15 to 0.20 g/cm³ each month, whereas the density for the North America stations in categories 1, 2 and 3 (except from November to January for category 2) increases by approximately 0.10 g/cm³ each month. The stations in category 4 in both regions show much less change throughout the winter. This uniformity may result because their densities are initially quite high, and because they are located at latitudes where the effects derived from solar radiation toward snow-cover compaction during mid-winter are negligible.

Density variations due to exposure

Of the 105 Soviet stations that conducted SCD observation in open field areas (Table 1), 29 also made concurrent measurements in forest clearings (i.e. partially exposed areas), and 37 also made measurements under a forest canopy (i.e. protected areas). An inspection of the SCD variations that one might expect due to these environmental differences was made. The test was conducted by comparing each of the Soviet average seasonal adjusted *field* SCD's with 1) the associated value obtained at the forest clearing site, and 2) the associated value obtained under a forest canopy. Plots of these data sets are shown in Figures 6 and 7 respectively.

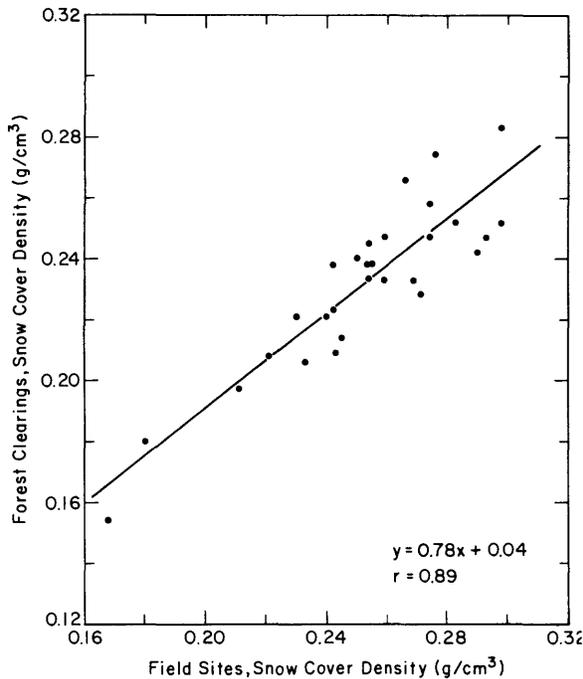


Figure 6. Comparison of U.S.S.R. average seasonal adjusted snow-cover density (g/cm^3) at open field sites vs at forest clearings.

A line of best fit in both figures was determined in a regression analysis of the data points. The results presented in Figure 6 show that average seasonal SCD measurements made at forest clearings range from about 4% (for category 1A locations) to nearly 10% (for category 3 locations) less than those made in open field areas. The results in Figure 7 show that measurements made under a forest canopy range from about 8% (for category 1A locations) to near 14% (for category 3 locations) less than those made in open field areas.

DISCUSSION

The results presented in this study provide an estimate of the average snow-cover density (SCD) regionally across the U.S.S.R. The study does not attempt to predict the density from point to point, from week to week or year to year. Included in the SCD map developed for the country were observed density values for 105 stations that were adjusted by +20%. This adjustment was based on recent Soviet literature that revealed systematic densitometer measurement errors resulting in low readings. This information, combined with a relationship found between wind speed and density, was then used to divide the U.S.S.R. into geographical areas that fall into one of five pre-

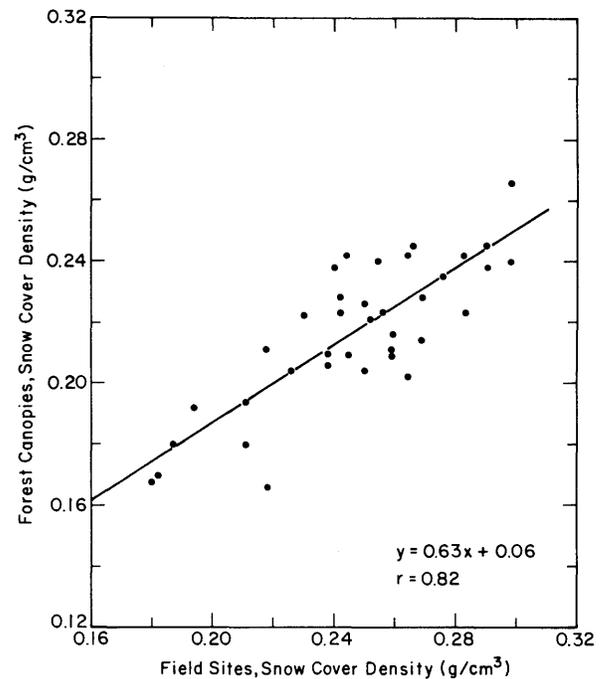


Figure 7. Comparison of U.S.S.R. average seasonal adjusted snow-cover density (g/cm^3) at open field sites vs under forest canopies.

defined ranges of density. Those regions with the least dense snow cover ($<0.21 \text{ g}/\text{cm}^3$) included the interior sections of the Soviet Union with light "winter" winds. The region with the greatest snow density ($\geq 0.31 \text{ g}/\text{cm}^3$) included the northern arctic coastal region and the northeastern fringe zones that are above the tree line. These areas are very exposed and subject to moderate to strong winds.

Within these regional ranges the snow cover and its characteristics can, however, be quite variable. Consequently, results of further studies on the subject provided some insight on the seasonal, monthly, and local variations that one might expect within the categories. For example, frequency deviations showed that, within each category, standard deviations (in density) of about $\pm 0.06 \text{ g}/\text{cm}^3$ can be expected. Monthly SCD values were also found to increase gradually from November to March and were higher in open field areas than under a forest canopy. Attempts to evaluate the variations in SCD in mountainous regions were not made. Snowpacks in areas of complex terrain range from the permanent snowfields at higher elevations to an occasional, brief snow cover in the warmer valleys (Tsomaia 1956). Due to the relatively small scale of the snow-cover density base map and the limited number of observed data, further delineation of the selected categories

was not possible. In addition caution should be used when interpreting Figure 4 for regions above 800 m, since the vast majority of the Soviet stations with observed densities and/or wind speed data are located below this elevation. Nevertheless the results presented in this study would be applicable of areas where the topography, vegetation and wind speed are reasonably uniform.

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APPENDIX A

COMPILATION OF SNOW-COVER DENSITY INFORMATION EXTRACTED FROM VARIOUS SOVIET ARTICLES

In the review of the literature on snow cover characteristics, 28 Russian articles provided specific information on SCD in the Soviet Union. Some of these reports were previously available in English copy and the remaining papers were translated by CRREL. In some cases, the original articles were excessively lengthy and contained material irrelevant to this study. In these cases, only the pertinent sections of the articles were translated.

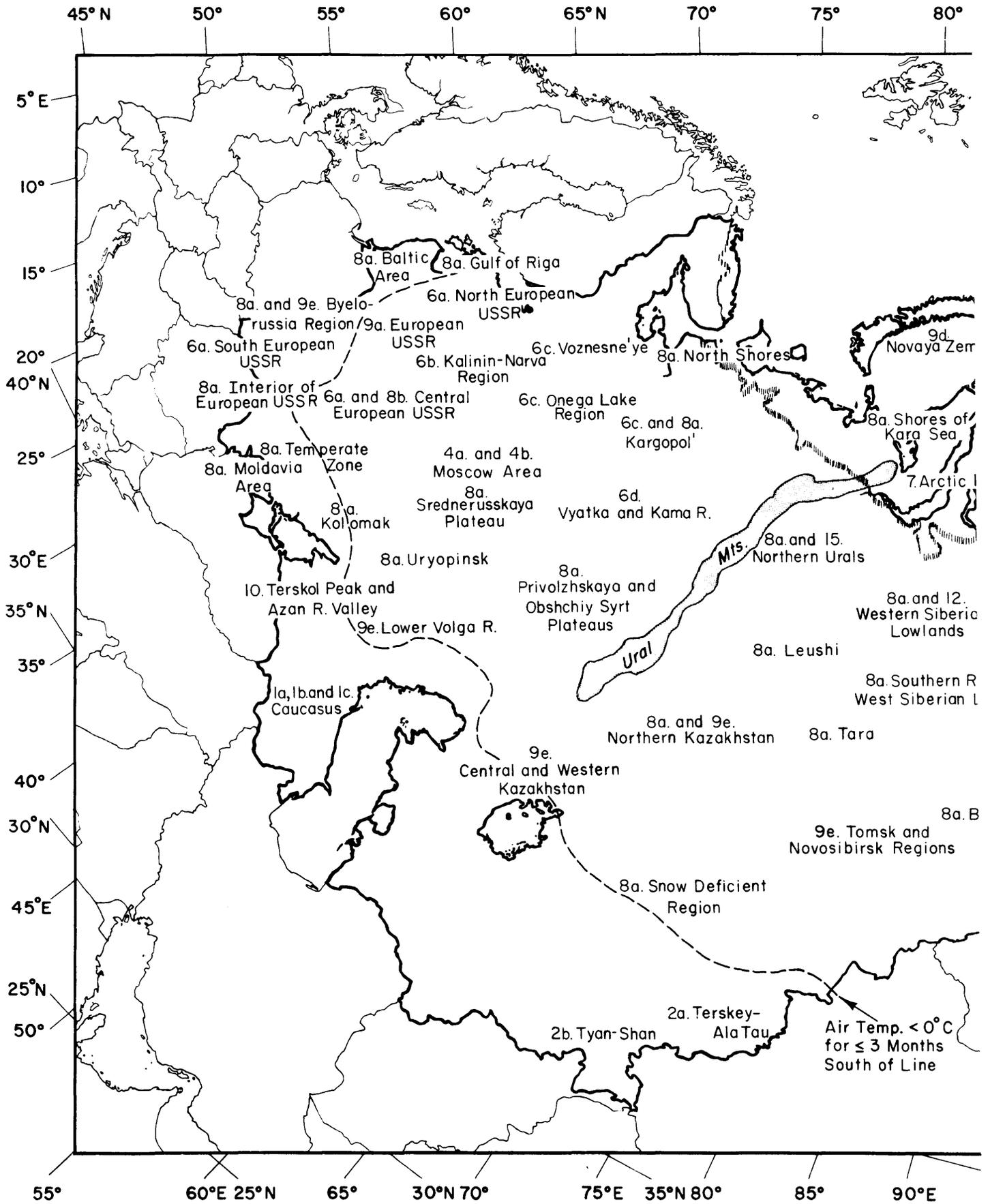
The snow density information that appeared germane to this report was extracted and included in this appendix. Although the bulk of the material was transcribed verbatim, numerous unnecessary sections, paragraphs, sentences, etc., were omitted from the original reports. Consequently, the listed summaries cannot be considered as fully completed quotations.

The information given in the 28 articles was placed into one of the following 15 general groups, most of which define geographical regions of the U.S.S.R.:

1. The Caucasus
2. Terskey - Ala Tau Mountain region
3. Southern part of the Buryat region
4. Moscow region
5. Amur Region
6. European part of the U.S.S.R.
7. Arctic Regions
8. Throughout the U.S.S.R.
9. Selected areas in the U.S.S.R.
10. Terskol Peak and Azau River Valley
11. Irkutsk
12. U.S.S.R. Taiga
13. U.S.S.R. Snow Surveys
14. Kamchatka
15. Ural Mountains

Some inconsistencies were noted in the articles. For example, the following statement is made in the Caucasus section (item 1a) "the conformity observed in the entire territory of snow cover density is lessening as the elevation above sea level increases." Whereas in the summary given under item 8a the statement, "In the mountains, it [i.e. snow-cover density] varies with the slope, orientation, snow-cover thickness and altitude above sea level. For example, the density increases in the Caucasian Mountains with altitude." Unless the statements are the results of incorrect translations, the observations are in direct conflict.

Despite the possible inherent erroneous information that may exist in the extracted material, excerpts of SCD values given in each of the articles were compiled and listed in Table A1. The general regions where these measurements were made are shown in Figure A1. This procedure makes it possible to associate the reported density values with the area to which they apply. Although the extracted SCD data cover many regions of the



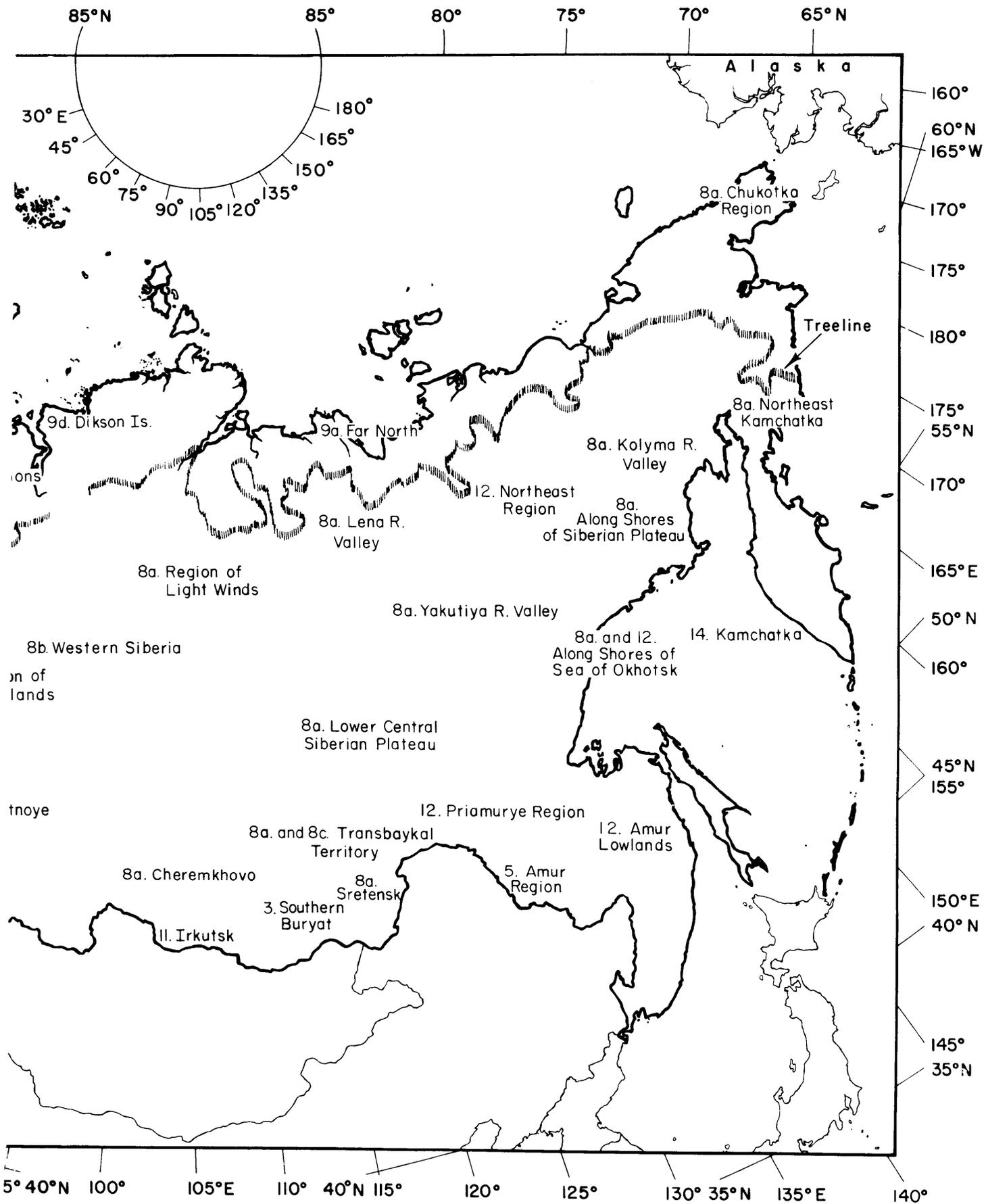


Figure A1. Location of geographical areas (or regions) described in Appendix A (see Table A1).

Table A1. Summaries of snow-cover density data from 28 Soviet references. Numbers correspond to sections in this appendix and locations in Figure A1.

*Excerpts of Snow-Cover Density Data from
28 Soviet References*

1a, 1b and 1c. The Caucasus

From 0.10 to 0.40 g/cm³ depending on month, elevation, slope, and exposure. Average value from November through April is less than 0.20 g/cm³. Density increases to 0.30 and 0.40 g/cm³ during thaw.

2a. Terskey-Ala Tau

Density increases from 0.20 to 0.25 g/cm³ between November and April.

2b. Tyan-Shan

During spring months, density ranges from 0.28 to 0.49 g/cm³.

3. Southern Buryat

During March 1956 density in steppes and forest-steppes was 0.13 to 0.20 g/cm³, with higher values (0.17 to 0.27 g/cm³) in drifted areas.

4a and 4b. Moscow region

Densities of 0.13 to 0.22 g/cm³ during November, and 0.15 to 0.20 g/cm³ in December. During thaw, values increase to between 0.22 and 0.34 g/cm³, and from 0.35 to 0.40 g/cm³ in drifts.

5. Amur region

Average seasonal densities of between 0.25 and 0.29 g/cm³. In March they range from 0.11 to 0.40 g/cm³.

6a. South European U.S.S.R.

At beginning of spring, thaw density is 0.22 or 0.23 g/cm³, but it is as high as 0.34 to 0.36 g/cm³ in open areas.

6a. North European U.S.S.R.

At beginning of spring thaw, density ranges from 0.22 to 0.28 g/cm³.

6a and 8b. Central European U.S.S.R.

At beginning of spring thaw, density ranges from 0.24 to 0.32 g/cm³. Values during March and April increase to 0.28 and 0.40 g/cm³.

6b. Kalinin-Narva Region

At the time of maximum water equivalent density ranges from 0.22 to 0.28 g/cm³.

6c. Onega Lake Region

Density in January is 0.20 to 0.28 g/cm³, and increases to 0.38 by April.

6c and 8a. Kargopol'

From November through March, density ranges from 0.13 to 0.26 g/cm³, with average values of 0.22 in open areas and 0.21 g/cm³ in protected areas.

6d. Vyatka and Kama Rivers

At time of maximum equivalent, density ranges from 0.26 to 0.32 g/cm³ in field areas, and 0.24 to 0.28 in forests.

7. Arctic regions

Density averages 0.30 to 0.35 g/cm³ all winter. It increases to 0.40 and 0.45 g/cm³ at the beginning of thaw.

8a. Baltic region

Average density less than 0.24 g/cm³.

8a and 9e. Byelorussia region

Average density is less than 0.24 g/cm³. In forests it ranges from 0.19 to 0.22 g/cm³, and in fields it reaches 0.25 g/cm³.

8a. Interior of European U.S.S.R.

Average density mostly near 0.24 g/cm³.

8a. Temperate zone

Density in forested areas 0.20 to 0.23 g/cm³, and 0.25 to 0.27 g/cm³ in unshielded areas.

8a. Moldavia region

Average density less than 0.24 g/cm³.

8a. Kolomak

Average density of 0.26 g/cm³ in both open and protected areas.

Srednerusskaya Plateau

Average density of 0.28 g/cm³.

8a. Uryopinsk

Average density of 0.27 and 0.28 g/cm³ in both open and protected areas.

8a. Privolzhskaya and Obshchii Syrt Plateaus

Average density of 0.28 g/cm³ with lower values in the river valleys.

8a and 9e. Northern Kazakhstan

At higher elevations density ranged between 0.27 and 0.30 g/cm³.

8a. Snow-deficient region

Density ranges from 0.14 to 0.16 g/cm³.

8a. North shores

Average density is 0.30 g/cm³.

8a and 15. Northern Urals

Average density exceeds 0.28 g/cm³, and reaches 0.35 and 0.42 g/cm³ in second half of winter. In forest zones values range from 0.25 to 0.30 g/cm³ in upper third of snow pack and 0.15 to 0.28 g/cm³ in lower layers. In unforested zones wind slab densities are 0.42 to 0.45 g/cm³.

8a. Leushi

Average density of 0.22 g/cm³ in open areas, and 0.19 in protected areas.

8a. Tara

Average density of 0.28 g/cm³ in open areas and 0.26 in protected areas.

8a. Shores of Kara Sea

Average density of 0.30 g/cm³.

8a and 12. Western Siberia Lowlands

Density ranges from 0.19 to 0.25 g/cm³.

8a. Southern Regions of West Siberian Lowlands

Average density of 0.23 g/cm³.

8a. Bolotnoye

Average density of 0.25 g/cm³ in open areas and 0.21 in protected areas.

8a. Region of light winds

Average density of 0.22 g/cm³.

Table A1 (cont'd).

8a. Cheremkhovo

Average density of 0.20 g/cm³ in open areas and 0.18 in protected areas.

8a. Lena River Valley

Average density of 0.14 to 0.16 g/cm³.

8a. Yakutiya River valley

Average density of 0.20 g/cm³.

8a. Lower central Siberian Plateau

Density varies from 0.18 to 0.20 g/cm³.

8a and 8c. Transbaykal Territory

Average density ranges from 0.14 to 0.18 g/cm³.

8a. Sretensk

Average density of 0.16 g/cm³ in open areas, and 0.14 in protected areas.

8a. Kolyma River Valley

Average density of 0.20 g/cm³, ranges from 0.14 to 0.26 prior to spring melt.

Along shores of Siberian Plateau

Average density of 0.28 g/cm³.

8a and 12. Along shores of Se of Okhotsk

Density ranges from 0.19 to 0.25 g/cm³.

8a. Chukotka region

Average density of 0.30 g/cm³.

8a. Northeast Kamchatka

Densities of up to 0.36 g/cm³ along the coast.

8a. Western Siberia

Average densities of 0.19 to 0.22 g/cm³ in forest zones, and 0.22 to 0.31 g/cm³ in forest-steppe and steppe regions.

9a. European U.S.S.R.

Density ranges from 0.20 to 0.28 g/cm³ in January, and 0.28 to 0.30 g/cm³ in March.

9a. Far North

Density ranges from 0.20 to 0.25 g/cm³ in November, and 0.30 to 0.33 in March.

9d. Novaya Zemlya

On the east coast density is 0.28 to 0.35 g/cm³ in January and up to 0.39 in May. On the west coast the wind is not as strong and density in March is not greater than 0.27 to 0.30 g/cm³.

9d. Dikson Island

Density in the spring was 0.38 g/cm³.

9e. Lower Volga River

Density range at end of February is 0.23 to 0.32 g/cm³, increasing to 0.40 during thaw.

9e. Central and Western Kazakhstan

Density at the time of maximum snow depth ranged from 0.30 to 0.35 g/cm³.

10. Terskol Peak and Asau River Valley

Snow cover with five layers: Top layer hoar frost density is 0.27 g/cm³, next lower layer consisted of fine-grained snow, next layer was packed snow at 0.40 g/cm³, next layer fine-grained at 0.33 g/cm³, and bottom layer was hoar frost.

11. Irkutsk

Average of three winter snow covers with three layers. Density of top 5 cm was 0.141 g/cm³, density of middle 5 cm was 0.214 g/cm³, and density of lowest 10 cm was 0.226 g/cm³.

12. Priamurye region

Density by the end of winter increased from between 0.13 and 0.15 g/cm³ to between 0.18 and 0.20 g/cm³.

12. Amur Lowlands

Density ranges from 0.19 to 0.25 g/cm³.

12. Northeast region

Density by the end of winter increases from 0.15 to 0.20 g/cm³.

14. Kamchatka

Density dependent on time of year, exposure, topography and snow depth. It ranges from 0.10 g/cm³ in October and November, to > 0.34 at maximum snow depth. Values are lower in interior sheltered areas and highest along coast and windy areas.

U.S.S.R., the heaviest concentration of information appears in the western (i.e. European) part of the country. Since Appendix A was included mainly for informational purposes, additional attempts to analyze the material were beyond the scope of the report.

1. THE CAUCASUS.

1a. Gurtovaya, Sulakvelidze, and Yashina (1960).

The average density of the snow cover [in this mountain region] from November to April is less than 0.20 g/cm³, and in areas sheltered from the wind it generally is not more than 0.10-0.15 g/cm³; the lowest density is found in February. Beginning with the end of April, the density of the snow cover, resulting from the thawing brought about by solar radiation, increases and by August reaches 0.35-0.40 g/cm³ at the lower boundaries and at the upper boundaries it is less than 0.30 g/cm³.

Because of the great variations in elevation and the physico-geographic variations in the area of the Caucasus, the snow cover density is very variable. It is first of all necessary to mention the conformity observed in the entire territory of snow cover density lessening as the elevation above sea level increases. Thus, from 0.40 g/cm³ at a 1000 m elevation, it decreases to 0.15-0.20 g/cm³ at the 3000 m elevation (February-March). In February-March, snow cover density of 0.30-0.40 g/cm³ is characteristic for the Western Caucasus. It is not constant, however, and changes within a great range depending on the meteorological conditions and the slope exposures. On the northern slopes, the snow cover density for practically the entire winter is, as a rule 0.05-0.15 g/cm³, which is below that of the southern slopes. At the beginning of the thaw and with clear weather, the difference in the snow cover density of these slopes may reach 0.20 g/cm³. A noticeable increase in density occurs uniformly on all slopes during wet snow falls and advective heating.

DISTRIBUTION OF SNOW COVER DENSITY (in g/cm³) IN THE
WESTERN CAUCASUS

Elevation above sea level, in m	Months	
	Dec-Jan,	Jan-Mar
2800	0.18-0.20	0.28-0.30
2300	0.21	0.33
2000	0.22	0.35
1700	0.23	0.34
1500	0.25-0.28	0.35-0.37
1000	0.30	0.40
500	0.35	0.40

According to the information of snow measuring surveys in the basins of the Bzipi, Koderi, Inguri, and Riono Rivers, a sufficiently clear conformity was developed for the change in density by time and elevation. The density of the snow cover decreases from 0.30 g/cm³ at an elevation of 1000 m, to 0.18 g/cm³ at an elevation of 2800 m, and [increases] from 0.30 g/cm³ in December-January to 0.40 g/cm³ in February-March. Snow density action is particularly evident during the spring months. In March, the density of the snow at the upper levels of this zone is on the average 0.10-0.12 g/cm³ higher than in December. In the lower belts of mountains, this difference decreases to 0.05-0.10 g/cm³.

In the eastern Priel'brus area the most characteristic density during the winter at the bottom of the Bakaan River Valley is 0.25 g/cm³ which decreases quite a bit during snow fall because of the freshly fallen snow. At the end of March and beginning of April, in the period of snow thaw, the density of the snow cover [increases] to 0.35 g/cm³. The difference in snow cover densities between the northern and southern slopes is [evident]. In January, the density of the snow, because of the great insolation on the southern slope, is 0.07 g/cm³ greater than on the northern, while in March, it is 0.09 g/cm³ greater. The density of the snow at the bottom of the valley has an [uniform] value in comparison with its densities on the northern and southern slopes.

The given values of snow cover density are averaged values for these regions. A detailed snow measuring survey provides considerable differences in its density because of the character of the vegetative cover and the microclimatic conditions. The density of the snow is usually less in the forest than on open areas. A certain increase in snow cover density is observed on the windward slopes because of its packing by the wind and wind-blown snow.

DISTRIBUTION OF SNOW COVER DENSITY (in g/cm³) IN THE
CENTRAL CAUCASUS

Elevation above sea level, in m	Months	
	Dec-Jan	Feb-Mar
3000	0.18	0.21
2500	0.20	0.21-0.28
2200	0.26	0.25-0.30
2000	0.24-0.26	0.31

An example of the clearest differentiation of the snow thickness on the stratigraphic line may be that of the following description on the snow cover in the upper reaches of the Baksanskiy ravine (Central Caucasus) during the middle of February 1957. At an average snow cover thickness of 45-50 cm, the uppermost 0-5 cm were presented as newly fallen, friable, snow consisting of fragments of flakes (density 0.10-0.15 g/cm³); lower, at a depth of 5-25 cm, there was deposited a firm, dense wind-blown slab (density of 0.36-0.40 g/cm³) of small-grain snow. The thickness of the wind-blown slab varied from 10 to 25 cm. At a depth of 25-50 cm, was a line of friable plutonic hoarfrost. The plutonic hoarfrost crystals were large (5-7 mm), prismatic, hollow inside, and vertically oriented. They were solidly connected into clusters. The vertical texture with suspended clusters of crystals and a large amount of space between them was depicted very clearly. Lower, at the edge of the grass cover, the crystals were larger (to 10 mm), and the amount of volume of the space increased. The solidly frozen together crystals abutted with grains and gradually assumed a cylindrical form and horizontal orientation. The porosity of the lower line was so great that the frozen together crystals gave the impression of a porous screen; the ground was thawed.

Changes in the density of the snow thickness by the structural lines in forested and in open areas take place dissimilarly. In non-forested sectors, the upper line of the snow cover has a relatively larger density (0.21-0.28 g/cm³) than the lower (0.15-0.23 g/cm³) because of solidifying by wind and radiational heating. In the forested sectors, because of sheltering of the line by the crowns of trees, there is no solidifying of the upper line. Therefore, in the shaded sectors of the forest, one observes stages in the distribution character of the density along the profile of the snow thickness - namely, the development of density from the upper lines to the middle and its decreases in the line of friable plutonic hoarfrost.

1b. Glazyrin and Denisov (1967). Snow cover density measurements made during one winter at a meteorological station situated on the Krestoviy pass in the Caucasus ranged from 0.18 to 0.26 g/cm³ during January through mid-March, and from 0.28 to 0.43 g/cm³ during late March through April.

1c. Zal'ikhanov (1967). In this article the distribution of precipitation in the mountainous regions of the Karbardino-Balkaria ASSR is analyzed in terms of seasons and altitude zones. It is shown that complex orography, a great range of altitudes above sea level, and proximity to the Black Sea create a complex picture of precipitation distribution. On the basis of analysis of snow-measuring surveys, it is demonstrated that a zone of great snowiness is located just 0.6-1.1 km to the north of the Primary Divide at an altitude of 2900-3200 m. These data differ considerably from the figures obtained previously by a number of researchers.

According to the data of G.K. Sulakv'el'idze, from November until April the average snow density in a zone of stable snow cover is below 0.2 g/cm³, and 0.10-0.15 g/cm³ in areas protected from wind. According to measurements conducted by us in April, 1964, the average density in open areas in a zone of constant snow cover is 0.20-0.25 g/cm³.

When snow falls on a surface of frozen earth, the contact of the snow cover with the surface of the earth is usually tenuous, which is conducive to movement of the snow on the surface of slopes, i.e., to the formation of avalanches.

The descent of a snow avalanche is importantly affected by the density of the snow. In studying the dependence of snow density on the elevation of the locality, the following features are noted:

(1) the density of snow cover decreases with an increase in absolute elevation of the locality. This is attributable to a decrease in the rapidity of the thawing process with increasing elevation.

(2) for the 2000-3000 m elevation zone, density of the snow cover, according to our data, is distributed with respect to the time of year in the following manner: when the snow falls, its density varies within the limits 0.10-0.20 g/cm³; in the first half of the winter, when the air temperature is relatively high, the snow condenses, and its density reaches 0.30 g/cm³. In open areas wind contributes to snow compaction;

(3) in February, as a result of large snowfalls the average density of snow cover again falls, reaching 0.20 g/cm³;

(4) in the spring, during the initial thawing of the snow cover, the snow density reaches its maximal value of 0.40-0.45 g/cm³. This density is apparently attributable to the fact that the water formed during the initial thawing does not reach the surface of the soil, but freezes in the snow mass, which leads to an increase in its density during the spring period;

(5) in April the density of the snow again declines to 0.30-0.40 g/cm³. This may be explained by the fact that thawing takes place very rapidly as a result of high, positive air temperatures, and the water flows away. Values for elevation and density of snow-cover measured during a 1964 expedition are presented in the following table. [Table A2].

Table A2. Elevation and density of snow cover in April, 1964, in the territory of Kabardino-Balkaria.

Elevation (m)	Adyl-Su		Adry-Su		Bashli'- Auzu-Su		Gara- Auzu-Su		Cherek- Bizengi j		Karasu-Balgarian (L'kezi)	
	Depth of snow cover* (cm)	Snow density (g/cm ³)	Depth of snow cover (cm)	Snow density (g/cm ³)								
2200	6	0.38	8	0.36	16	0.34	8	0.34	12	0.32	16	0.31
2300	12	0.37	13	0.36	30	0.27	44	0.31	15	0.33	18	0.32
2400	28	0.36	18	0.36	46	0.30	108	0.30	17	0.29	12	0.31
	30	0.36	16	0.35	52	0.34	51	0.29	16	0.28	17	0.33
2500	38	0.37	14	0.36	34	0.30	64	0.30	17	0.26	19	0.35
2600	36	0.34	25	0.35	30	0.29	69	0.29	20	0.27	24	0.34
2700	39	0.32	36	0.34	32	0.28	71	0.26	22	0.22	23	0.30
2800	42	0.30	37	0.32	37	0.26	90	0.27	27	0.26	27	0.29
2900	46	0.29	40	0.31	37	0.21	170	0.28	30	0.26	31	0.26
3000	42	0.27	42	0.30	42	0.24	146	0.27	32	0.25	30	0.21
3100	40	0.26	44	0.28	42	0.23	139	0.26	42	0.26	26	0.22
3200	36	0.26	43	0.27	46	0.24	141	0.28	49	0.24	20	0.21
3300	34	0.21	42	0.27	15	0.23	149	0.27	34	0.21	29	0.21
3400	-	-	-	-	21	0.21	120	0.20	22	0.20	-	-
3500	-	-	-	-	19	0.11	107	0.15	19	0.20	-	-

*The depth of snow cover in each valley increases with an increase in the absolute elevation of the locality up to a certain elevation limit, and then it begins to decrease.

2. TERSKEY - ALA TAU MOUNTAIN REGION

2a. Iveronova (1960). The snow cover density during the winter has a clearly defined course. In the period of the fall "striated" landscape, there occurs a significant growth in the snow cover density; it changes little from December through February (the absolute value about 0.20 g/cm^3); it [increases] again strongly in March and reaches a maximum in April at the height of thawing. The density of the snow during the winter (November-March) at an elevation of 3250 m, close to the end of Karabatkak Glacier in a valley bottom (low grass alpine meadow), varies in insignificant limits (on the average, $0.20\text{-}0.22 \text{ g/cm}^3$). It begins to [increase] in April and reaches its maximum [density] during the height of the thaw - in May.

The density of the snow cover during the period of its continuous deposit has a rather clearly defined course: in the fall, in September-October, it increases, by November or December it generally falls a bit and then until May, to April in some years, it remains rather constant varying in small limits and on the average reaches rather large absolute values ($0.23\text{-}0.25 \text{ g/cm}^3$). In May (or in April) the density of the snow begins to increase sharply and reaches a maximum at the height of thawing, in July, and occasionally in June.

2b. Sadvakasov and Kozik (1970). In the western spurs of the Tyan'-Shan' Mountains, we find the Kyzylcha River, forming a tributary of the second order to the Akhangaran (Angren) River. The Kyzylcha Basin has an area of 53.4 km^2 ; the height of its individual parts will fluctuate from 1200 to 3800 m above sea level; for more than 15 years, snow surveys have been conducted in the basin at permanent snow-measuring points during the cold season. In 1958 in the basin's center at a height of 2075 m, we established the Kyzylcha Snow-Avalanche Station. With its organization, the snow surveys became more regular and detailed; specifically, fairly detailed spring snow surveys from March to April were initiated in the basin of a ravine with an area of 0.21 km^2 , specially chosen for water balance observations. The elevation mark of the ravine floor was about 2000 m above sea level, with average height of watershed relative to the closing line of the ravine amounting to 192 m. The meteorological site of Kyzylcha Station is located about 1 km from the ravine and is separated from it by a water divide occurring approximately at the station's altitude.

The snow survey data for six spring periods from 1960-1965 [were reviewed]. In the snow surveys, it was customary to take measurements of the depth and water content of the snow cover in several high zones based on horizontal traverse routes, tied into the terrain by permanent rods. The snow surveys were started in March at the beginning of the spring snow thawing and were ended when more than 90% of the ravine's surface had become free of snow. This usually occurred in April, rarely in May. The time spans between successive snow surveys usually comprised 1-3 days. Depth of snow in the traverse routes was determined approximately every 20 m while water content was estimated every 100 m. During the 6 years

indicated, data from 74 snow surveys were examined. Based on the data for each snow survey, the density value for the snow cover in the ravine during these spring months ranged from a minimum of 0.28 to a maximum of 0.49 g/cm³. The average density observed was about 0.38 g/cm³.

3. SOUTHERN PART OF THE BURYAT REGION.

Nefed'yeva (1960). The density of the Zabaykal'ya and [Gusino Lake depression] snow covers is comparatively low (0.15-0.18, rarely 0.20). It changes little during the course of the winter which is explained by the stable frosts, infrequent snow storms, and lack of thaw. By the end of March in 1956, in the majority of the steppes and forest-steppe regions of the Buryat, the snow density was 0.13-0.20. Direct observations have shown that in the open steppe regions, the snow cover, in spite of the low average density, is distinguished by rather uneven depositing explained by the wind redistribution of the snow.... The amount of water reserve in the wind blown snow drifts increases in comparison with the average by 5-10 times and the snow density increases by 30 to 35 percent.

4. MOSCOW AREA.

4a. Nefed'yeva (1960a). Frequently, the snow layers were separated by ice lenses - the remains of winter thaws. The total density of the thickness was not great and varied from 0.18 to 0.27 g/cm³.

In the period from the middle of March to the middle of April there also was an increase in the reserves of wetness with heavy snow falls. The daily thaws observed during this time facilitated an increase in the density of the snow cover to 0.22-0.34 g/cm³, and in the wind blown drifts and edges to 0.35-0.40 g/cm³.

The maximum water reserves were observed in areas of snow drifts which should be a factor of both the great thickness of the snow cover (up to 110 cm) and its increase in density (to 0.29 g/cm³). In open fields around Moscow observed snow cover densities ranged from 0.13 to 0.22 g/cm³ during November and from 0.17 to 0.24 g/cm³ during December. In small open sections in wooded areas around Moscow, densities ranged from 0.13 to 0.18 g/cm³ in November and from 0.16 to 0.20 g/cm³ in December. In wooded areas under trees observed densities ranged from 0.13 to 0.18 g/cm³ in November and from 0.15 to 0.20 g/cm³ in December.

4b. Sabo (1962). Study of the formation of snow cover in the forest has been carried out by us during 1956-58 in the Taldom district, Moscow oblast, for the purpose of determining the water balance of swamped forest lands. The studies have been carried out in the same twenty areas during the period of maximal snow supplies. With the help of detailed snow-measuring surveys in different types of forests, with standing timber of varying age and fullness, as well as in field areas, determinations were made of the depth of snow cover, its density and water supplies, referred to hereafter as the "snow supply." A description of the areas and results of the studies with respect to snow depth and snow cover density are presented in the following table [Table A3]:

Table A3. Depth and density of the snow cover in periods of maximal snow supply, 1956-1958.

Land	1956		1957		1958	
	Depth (cm)	Density (g/cm ³)	Depth (cm)	Density (g/cm ³)	Depth (cm)	Density (g/cm ³)
Pine Forest	75.8	0.256	72.0	0.258	73.1	0.184
	60.7	0.221	56.0	0.253	63.7	0.230
	61.0	0.218	46.5	0.300	53.0	0.255
Spruce Forest	-	-	-	-	46.7	0.248
	54.5	0.216	46.1	0.297	56.3	0.243
	52.1	0.213	36.6	0.276	48.3	0.248
	47.9	0.217	-	-	-	-
Birch Forest	65.8	0.235	-	-	-	-
	62.7	0.230	53.3	0.263	61.8	0.239
	57.5	0.229	48.7	0.325	56.0	0.232
	62.5	0.228	55.8	0.285	59.1	0.269
Birch Forest after maint. fellings	63.9	0.225	54.8	0.258	60.7	0.232
	63.8	0.234	55.0	0.275	60.3	0.257
	65.2	0.240	57.9	0.265	62.9	0.243
	67.2	0.241	61.2	0.253	63.5	0.249
Glade in forest	61.0	0.240	52.9	0.304	57.5	0.267
	68.0	0.281	59.5	0.310	64.2	0.248
	-	-	-	-	45.3	0.252
Field	50.0	0.241	-	-	45.7	0.256
	44.8	0.265	-	-	43.6	0.264

5. AMUR REGION.

Nikol'skaya and Grigor'yev (1960). The average density of the snow cover on the Zeysko-Bureinskaya Plain was 0.29 g/cm³, and the largest was 0.39-0.40 g/cm³. It was observed on sectors with lightly rolling surfaces, among wormwood growths.

At the beginning of March the density of the snow cover approached 0.40 g/cm³ in the Kozloki River Valley. However, the seasonal average density of the snow cover in this area was 0.27 g/cm³. The average density in March in the Burei River Valley was 0.25 g/cm³, while the minimum and maximum were 0.11 and 0.40 g/cm³.

6. EUROPEAN PART OF THE USSR.

6a. Kus'min (1960). Among the more important characteristics determining the water, thermal, radiational, and many other snow cover properties, are the structure and density of the snow. The extreme limits in the variation of snow density are 0.01 and 0.70 g/cm³. The density of dry newly fallen snow which has not lost its initial structure changes from 0.01 to 0.20 g/cm³, small-grain from 0.17 to 0.43 g/cm³, and large grain from 0.32 to 0.49 g/cm³.

The average density of a dry snow cover by the beginning of the spring thaws (at the end of winter) changes over the territory of the European part of the USSR from 0.18 to 0.35 g/cm³ and, on the average, is 0.28 g/cm³. In the north of the European territory of the USSR, the average density of the dry snow cover by the beginning of the thaw changes from 0.22 to 0.28 g/cm³, and in the central area, from 0.24 to 0.32 g/cm³, and in the south from 0.22-0.23 to as high as 0.34-0.36 g/cm³. Snow density in the forest, as a rule, is lower than on the open sectors.

The density of the melting snow cover and the different forms of melting snow were characterized in a laboratory study with the following results. The density of the newly fallen snow, having an initial density of from 0.13 to 0.18 g/cm³, after being dampened by water to its full water capacity and with a free drain of the surplus gravitational water, increases to 0.55-0.60 g/cm³; the density of the granular snow with an initial density of 0.23-0.45 g/cm³ increases after wetting to 0.40-0.50 g/cm³. The density of the naturally melting snow changes at the beginning of thawing from 0.18 to 0.35 g/cm³; at the height of the thaw, from 0.35 to 0.45 g/cm³; and reaches 0.50 g/cm³ by the end of the thaw season. The density of the snow cover in the characteristic dates of [seasonal] thaw varies as follows: at the date of the beginning of snow thawing from 0.24 to 0.41 g/cm³, on the date of the appearance of thaw patches from 0.25 to 0.45 g/cm³, and on the date of snow dwindling on 50% of the area from 0.29 to 0.48 g/cm³.

6b. Vershinina and Volchenko (1974). The water equivalent of snow accumulated over the winter is the main factor determining the magnitude of the spring flood runoff in regions deficient in moisture, since soil moisture varies little from year to year. For this reason, it is particularly important to compute the maximum water equivalents and their variations in these regions.

The southern boundary of the region studied extends along the Smolensk-Vilnius line, the eastern boundary along the Kalinin-Borovich line, the western boundary along the Daugavpils-Riga-Narva line (with the exception of the Estonian SSR), and the northern boundary along the Narva-Leningrad line.

Regular snow surveys along triangular snow courses in the fields are known to have been initiated in 1936-1938. However they were temporarily discontinued in 1941-1944 and, therefore, continuous snow cover observation records are available only from 1945 to 1949. To compute the long-period mean maximum water equivalent and its variation (coefficients of variation), we used the observation period from 1949 to 1971.

To determine the maximum water equivalent in the forest, the results of snow surveys at 60 points with observation records from 15 to 20 years were used. The results of observations of water equivalent in the forest at each point were then reduced to a single 18-year period....

The average snow depth in the forests of the region under study varied over the observation period from 25-30 cm in the Latvian SSR to 50-60 cm in its eastern part, and in the Valdai Hills in particular. In snowy winters, snow depth may reach 80-100 cm in deciduous and mixed forests. The year-

to-year variations in average snow depth in the forest are fairly large; the coefficient of variation ranges from 0.25 to 0.45.

Snow density in a forest depends largely on the species composition and density of the forest. For this reason, there is no distinct pattern in the distribution of average snow densities in the forest over the region under study, since snow courses in the same area may be located in different types of forests. During the establishment of maximum water equivalent, the average snow density in the forest, just as in the field, varies insignificantly over the territory, from 0.22 to 0.28 g/cm³, and averages 0.25 g/cm³, i.e., 0.03 g/cm³ less than in the field.

6c. Molchanov (1946). The first snow in the Onega Lake District of European USSR is observed during the last ten days of September in the northern section but somewhat later or in the first ten days of October in the southern section. The very latest occurrence of a snow cover in the northern and eastern portions is the first ten days of November, but in the last ten days of November in remaining portions.

A continuous snow cover is established first, on an average, in the middle of October (11 to 20th Oct.) in the northern portion, but about ten days later (21 to 31 Oct.) in the southern portion; only at Klimetskiy Ov and Petrozavodsk (in the central portion) where the climate is influenced considerably by the lake, is the establishment of the cover delayed until the first ten days of November. The disappearance of snow cover in open places occurs (on an average) during the last ten days of April; in forests, it disappears much later or as late as the end of May and in certain cases even up to June.

In exceptionally warm years and winters with little snow, the cover in open places disappears before the 20th of April, and even as early as late March in the southern portion of the basin. Conversely, in winters with heavy snow, the cover lasts until mid-May (11 to 20th) and at some northern stations until the end of May.

During the spring cold waves (return of cold), a post winter snow is possible. For example, in 1930, when snow fell at the beginning of June, the snow remained several days. On an average, the duration of snow cover is 184 to 206 days annually in the northern portion and 178 days at the southernmost station (Voznesen'ye); Klimetskiy Ov, under the lake influence has the shortest duration or a total of 169 days annually.

Generally, the snow depth in late October is one to two centimeters at all stations except at Klimetskiy Ov (lake influence) where such a depth is not recorded until the middle of November (11 to 20th). The depth gradually increases from November through March when maximum depths are recorded everywhere in the basin. The greatest depths occur in the northern and eastern portions of the basin where the mean depth is 54 to 60 cm and the absolute maximum is 92 cm at Morskaya Masel'ga. The absolute maximum depth for the entire basin, however, is in Pudozh (eastern portion) where the depth reached 108 cm. In the western and southern portions, the deepest average depths vary from 30 to 47 cm or about 20 cm less than in the northern portion. Depths in the central portion (lake influence) are relatively smaller.

As for snow density, freshly fallen snow is very [light], but the snow cover density increases as winter progresses. In January, the density is 0.20 to 0.26 g/cm³, but in April (when the thaw sets in) it is 0.38 g/cm³. In winters with great snowfalls, and considerable thawing, the snow cover density may vary within considerable limits. The following snow-cover density data were derived from observations made from 1928 to 1933:

Table A4. Snow cover density data, 1928-1933.

<u>Station</u>	<u>Location (lat. & long.)</u>			<u>Elev. (m)</u>		
1. Voznesen'ye	61°01'N	35°29'E	38			
2. Kargopol'	61°30'N	38°57'E	126			

Stations	November			December			January		
	1 to 10th	11 to 20th	21 to 30th	1 to 10th	11 to 20th	21 to 31th	1 to 10th	11 to 20th	21 to 31th
1. Voznesen'ye									
a. Max	0.12	0.15	0.19	0.18	0.16	0.17	0.19	0.32	0.26
b. Avg	--	0.11	0.14	0.11	0.13	0.16	0.22	0.21	0.21
c. Min	--	0.09	0.08	0.08	0.10	0.10	0.15	0.14	0.10
2. Kargopol'									
a. Max	--	--	0.17	0.25	0.19	0.22	0.26	0.30	0.33
b. Avg	--	0.14	0.13	0.21	0.14	0.17	0.21	0.21	0.25
c. Min	--	--	0.09	0.17	0.10	0.12	0.15	0.18	0.18

Stations	February			March			April		
	1 to 10th	11 to 20th	21 to 28th	1 to 10th	11 to 20th	21 to 31th	1 to 10th	11 to 20th	21 to 30th
1a.	0.20	0.26	0.23	0.26	0.25	0.35	0.36	0.36	0.50
1b.	0.17	0.21	0.20	0.21	0.21	0.26	0.30	0.32	0.38
1c.	0.15	0.16	0.18	0.19	0.20	0.20	0.24	0.26	0.26
2a.	0.32	0.31	0.33	0.35	0.28	0.27	0.39	0.42	0.40
2b.	0.24	0.25	0.25	0.26	0.23	0.24	0.29	0.36	0.38
2c.	0.20	0.21	0.21	0.21	0.21	0.23	0.24	0.25	0.24

Stations	May		
	1 to 10th	11 to 20th	21 to 31th
1a.	0.33	--	--
1b.	--	--	--
1c.	--	--	--
2a.	--	--	--
2b.	0.31	--	--
2c.	0.27	--	--

6d. Vershinina and Volchenko (1979). The boundaries of this study area are 56 and 60°N, 47 and 56°E. This area includes the basins of the Vyatka River upstream of Vyatskiye Polyany city (basin area of 124,000 km²), and

the upper courses of the Kama River upstream of Kay village (13,000 km²) and of the Vetluga River upstream of Bystri village (2680 km²).

This territory occupies the eastern part of the East European plain with elevated, dissected interfluves and broad river valleys with gentle terraced slopes. The lowest and most swampy (13-18%) area is the southwestern part of the basin of the Vyatka River, where the lowland crossed by its tributary, the Pizhma River, has elevations of about 100-145 abs. m and the drainage divides rise no more than 170-180 m.

A belt of flattened uplands, the so-called Northern Hills, with elevations somewhat higher than 200 m, extends in the northern part of the region along the drainage divide with the basin of the Northern Dvina River.

The flat Upper Kama upland, deeply dissected by rivers, is located in the upper courses of the Kama, Vyatka, and Cheptsy Rivers. The elevations of the broad and flat drainage divides here do not exceed 330 m.

A stable snow cover usually forms in the study region in the second decade of November and increases most rapidly in early winter.

The water equivalent is usually highest by the end of March in the southern and central regions and by mid April in northern regions.

At the time of establishment of the maximum water equivalent, snow depth in the field averages from 30 to 50 cm (on the right bank of the Lower Vyatka) to 70 to 80 cm in the upper courses of the Vyatka and Kama Rivers. The temporal variations in average snow depth in the field range from 0.26 to 0.30, depending on the water equivalent and local topography.

Snow density in the field varies insignificantly in the period of establishment of the maximum water equivalent from 0.26 to 0.32 g/cm³, and averages 0.28 g/cm³ in most areas. Its temporal variability is also low, averaging 0.10.

The average height of snow cover in forests during the entire period of observation, and within the limits of the experimental territory, changes from 50-60 centimeters in south-eastern regions to 80-90 centimeters in the upper reaches of the Vyatka and Kama Rivers. From year to year changeability of the average figures for the height of snow cover in forests is negligible. The average density of snow cover in forests during the acknowledged periods of maximum snow reserves just as for fields changes only slightly from 0.24 to 0.28 g/cm³, comprising on the average 0.26 g/cm³ which is 0.02 g/cm³ lower than in fields.

The average figures for the maximum water reserves in the snow cover within the limits of the experimental territory and during the observation period change from approximately 120-150 mm in southwestern regions to 230-240 mm in the Verkhnekamsk upland area. During the winters of thick snow cover the snow reserves in forests reach the figure of 250-260 mm. Approximately 10-20% more snow is accumulated in forests than in fields.

7. ARCTIC REGIONS.

Bogorodskii (1975). The density of snow cover in the Arctic varies within

great limits depending upon the time of year, the temperature and wind conditions, the thickness of snow cover, and so on. Freshly fallen snow is distinguished by the least density, but depending upon the conditions under which it falls its density may vary by 10 times --from 0.03 g/cm³ (fluffy snow in windless weather) to 0.30 g/cm³ (moist snow).

In its annual trend, the mean snow density [increases] rapidly in the first 2 months and then remains almost constant (0.30-0.35 g/cm³) all winter. At the beginning of the thaw the density rises sharply and attains 0.40-0.45 g/cm³.

The density of snow cover is not uniform in depth. According to the author's observations it increases by two times at a depth of 3 m from 0.22 to 0.45 g/cm³, the growth in density occurring exponentially:

$$\rho = 0.11h^{0.243}$$

where ρ is density, g/cm³, and h is depth, cm.

Temperature has little effect on snow cover compaction: at -2°C the density of freshly fallen snow is 0.10 - 0.11 g/cm³, while at -10°C $\rho = 0.07 - 0.08$ g/cm³. However, the snow cover compacts rapidly as the wind increases as follows:

Wind Speed m/s	Snow Density (g/cm ³)
0 - 2	0.05 - 0.06
3 - 7	0.14 - 0.17
8 - 15	0.18 - 0.22
<u>≥ 16</u>	0.23 - 0.33

8. THROUGHOUT THE USSR.

8a. Lipovskaya (1968). The principal factors, governing the distribution of the snow cover densities throughout the USSR, are reviewed. In general, the snow cover density was found to increase with the weight of the snow, the wind velocity and the temperature changes. Within the range of 0.27 to 0.36 g/cm³, the density rises by 0.04 g/cm³ for every degree of positive temperature. It also depends on the length of time the snow cover exists. Thus, the density of snow cover triples from 0.05 to 0.15 g/cm³ while its thickness decreases by 8 to 10 times. In temperate zones, the snow cover density is from 0.20 to 0.23 g/cm³ in forested and from 0.25 to 0.27 g/cm³ in unshielded stretches. In the mountains, it varies with the slope, orientation, snow-cover thickness and altitude above sea level. For example, the density increases in the Caucasian Mountains with altitude. The relationship between the snow cover and the topographic relief is shown in Figure A2. The density increases from 0.14-0.16 g/cm³ in the snow-deficient southern regions, to 0.30 g/cm³ on the north shores and to 0.36 g/cm³ on Kamchatka. Fluctuating between 0.20 and 0.22 g/cm³ in the western parts of the European USSR, the density rises to 0.24 g/cm³ on the islands of the Gulf of Riga. It is mostly 0.24 g/cm³ in the interior of the European USSR. In the West Siberian lowlands, the density varies from 0.23 on the south to 0.30 g/cm³ on the shores of the Kara Sea, averaging 0.26 g/cm³ over most of that territory. The narrow strip of [low?] densities, under 0.22 g/cm³ in the south of this territory coincides with the belt of

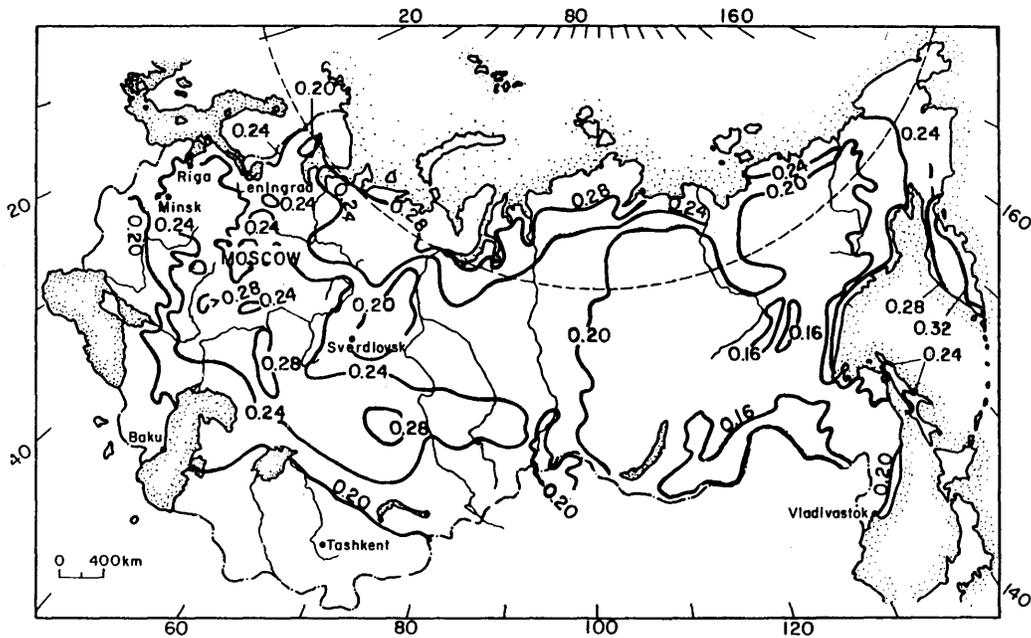


Figure A2. Average density of snow-cover (greatest 10-year depth for U.S.S.R.) from Liovskaya (1968).

winds having velocities of less than 5 m/sec. From there on eastward, the densities increase to 0.20–0.24 g/cm³ on shores of the Sea of Okhotsk.

The greater part of the European territory of the USSR consists of regions having average snow density of 0.24 g/cm³ and more. These regions occupy the area from 47 to 57°N. Lower density values are observed only in northwestern USSR, in the area near the Baltic, in southern Byelorussia, Moldavia, in the southern Ukraine, and along the Volga. On plateaus (Srednerusskaya, Privolzhskaya, and Obshchiy Syrt) density increases to 0.28 g/cm³. Density has a lower value in river valleys because of the considerable protection.

The greater density in the Cisurals and in the northern Urals (more than 0.28 g/cm³) is consistent with the distribution of snow depths and is brought about by the compaction of snow under its own weight.

Beyond the Urals the lowest values, 0.20 g/cm³, are noted on the Western Siberian plateau, which is explained by the decrease in snow depths and the heavily forested character of the territory (more friable snow). The range of density variations on the Western Siberian Plateau is as follows: from 0.22 g/cm³ in the south to 0.30 g/cm³ along the banks of the Kara Sea. Throughout most of this region density is 0.22 – 0.26 g/cm³. The narrow belt of decreased density (less than 0.22 g/cm³) in the south can be connected with the decrease in snow depths and is consistent with the lower average annual wind velocity, 2.5 – 4 m/s.

All of the northern Kazakhstan, particularly its elevated part, is characterized by an increase in density to 0.28 g/cm³, in spite of the insignificant depths of the snow cover. In this case the compaction of the snow is connected with the high wind velocities – all this territory belongs to the second, and partially to the first, wind regions with wind

velocities of 4-9 m/s. Local relief also has a considerable effect on compaction of snow.

On the lower part of the Central Siberian Plateau density varies from 0.18 g/cm³ to 0.20 g/cm³, and along the shore, to 0.26 g/cm³. Due to the great depths of snow in the mountainous part of the Asian Territory of the USSR, density values must also be high; however, data on these regions are practically nonexistent, and the stations at hand give information on snow density only in river valleys.

Particularly low density values, 0.14-0.16 g/cm³, are observed in the Lena Valley. In Yakutiya and Kolyma stations are also located only in the lower regions and in river valleys. Density distribution in these regions is consistent with depth distribution and mean annual wind velocity; density values here are on the order of 0.20 g/cm³*

As we progress to the east density increases to 0.30 g/cm³ at Chukotka. Along the coast of the Okhotsk Sea density is 0.20-0.24 g/cm³. The density gradient in the area near the coast is very large.

In a large part of the Transbaykal territory density is insignificant, 0.16-0.18 g/cm³, which is due to the low depths and the low wind velocity. In the mountainous part - the Stanovoy upland and the northern regions of the Yablonovyy upland - we can assume an increase in density of the snow cover. The near-coastal plain region of the Far East record densities of 0.16 to 0.20 g/cm³.

The greatest density for the entire USSR is noted on Kamchatka, particularly in the northeast along the coast, up to 0.36 g/cm³, where the snow is also the deepest.

The results of comparing densities in open and protected areas are presented in the following table [Table A5]:

Table A5. Density of snow cover in protected and open areas, g/cm³.

Station	Area		Location	
	protected	open	Lat. (N)	Long. (E)
Kargopol'	0.21	0.22	61°30'	38°58'
Bolotnoye	0.21	0.25	55°41'	84°23'
Leushi	0.19	0.22	59°39'	65°47'
Tara	0.26	0.28	56°54'	74°23'
Cheremkhovo	0.18	0.20	53°09'	103°05'
Sretensk	0.14	0.16	52°15'	117°43'
Uryupinsk	0.28	0.27	50°47'	42°00'
Kolomak	0.26	0.26	49°50'	35°40'

*During a visit to the upper reaches of the Vstrecha Creek Basin at the Kolyma Water Balance Station, Slaughter and Bilello (1977 - see list of cited references in text) obtained the following information. Snow depth measurement transects are made at 10-m intervals, along six marked lines running from ridgeline to ridgeline; snow water equivalent is measured at 100-m intervals. The frequency of snowpack measurement, depending on site, is every 10 days from date of first snow, monthly or annually at maximum accumulation. Spring snowmelt at the station begins in mid-April; snowpack density prior to spring melt ranges from 0.14 to 0.26 g/cm³.

8b. Kopanev (1978). Snow cover density is one of the important characteristics in scientific and practical studies of snow cover. Many factors determine the spacio-temporal fluctuations of snow density. Constant physical processes, the intensity of which depends on the nature of the precipitation and the thermo-cycle between the snow cover and the layer of atmosphere near the ground, begin to operate as soon as the snow layer is formed. Liquid precipitation usually solidifies snow cover, whereas thaws, followed by the melting of the snow, have a loosening effect; at low temperatures, however, a thaw may cause such a high density snow cover that a snow-crust will be formed.

The density of fresh snow depends on the air temperature during the falling of the snow (at relatively high temperatures there is a considerable increase in snow density), as well as on the size of the snow-flakes (small snow-flakes on windy days contribute to the formation of a relatively dense snow-cover). Snow density increases with the increase in the wind velocity even if all other conditions remain the same.

Snow density in areas exposed to wind is particularly affected; the wind is also the determining factor in regions where liquid precipitation or thaws occur rarely. Snow density in places well sheltered from the wind is not great since there it depends solely on the weight of the snow cover itself. Relatively dense snow covers are found in places exposed to the wind.

Recently published systematized data on the snow density of several areas of the Soviet Union show the average snow density, on days when the snow cover was the highest in Western Siberia, fluctuates from 0.19 to 0.22 g/cm³ in the forest zone and from 0.22 to 0.31 g/cm³ in the forested-steppe and steppe regions.

The work of V.I. Lipovskaya is of great interest since she was the first to make a map of the average snow density for the entire territory of the Soviet Union. The map reflects changes in the snow density in areas of maximum snow cover. The result of processing the snow measurements obtained from 800 posts and stations confirm the data about density distribution in different regions of the country. Regions of unstable snow cover (0.13-0.16 g/cm³) are characterized by minimal snow cover, which is the highest in the northern regions of Kazakhstan (0.24-0.28 g/cm³).

Average snow cover densities collected on the basis of a survey (1935-1965) conducted during the last ten days of each month show that the density increases, from month to month, simultaneously in all areas of the country. Over the largest part of the territory of the Soviet Union the greatest snow density was observed in forest clearings in March and in regions with prolonged cold periods in April.

During cold periods, differences in density between field and forest areas gradually become negligible and are least when snow precipitation reaches its peak (which usually takes place before the spring thawing).

There is no distinct zonal character in the density distribution in the months of March and April. In the European part of the USSR (ETS), for example, only a slight tendency toward growth is observed in a north-south direction. Meridional changes however, are more evident. In the central

European part of the USSR the density of snow in March and April in open spaces comprises 0.28-0.40 g/cm³, while in Yakutia the figures are 0.19-0.22 g/cm³. Absolute figures for forested areas of the mentioned regions show little difference.

Snow density in field areas is greater than under forest canopies and forest clearings. Density figures for clearings in forests and wooded areas are not the same in different regions of the country, i.e., the density in forest clearings may be equal to or less than the density in wooded areas. In the months of March and April the density in forest clearings in some areas does not differ significantly from the density in wooded areas. This is true of the Southern and Northern areas of ETS and of Eastern Siberia.

It may seem that the snow cover under the protection of trees in a forest is less affected by wind, solar radiation, etc., than that in open clearings. Fluctuation in the density distribution in forest clearings and in wooded areas is accounted for by inadequate methods of measuring snow density. Besides, changes in the height of the snow cover in forest clearings may occur as a result of thaws. If snow falls after a thaw in areas protected by trees where the old snow has been almost entirely preserved, the density may be somewhat higher than that in most of the forest clearings. The differences in the density between field and forest areas (for the months of March and April) are not high - particularly in the regions of Eastern Siberia.

8c. Zavarina (1976). The greatest snow cover density in the USSR is naturally seen in those regions where the snow cover is deepest. In regions with little snow, the density is lower than that in regions with much snow by a factor of 2-2.5.

Throughout the country, the average density values range -- under the influence of different factors -- from 0.14 (in the Trans-Baykal) to 0.36 g/cm³ (on the Kamchatka Peninsula); with a noticeable increase in snow cover density as the latitude increases. Snow density should increase as the air temperature rises. Consequently, its high value in the northern regions in comparison with the southern ones is explained by the action of other factors: first, the extended duration of its existence; second, the effect of a stronger wind. The joint effect of these factors overshadows the effect of air temperature.

Because snow cover depth increases as the altitude of an area above sea level does, so does the density of the snow. In winters with much snow, the change in the average density per 100 m of altitude reaches 0.02 g/cm³.

The degree of exposure of an area also affects snow cover density. As a rule, the density of the snow in exposed sections of forest-steppe and steppe regions is somewhat higher than in protected sections; this is apparently related to wind action. The difference between the densities reaches 5-10 percent.

In the southern region of the USSR, on the other hand, snow density in exposed sections is even somewhat less than in shielded ones. This can be

explained by the fact that here the difference in the duration of the snow cover's existence has a greater effect.

The average values characterize snow cover density quite well, since its variability from year to year is not great. For example, the coefficient of variability, as calculated for the territory encompassing the upper reaches of the Lena River, equals 0.09, while in the mountainous region of Kirgiziya it is even greater, but still does not exceed 0.18.

9. SELECTED AREAS IN THE USSR.

9a. Rikhter (1954). In the Arctic and the southeastern European portions of the USSR where snowfall is usually accompanied by strong winds or gales, the density of the snow cover tends to be higher than in the south and west. In eastern Siberia, with its lasting frosts and negligible wind velocities, the density of the snow cover remains very low all winter long.

The density of the snow becomes greater steadily all winter long, with the result that in spring the density of the snow cover is at its greatest in all its layers. The density of the snow cover as a whole rises an average of 10 to 12 percent per month.

The density of the snow cover increases most rapidly during warm spells and thaws, when it may change noticeably in 24 hours. The process whereby density increases proceeds without interruption during both warm spells and frosts, but it operates more slowly during frosts than warm spells. Contrary to the case of newly fallen snow, the density of the snow cover increases without interruption from fall to spring.

Snow cover may be divided into the following five groups by density: 1) very loose, with a density of 0.01 to 0.1; 2) loose, with a density of 0.1 to 0.25; 3) medium, 0.25 to 0.35; 4) dense, 0.35 to 0.45; 5) very dense, over 0.45.

When snow reaches a density of 0.32 to 0.35, it will sustain a pedestrian without skis. At 0.35 to 0.38, the foot hardly leaves a mark. At over 0.4, the snow will sustain a horse, and the human foot leaves no mark whatever.

The wheels of a horse-drawn wagon no longer break through when the density is over 0.3 or 0.35, but heavy trucks require a density of not less than 0.5.

Systematic records of snow density have been conducted by only a very small number of weather stations, and even these observations have, for the most part, not been processed and published to date.

The density of snow cover is exceedingly unstable both in time and space and is subject to extreme fluctuation (from 0.01 to 0.7) as a result of climatic conditions. These circumstances make it impossible to compile any more or less complete description of the geographical distribution of average snow cover density throughout the USSR., or of changes therein in time. We can provide only the following partial data on average density, month to month.

November. At the beginning of the period of snow accumulation when the snow has as yet not been packed down and redeposited by the wind, density does not vary greatly and is approximately 0.15. It is somewhat higher only in the far north (0.20-0.25), where it is packed down by strong winds and storms immediately after it has fallen.

December. The average density over most of the USSR now ranges between 0.16 and 0.20. It is somewhat above that in highland areas, and also in southeast European USSR and in the far north. In the central areas of eastern Siberia, it is insignificant and usually does not exceed 0.15.

January. The average snow density in European USSR is about 0.20, rising toward the north and the southeast to 0.25-0.26. In the Arctic proper, it reaches and exceeds 0.3. Minimum density (0.15, approximately) is recorded in eastern Siberia, where snowstorms and thaws are equally almost entirely lacking in January.

February. The average density of snow in European USSR is about 0.23-0.24. Higher average densities are recorded in the north where storms, tending to pack it down, prevail in the southeast (0.27-0.28), in the central Russian Upland (0.26-0.27) and in the west (0.25-0.26) where this is due to repeated thaws. The figure is at its highest in the Arctic, where it reaches 0.30-0.32. It is lowest in central and eastern Siberia, at 0.15.

March. The average density of the snow over most of the territory of European USSR is 0.28. In the southwest and west, where rapid thawing has begun, the average density increases to 0.30. Approximately the same order of magnitude is attained in the Arctic, at 0.30-0.33. In eastern Siberia it is now about 0.20.

April. In most districts of the USSR during this month, rapid packing down and thawing of the snow takes place, with its density climbing to 0.35. The same average holds in the Arctic as well. The lowest average density at this time is about 0.22, recorded in eastern Siberia.

May and June. Snow on the ground remains only in the far north and the Arctic, where it attains the highest of average densities (0.4-0.6 g/cm³).

9b. Trifonova (1962). The variability and precision in determining the density of a snow cover under various physical-geographic conditions was evaluated. For this purpose [the following snow cover density values] were obtained from snow surveys made in different regions of the USSR:

Snow survey region	Elevation m	Density g/cm ³
Kzylsu River	125	0.30
Karakol River	91	0.23
Koluton Station	1050	0.14
Valday Station	636	0.23
Oksochi Station	350	0.19
Koltushi Settlement	92	0.32

1. The following expedition snow surveys of the State Hydrological Institute were conducted in Kazakhstan (Tselinogradskaya Oblast) at the end of March 1956:

a) the snow survey route over a triangle with a total length of around 30 km within the limits of the basin of the Kzylsu River. The measurements of snow cover density were measured every 250 m (a total of 120 measurements).

b) snow surveys along two routes 10 and 13 km long along the Karakol River Valley (the first route ran along the right bank of the river, and the second along the left). The measurements of snow cover density were made every 250 m (a total of 900 measurements).

c) the snow surveys over 13 parallel routes crossing the channel of the Karakol River. The routes were located a distance of 100 m apart.

d) the snow surveys along 10 routes located in the region of Koluton Station. The total length of the routes was 10 km. The measurements of the snow density were made every 50 m. During the winter of 1954-1955, 5 snow surveys were made.

2. The data of the snow surveys by the Valday NIGL (nauchno-issledovatel'skaya geofizicheskaya laboratoriya) were made in the basin of the Polomet' River in February 1953. A total of more than 600 measurements were made for snow cover density.

3. The snow-measuring survey routes near Oksochi Station (the Northwestern UGNS [upravleniye gidrometeorologicheskoy sluzhby; administration of the hydrometeorological service]), in the basins of the tributaries of the Gridenki River (the Smolechanka and Oksochi Rivers). The snow surveys were made over two routes with a total length of 5, 100 m during two winters (1955-1956, 1956-1957). Snow cover density was measured every 100 m.

4. The snow survey route in the region of Koltushi Settlement of Leningradskaya Oblast consisting of two parallel routes around 500 m long each. The measurements of snow cover density on the routes were made every 10 m.

9c. Dmitrieva (1950). Systematic observations of snow density were begun in Russia in 1903-04. However, sufficient data on snow density were not collected even for such regions of European USSR as the Ukraine and the Northeast at the time the quantitative surveys of water resources were made during 1931-35. Data on density are scarcest for Siberia, especially the mountain regions....

The density of freshly fallen snow depends on snowflake size. Snow falls composed of large snowflakes reached densities of 0.056, medium flakes 0.091, and small ones 0.135 g/cm³. Fresh snow then rapidly settles to a density of 0.14-0.16 g/cm³.

The pressure of fresh snow on underlying layers is the first cause of increase in snow density. However, the snow crust under natural conditions

usually resists the action of gravity and contributes to a more regular distribution of load. The sublimation processes diminish the density of lower snow layers. In the forests of the temperate zone, where wind influence is low, the mean snow density near the end of winter does not exceed 0.20-0.23 g/cm³. Thus the mean snow settling during 3-4 winter months is negligible if the density of fresh snow is 0.14-0.16 g/cm³.

The generally well-known compaction effect of winter and spring thaw weather is also important. Some meltwater remains in the snow as a result of snow melting and causes an increase in the density of the snow. Obviously, considerable snow-cover compression will occur during alternating frost and thaw weather, as well as in areas where runoff of meltwater is hampered (in hollows and on slight slopes).

Facet fusion of snowflakes, subsequent settling, and compression of snow cover accompany the snow-melting process. Observations showed the dependency of snow density of the sum of positive mean diurnal air temperatures accumulated to the data of observations. These observations indicated that an increase in density of 0.004 over a density range of 0.27-0.36 g/cm³ corresponds to one degree of positive air temperature.

9d. Formozov (1939). In the northern portion of the European part of the USSR the snow cover lasts more than 200 days per year, and in many places in Siberia for 250 days and more. The average maximum snow depth in many places in Northern Siberia is 100 to 140 cm. In Teriberka (Murmansk coast) the number of days per year with snow-cover is 211; in Pustozersk 223 days, with an average maximum thickness of 72 cm. In New Port (southern part of the tundra on the Yamal Peninsula) the snow cover remains for about 9 months from the first 10 days of October until the end of June. Average depth in April and May is 65 cm. On Dickson Island snow remains for an average of 266 days and maximum depth is about 30 cm. Maximum snow cover is especially deep in the tundra of the western part of Siberia ranging from 100 to 140 cm. On the north coast, near the mouth of the Yana River, the thickness of the cover reaches 20 cm. At Anadyr Krai and on the Chukchi Peninsula it reaches 60 to 70 cm, and at Kamchatka as much as 90 to 100 cm. Extreme roughness of the weather which is peculiar to the winter season in the tundra, gives an unequal distribution of the snow cover and a high snow density. These conditions are linked with the steadiness and strength of the wind.

The density of snow on Novaya Zemlya (Matochkin Shore) in January is equal to 0.28 to 0.35 g/cm³, in May it reaches 0.39 and June 0.40 to 0.50 g/cm³. The west side of the Island, along with winds of less strength, the density in March was not greater than 0.27 to 0.30 g/cm³. At Marre-Sale in April 1921 the density was 0.36 and for greater Lyakhov Island (May 1931 and 1934) it ranged from 0.34 to 0.38 g/cm³. On Dickson Island toward the spring it was 0.38 g/cm³.

The depth of the snow cover in the broad leaf forest of European USSR increases from 0 to 5 cm in the southwest (Onesti region) to 40-50 cm toward the north and east central region (i.e. area along the Volga River).

Snow depth at Karelia averages 40 cm and increased to 80-90 cm eastward near the head waters of the Pechora River and to 95 cm at Shogor. In

western Siberia along the lower parts of the Ob' and Toz Rivers there is an abundance of cyclonic activity causing large accumulations of snow. The average maximum snow depth here reaches 100-140 cm. Maximum snow depth observed at Surgut is 80 cm, at Turukhansk 109 cm and central Siberia 50 to 70 cm.

Irkutsk Oblast, Yakutia, and Transbaikalia is a region of less snow; toward the month of March snow cover reaches an average of 37 cm at Yakutsk, 27 cm at Verkhoyansk, and at Irkutsk (in February) only 22 cm. Snow depth along the Pacific Coast, for example at Okhotsk, averages 50-60 cm at Kamchatka 100 cm: 5 cm at Blagoveshchensk, 10 cm at Vladivostok but 50 to 70 cm along the lower Amur River.

Toward the end of the winter 1892-93, the density of snow in Lesnoye in the areas between trees was 0.11 and in the glade it was 0.13 g/cm³. During most of the winter the density did not exceed 0.18 and only toward April did it reach 0.32 g/cm³.

The snow cover in the great plains and deserts of Russia is compacted by high winds and blown into the hollows and gullies leaving slopes, and hills bare. The density in these areas is higher than in the forest regions, averaging around 0.29 g/cm³.

9e. Kuz'min (1960). The particular conditions and distribution of the snow cover and its properties were studied for the Siberian forest-steppe region. In the case of strong winds (10-15 m/s), the snow is completely removed from exposed area. The snow is pulverized by the winds blowing from Kazakhstan dry steppes, becoming amorphous and mixed with a considerable amount of dust. Pure snow, having a normal structure, is found only in sections within forests. Exceedingly uneven deposition of the snow is observed.

Charts of the depth, the density, and the water equivalent of the snow cover were plotted from the data of snow surveys by triangulation on open country, in 1935-1953, for the southeastern part of the West Siberian Lowland (the Tomsk and Novosibirsk Regions, the level part of the Altai region, and the extreme north of the Kemerov Region). The many-year average of the maximum depth of the snow cover is from 22 to 50 cm in steppe and forest-steppe zones, and from 50 to 70 cm in the forest zone. Accordingly, the many-year averages of the water equivalent of the snow cover in the ten-day period with maximum snow depth are from 70 to 110 mm in steppe and forest-steppe zones, and from 100 to 140 mm in the forest zone. The norms of snow depth and water equivalent of the snow cover may be almost double those given above in snowy years, and half to a third less in years with sparse snow. The mean density of the snow cover in the period when maximum depth is attained varies in the forest zone from 0.19 to 0.22, and in steppes and forest-steppes from 0.22 to 0.31 g/cm³.

The mean density of the snow cover is 0.27-0.30 g/cm³ in North Kazakhstan, and 0.30-0.35 g/cm³ in central and West Kazakhstan. The snow-measuring stations from which readings were used between 31 m and 675 m above sea level. The water equivalent of the snow cover increased with the altitude of the station by 15-17 mm per 100 m height on the average.

The largest water equivalents of the snow cover at the beginning of the thaw are observed in the higher areas of the investigated region of Kazakhstan, where they reached 100-120 mm; on plains they decrease to 50-60 mm, and in the southwest they may be as low as 30 mm.

The average maximum depth, and daily values of the depth, density and water equivalent of the snow cover for each station from data of snow surveys by triangulation at 52 stations in Lower Volga region were obtained over periods of 6 to 10 years (1926-27/1936-37). The many-year averages of the maximum depth of the snow cover in this district decrease in a north-south direction to 10 cm. The many-year average density of the snow varies over this area, at the end of February, from 0.23 to 0.32; during the thaw it increases to 0.40. The maximum winter water equivalents of the snow cover are observed in the northern part of the area, where they attain average values of 80-85 mm, with a maximum at Karabulak (112 mm), decreasing southward down to 25-50 mm.

While the forests of Byelorussia are characterized by a relatively even deposition of snow and by its density (0.19-0.22 in forests and 0.25 in fields), the Soviet Arctic is distinguished by an extremely irregular distribution of the snow cover and a relatively higher density. Thus, the snow survey carried out in April 1934 on Dickson Island showed that the maximum depth of the snow cover (1982 cm) was 5.2 times the average value (35 cm), while the minimum depth (4 cm) was only 11% of the average value.

The depth of the snow in the Arctic is generally quite small, because of the small amount of precipitation. At the end of the winter the average depth varies between 10 to 80 cm, according to the data of surveys at 15 stations. In different years this value may vary considerably, and be much greater or much smaller (by a factor of 6 or more). This is caused by transport processes due to the strong winds, and not because of changes in the amount of the precipitation, this being fairly constant. The snow density in the Arctic is about 0.25 mm at the beginning of winter and increased up to 0.36 mm or even more, at the beginning of thaw.

The onset of the season with a stable snow cover in the European Territory of the USSR (ETS) is abrupt only in specific years. Usually, it is preceded by an intermediate period, from original formation of the snow cover until its definite consolidation. Generally two "temporary covers" may form, during this period, but in certain years up to 5-6 may occur.

The many-year average of the duration of such a "pre-winter" season is 20-30 days. The period of a stable snow cover lasts, on the average, from 80 days in the southwest of the ETS, to 175 days in the northeast. The period of the thaw (in non-homogeneous areas) lasts from 8-9 days in the south to 15-19 days in the north of the ETS. After the first disappearance of the snow in spring new formation of a continuous snow cover (post-winter) is frequently observed.

The total duration of the period from the disappearance of the permanent winter snow cover to the last day on which a snow cover can be observed, is, on the average, from 3-8 to 16-22 days.

The many-year average of the water equivalent of the snow cover in the ETS, at the end of the season with a stable snow cover (in the ten-day

period with the maximum depth for the winter) increases from the southwest to the northeast. It reaches 160-220 mm on the western slopes of the Urals, and on the adjacent areas to the west.

10. TERSKOL PEAK AND ASAU RIVER VALLEY.

Kuvaeva, Sulakvelidze, Chitadze, Chotorlishvili and El'mesov (1967). At the edges of layers of differing snow density, an acute modification in the character of the snow cover takes place; crystals with distinct traits of metamorphism are formed in an above-lying layer. Their appearance is evidently connected with the movement of water vapor. To confirm this, an observation was conducted on Terskol Peak (43°18'N, 42°32'E) from March 13 to May 5, 1959. During that period several layers were distinctly evident in the snow mass: a layer of deep hoar-frost was at the bottom; a layer of fine-grained snow above it (with the density of 0.33 g/cm³), a layer of packed snow (0.4 g/cm³), a snowstorm type fine-grained snow still higher, and a layer of deep hoarfrost of 0.27-g/cm³ density above it.

During the entire experiment, lasting 36 days, the amount of vapor transferred from the layer above the packed snow was twice as much as the amount of water vapor passing from the layer below the packed snow. This means, that on sunny days and in layers of packed snow close to the surface (up to 30 cm) intensive evaporation takes place above it due to solar radiation and conditions are created for the formation of deep hoarfrost.

Experiments were conducted to see how the movement of water vapor affects the density of the snow cover. Evaporation was found to occur in all layers at the bottom of the valley and the volume of evaporation increases toward the snow surface. Within a month, the snow density in that area decreased by 0.003 -0.004 g/cm³. Evaporation constantly occurred in the lower layer of the northern slope whereas the upper layer remained unchanged for a long time. In 25 days the snow density in the lower layer decreased by 0.005 g/cm³, and in the upper by less than 0.001 g/cm³. Less evaporation from the surface layer of the northern slope can be accounted for by the decrease in temperature index in the direction from ground to snow surface. The southern slope presented quite a different picture. For the majority of the observation period the snow mass increased (by 0.001 g/cm³ in 20 days). This was caused by an increase in temperature in the upper layer from solar radiation and on the ground an increase in temperature owing to radiation absorption (the snow [depth] did not exceed 30 cm). As a result, temperature conditions were created inside the snow mass such that water vapor was forced to move from the ground into the snow mass and from the upper layer of snow to the lower.

Evaporation from the lower, 10-centimeter layer in the Azau river valley, on Terskol Peak and the Ice base was measured throughout 1960/61 winter. Average evaporation was equal to that registered in 1958-59 winter. On the whole, during that winter, snow density in the lower 10-cm layer at the altitude of 2200 m above sea level decreased by 1.5-2%, and at an altitude of 3100 m above sea level - by 2%, and at 3700 m altitude by 2-2.5%. The evaporation from the lower layer of the snow mass at different altitude zones changed insignificantly, mainly owing to the fact that the water vapor movement through the lower layers slowed down because of a large accumulation of snow that can be accounted for by longer winter periods at higher altitudes.

The results of the experiment show that there is practically no decrease in snow density owing to evaporation processes in the Elbrus area (up to an altitude of 3700 m above sea level). The snow density decrease caused by 3% evaporation during the entire winter period is more than compensated for by the sinking of the snow cover as the result of which the cohesion between particles in the lower layers remains the same.

11. IRKUTSK.

Rozental (1904). Snow density measurements of freshly fallen snow, and at various depths within the snow cover were made at Irkutsk, USSR, during the winters of 1899, 1902 and 1903. The results of the study showed that the density of fresh snow depended upon the structure and size of the original snowflakes. The density of new snow consisting of large flakes ranged from 0.044 to 0.066 g/cm³; for medium sized flakes from 0.078 to 0.118 g/cm³; and for small flakes from 0.130 to 0.141 g/cm³. The average density of fresh snow for 13 separate storms at Irkutsk was 0.093 g/cm³.

The snow cover densities for the top layer of snow (surface to 5-cm depth) following each new snowfall increased from values of 0.059 to 0.133 g/cm³ on the first day of accumulation to between 0.175 and 0.263 g/cm³ after periods of 5 to 14 days of time. According to consecutive measurements made after five new snowfalls the density increase per day in the top layer of snow cover is approximately 0.01 g/cm³. Deviations from the norm were associated with surface wind speeds of 10 m/sec and/or melting.

Average snow cover profile densities for the three winters of observation at Irkutsk were given as follows: for the layer of surface to 5-cm depth 0.141 g/cm³; for 5 to 10-cm depth 0.214 g/cm³ and for below 10-cm 0.226 g/cm³. The greatest observed density (0.385 g/cm³ occurred twice in 1899 - on 24 January and 18 March. In the first case the sample was taken from a snow pile at a depth of 15 to 27 cm, and in the second case it was taken from the surface during melting. Incidentally, on 18 March 1899 a second snow surface measurement taken in the same area except that the snow cover is almost entirely shaded throughout the day, the density was much lower (0.263 g/cm³).

12. USSR TAIGA.

Kolomyts (1964). In analyzing the data supplied by research studies of the snow cover in the mountainous taiga of the Zabaykal regions over a period comprising the last few winters, and adding to this data an examination of the few published studies on the snow cover in other regions of Siberia and the Far East, we are made aware of the sharp differences between the snow cover in Siberia and the North Eastern regions and that in the European and Maritime. Differences have been revealed not only in the depth of the snow and the length of time it covers the ground, but to even a greater extent in its stratification and density (general as well as strata density), and accordingly in the nature of the evolution of the snow mass. This leads to the thought that the winter conditions of one or another geosystem can be determined by the character of the development (the stratification) of the snow cover. By the latter we mean that combination of elements, the stratification, depth and density of the snow, which reflect the entire complex of winter conditions of the given regional or cartographical unit.

Of all parametric data, that which is of decisive importance is the stratification of the snow, connected as it is with the most complicated and many-sided aspects of the winter regime of the given geosystem.

The most significant provincial peculiarities of winter in various regions of the USSR taiga conform to types and subtypes of development. The criteria for determining these types are: 1) seasonal changes in the snow mass determining the character and speed of its melting, and 2) the tendency and intensity of the metamorphism of the snow mass, reflecting the exchange of heat between various components of the landscape in winter. For USSR territories characterized by plains and low mountains, three basic types of the snow mass have been determined; we propose calling them: 1) syngenetic, 2) epigenetic, and 3) transitional types.

1. The syngenetic type is characteristic of large regions with a moderate continental winter, where the accumulation of snow continues throughout the entire winter season (as in the taiga of the European part of the USSR, the Maritime Region and the eastern coast of Kamchatka). The snow cover offers summarized information about the winter regime, simultaneously (synchronous) with its own accumulation, and therefore may be considered a peculiar kind of chronicle of the winter season. The order of the basic stages of winter is recorded by marked changes in the snow mass: in its horizon, stratification, and simultaneous growth. In this lies the essence of its syngeneses (analogous with permafrost).

The snow cover is deep: the maximal mean depth is from 50-70 cm and more. With moderate frosts this creates a relatively low temperature gradient within the snow mass, an average of 0.1-0.3°C/cm and accordingly, a low degree of secondary super-crystallization of the mass. The dominant process in these conditions is a breaking down of the structure, i.e., firm snow formation: the transformation of initial idiomorphic crystals into round seed of various sizes. This process is attended by the settling of the snow and an increase in density; it takes place most energetically at times of thaw. The result is that by the end of winter, the density of the snow near the southern boundaries of the European taiga reaches 0.27-0.35 g/cm³ and in the northwest 0.40-0.45 g/cm³. In accumulations of old snow we find the formation of secondary crystals - deep hoarfrost (snow sand), lying off to one side of the general development of the snow.

The syngenetic type of layer may be divided into two subtypes: depending on the degree of the development of texture and structural characteristics of the snow cover, i.e., the correlation of the rate of snow accumulation on one hand and the intensity of its evolution on the other. The layers in the first subtype are distinguished mainly by their texture characteristics. This subtype is most characteristic of the syngenetic type and is particularly often encountered in the taiga area of the Primorie. It is formed in the process of periodic (spasmodic) snow accumulation as a result of frequent alternation of snow falling and snow storms, frosts and melting periods. The second subtype is not characterized by such clear distinction of layers which are recognized mainly by their structural peculiarities. This is an initial shift to the second type. It may be regarded as the beginning of the modification toward the second type. This subtype is spread in the north of the ETS and in the Priuralye as well as along the Pacific shoreline of Kamchatka.

2. The epigenetic type is widely spread in the taiga areas with severe and comparatively snow-free winters and extremely uneven seasonal solid precipitation in Middle Siberia, Jakuria, central mountainous areas and depressions of the Altai, the Sayan and Zabaykalye regions as well as on the left-bank of the middle reaches of the Amur. In these regions the main bulk of snow falls during the first 1.5-2 months of winter (before a stable anticyclone period of air mass sets in); that is why the peculiarities of the winter season develop in an already formed snow layer of one age-period (epigenesis). The height of snow does [not] exceed 20-30 cm; this factor with the continued domination of low temperatures contributes to high temperature [gradient] in the snow cover (of up to 1.5-2°C/cm).

Slight transportation of snow during snow storms and almost no melting periods help form the initial homogeneity of the snow. The formation of the snow cover is closely connected with winter mass- and thermo-exchange between ground layers and the above-ground layer of air as a result of which these processes predominate in the vertical direction. Extremely high activity of the latter accelerates structural metamorphosis of snow under the action of sublimation. The following accompanying phenomena were registered: intensive growth of crystals, the appearance of columnar-type formations and of secondary (epigenetic) differentiation of the snow cover, connected, evidently, with sharp, periodic (daily) temperature fluctuations and deep penetration of direct solar radiation into snow cover. By the end of winter the entire body of snow has modified its crystal structure completely and is then almost entirely composed of deep hoar crystals reaching in diameter from 3 to 5 mm. There are no sharp demarcations between layers; the differences are of a structural nature and can be traced only by the size of the crystals, the cohesive force (solidity) volume weight and other physical and mechanical characteristics of snow. Judging by isolated data, snow cover of the Northeastern taiga regions of the Soviet Union bears a similar structural resemblance.

The distribution of snow in the forest-steppe regions of Zabaykalye and on forestless areas of Central Jakutia as well as in the depressions of Verkhoyanye, Priabaykalye and Stanovoye upland has its own peculiarities. Energetic transportation of snow in the beginning of winter stimulates the initial differentiation of the snow cover which continues to take place. The mirroring effect of the upper layers affected by snow storms upon the migration of water vapor stimulates constructive metamorphism of the below lying layers and this makes it possible to classify this modification process of epigenetic type of deposition as a subtype.

The epigenetic type as compared with other types has the lowest snow density. In Zabaykalye and Priamurye the density by the end of winter comprised from 0.12 to 0.15 and to 0.18-0.20 g/cm³ and in the Northeast of the Soviet Union - from 0.15-0.17 to 0.20 g/cm³.

3. The transitional type is predominant in the taiga region of the West Siberian lowland, along the Okhotsk shoreline and partly on the left-bank of the Amur lowlands. The height of the snow cover is quite large here - 40-70 cm. The snow density ranges from 0.19-0.20 to 0.23-0.25 g/cm³. In structure and composition of snow this type is evidently closer to the second subtype of the syngenetic type, but differs from it by a lesser initial differentiation of the mass and by clearer traces of the evolution of constructive metamorphism.

Alpine regions of Siberia and the Far East are distinguished from the plain regions of taiga by differentiation in the characteristics of snow cover not only in the horizontal layers but in vertical as well. The snow cover in the foothill area as well as in the center of mountains develops on the whole in the same way as in surrounding taiga areas of the plains and depressions. However, "taiga uplands" including the upper parts of the mountain taiga and bare spaces is characterized by very specific snow distribution affected by transportation of snow during snow-storms as well as by the factor of snow accumulation that sets in during pre-spring and spring months. Two different layers develop in the stratigraphic column of snow; the lower layer is formed by intensive snow falls at the beginning of winter. By winter it reflects (in its structural peculiarities) the results of metamorphism processes in the snow cover taking place during the mid-winter months when snow falling sharply decreases (epigenesis). The upper layer of snow cover determines the spring and summer seasons (syn-genesis). In this way the snow cover has two segments and may be classified as alpine polygenetic (episyn-genetic) type with the predominance of constructive metamorphism of crystals.

13. USSR SNOW SURVEYS.

Vershinina (1969). This paper examines the accuracy in determining snow water equivalent at a point, over a course and over an area, from data of special snow surveys carried out by the State Hydrological Institute in different physico-geographical regions of the Soviet Union. The data used were collected by continuous, landscape-course and course snow surveys during 1963-1968.

An estimate of the instrumental errors in determining snow water equivalent is shown in the following table [Table A6]:

Table A6. Instrumental errors in determining snow water equivalent.

Snow density, g/cm ³	Snow depth, cm	Snow-density measurement error g/cm ³	Water equivalent of snow, mm	Error in determining water equivalent of snow, %
0.10	5	0.040	5	40
	10	0.020	10	20
	15	0.013	15	13
	20	0.010	20	10
	30	0.007	30	7
	40	0.005	40	5
	50	0.004	50	4
0.20	5	0.060	10	30
	10	0.030	20	15
	15	0.020	30	10
	20	0.015	40	8
	30	0.010	60	5
	40	0.007	80	4
	50	0.006	100	3

Table A6 (cont'd). Instrumental errors in determining snow water equivalent.

Snow density, g/cm ³	Snow depth, cm	Snow-density measurement error g/cm ³	Water equivalent of snow, mm	Error in determining water equivalent of snow, %
0.30	5	0.080	15	27
	10	0.040	30	13
	15	0.027	45	9
	20	0.020	60	7
	30	0.013	90	4
	40	0.010	120	3
	50	0.008	150	3

The accuracy in determining snow water equivalent over courses depends on the nonuniformity of snow distribution and the number of measurement points. Studies show that the errors in determining maximum snow water equivalent in northern and northwestern areas of European USSR amount to 5-6 %, and in central and southern areas to 7-8%.

14. KAMCHATKA.

Vinogradov (1964). The eastern littoral includes the zone situated along the eastern coast of the Kamchatka peninsula. The winter period is characterized by relatively high air temperatures and heavy precipitation. The latter is often accompanied by strong winds of 29-40 m/s. As a rule, the determination of snow cover is carried out simultaneously throughout the entire region: in 1960 it took place during the last ten days of October. In the southern area of the region, within the Avachin depression, the nonuniformity in distribution of snow cover characteristic of Kamchatka is observed. The smallest depth of snow cover during the period of maximum accumulation is recorded in the area of the villages Yelizovo and Koryaki (50-60 cm). At the same time, the basin among the medium-height ridges where the village of Paratunka is located is characterized by the deepest snow cover, reaching 220-240 cm in hard birch forest and 160-170 cm in open areas without forestation. The density of the snow gradually increases from 0.10 to 0.12 g/cm³ in the beginning of the winter to 0.35 to 0.40 g/cm³ at the time of maximum snow accumulation.

On the upper relief layer, on plowed land, accumulation occurs extremely slowly. A significant portion of the snow is carried away by wind, during the first ten days of December the depth of snow cover was 29 cm. A snow mass section consisted of two strata: freshly fallen snow, and weakly crystallized, finely granulated snow. At the same time, in areas covered with an undergrowth of hard birch, the depth of snow cover reached 57 cm at a density of 0.20 g/cm³. In hard birch thickets, the depth of snow cover was 168 cm, which was the maximum amount of snow accumulation on a flat surface in the area of Petropavovsk. The density of the snow is 0.29 g/cm³, and the water supplies in the snow mass, 487 mm. In an open area occupied by a field, the depth of snow cover was 90 cm. As a result of great wind compaction, density was 0.31 g/cm³, and the water supplies, 279 mm.

Wind conditions exert a great influence on the distribution of snow cover. In areas where prevailing wind is strong (more than 29 m/s) and northeasterly, the most snow-covered slope is that with southern exposure, where accumulation takes place more extensively due to the snow masses blown from the ridge. The maximum snow cover was recorded in the middle of March, when its depth on a southern slope was 128 cm, its density 0.31 g/cm³, and its water supplies 396.8 cm.

Snow cover on flat offshore bars is subject to extreme variation and shifting. These areas are characterized by extremely deep snow cover; in the period of maximum snow accumulation it has a depth of 52 cm. As a result of strong winds often encountered on the littoral, the density of snow cover approaches its maximum in February, and in the middle of March it was 0.34 g/cm³. In a snow mass section a series of strata of finely granulated, friable snow and hard-frozen snow was observed. In open areas of the maritime plain, in the base of a snow section taken during the beginning of April was observed ice similar to that on the gulf of the sea. During the period of maximum snow accumulation in the middle of March, the depth of snow cover on the surface of the gulf was 38 cm, and in open areas of the maritime plain, 75 cm. Such a difference may be explained by the more extensive transfer of snow on the surface of the gulf. The maximum snow cover was found in occasional groups of willow that favored the accumulation of snow. The density of snow in the willow clusters was 0.30 g/cm³, somewhat lower than that in open areas. Dissipation of the snow cover begins during the end of May.

During the period of maximum snow accumulation (from the second half of February to the second half of March) the greatest differences in depth of snow cover were to be found in the different standard areas. The maximum depth of snow cover, 182 cm, was recorded on a slope of north-north-western exposure with an angle of inclination of up to 20° under cover of a hard birch forest. The smallest depth was observed on the flat, open surface of the Kamchatka River, and was equal to 155 cm. Density of the snow cover continued to increase, a certain degree of stabilization being observed in open areas at 0.30-0.31 g/cm³. During this time the density in areas of forestation increased from 0.20 to 0.30 g/cm³. Water supplies increased constantly as a function of the depth and density of the snow cover; although during the period of snow accumulation the water supplies in open areas were lower than those in forested areas, during the period of maximum snow accumulation the situation was reversed, and in open areas they were somewhat higher.

The structure of the snow mass in the period of maximum snow accumulation in the different standard areas had certain common features. In open areas, a layer of buried snow from snowstorms was observed that was absent under cover of forest. In the period of snowfall (April-May), extensive fusion of the snow crystals and a decrease in the depth of snow cover occurred as a result of average daily temperatures exceeding 0°C. In open areas the snow cover dissipated by the beginning of June. Considerably slower snow thaw occurred under cover of hard birch forest, where dissipation of the snow cover occurred during the second half of June. The highest values of snow cover density, 0.45 g/cm³, were recorded for open areas; in hard birch forest the density did not exceed 0.43 g/cm³. During the second half of April complete fusion of the snow granules occurs, and

the snow mass acquires a uniform structure. In the beginning of May, throughout its profile, regardless of the depth of snow cover, the snow mass consists of fused snow; at its base is observed a layer of water.

The climate of the Esso region has features of continental nature, expressed in terms of a relatively dry and hot summer and a cold winter. Average monthly air temperatures are near or below -15°C . In December and January it is not uncommon for the air temperature to fall below -40°C at night. Wind conditions have no effect on redistribution of snow, since weak winds with an average velocity of 2-4 m/s prevail. The maximum wind velocities do not exceed 10 m/s, and these occur only infrequently and are of short duration. On October 20 the depth of snow cover was 12 cm at a density of 0.12 g/cm^3 . Accumulation over the course of the first month of its existence took place extremely slowly, and on November 20 in open areas it had a depth of 16 cm at a density of 0.15 g/cm^3 .

The accumulation of snow cover in the Esso region occurred relatively uniformly and concluded in the middle of January. By this time the depth of snow cover had stabilized, and subsequently it underwent very little change prior to the beginning of snow thaw. During the course of the winter period the snow cover in forested areas is 10 cm greater than that in open areas, this difference remaining fairly constant. In forested areas constant accumulation of snow cover occurred until January 20. Subsequently, during the course of February and March, no increase in depth is observed, but fluctuations of up to 7 cm are recorded due to wind effects. By January 20 the depth of snow cover was 58 cm at a density of 0.17 g/cm^3 . Within a month its depth grew to 63 cm, and its density also increased to 0.19 g/cm^3 . March was characterized by a decrease in the depth of snow cover and stabilization of its density at $0.21\text{--}0.22\text{ g/cm}^3$. During the period of maximum snow accumulation the depth of snow cover reached 71 cm, which is the maximum figure for this region. The density of snow cover at this point was 0.23 g/cm^3 . Subsequently, during snow thaw the density increases to 0.26 g/cm^3 as a result of dampening of the snow mass.

In open areas accumulation of snow cover occurred in a manner similar to that described during February and March. Extensive accumulation does not occur, but the depth of snow cover does not remain constant; it increases slightly. The absence of strong winds capable of transferring snow results in a relatively uniform distribution of snow cover. In the period of maximum snow accumulation, the snow cover in open areas has a depth of 59 cm. The density of snow cover remains constant at 0.22 g/cm^3 as of the beginning of February. During snow thaw the density increases to 0.24 g/cm^3 .

In the middle reaches of the Anavgaj River the right edge of the gently sloping valley rises to the planated lava surface of the Central Range in the area of the Anaun volcano. Here a larch forest extends up to the upper boundary of the forest belt, and a belt of hard birch occurs. The snow cover is friable and occurs relatively uniformly. On a slope of eastern exposure with an angle of inclination of 10° , under cover of the larch forest the depth of snow cover is 82 cm at a density of 0.30 g/cm^3 . Stable snow cover was formed on October 29, and its principal accumulation occurred during November and the first half of December. On December 20 in

a level, open area the depth of snow cover was 54 cm at a density of 0.25 g/cm³. Within the accumulation period the depth of snow cover remains fairly constant. During the first ten days of February, after almost two months' existence of snow cover, its depth underwent hardly any change. On the surface of the hollow, the depth of snow cover was 54 cm at a density of 0.26 g/cm³.

During the period of maximum snow accumulation, the most snow covered types of relief are depressed areas. In the hollow, the depth of snow cover was 72 cm at a density of 0.22 g/cm³. Elevated areas exposed to the wind had depths of snow cover approximately half as great; on the weakly defined divide it was 47 cm, at a density of 0.25 g/cm³. Snow melting occurred at a reduced rate in April and very rapidly in May. On the level surface of the divide, from April 20 to April 30 the depth of snow cover decreased from 31 to 24 cm, and by May 10 no snow cover remained. In the hollow snow thaw occurred more slowly, and during the same period the snow cover decreased respectively from 58 to 49 cm and from 49 to 8 cm.

In the area of the settlement Ust-Bolsheretsk snow cover existed for 194 days in 1960-61. The maximum depths of snow cover in depressed areas were 72 cm, and in elevated areas, 47 cm. As a result of wind action the density of snow cover approaches its maximum, registering 0.25 g/cm³, in the middle of December. The area of the settlement Sobolevo, located to the north of Ust-Bolsheretsk, also has a plain relief. Stable snow cover was formed on October 20 after prolonged snowfall. Formation and accumulation occurred at a fairly rapid pace: on October 31 the depth of snow cover was 29 cm at a density of 0.17-0.18 g/cm³, and its distribution was uniform. By November 20, the depth of snow cover had increased to 42 cm in open areas and 47 cm in hard birch forest; its density remained unchanged. Subsequently, a slower increase in the depth of snow cover takes place, but its density increases rapidly. On December 20 the depth of snow cover in open areas is 58 cm at a density of 0.25 g/cm³, and in forested areas, 64 cm at density of 0.22 g/cm³. A fairly constant increase in the depth of snow cover occurs until the middle of January.

In the period of maximum snow accumulation, the most snow-covered areas are those with forestation. On the level of a terrace covered with hard birch forest was recorded the maximum depth of snow cover for this area - 92 cm, at a density of 0.26 g/cm³. A level surface with herbaceous grassland has a depth of snow cover of 77 cm at a density of 0.27 g/cm³. Snow thaw occurs rapidly. In open areas the snow cover dissipated on May 10, and in forest, on May 17. Snow cover in the area of the settlement Sobolevo existed during the winter period 1960-61 for 202 days in open areas and 209 days in forested areas. The maximum depth of snow cover on a level surface covered by hard birch forest was 92 cm, and in open areas, 77 cm. The area of the settlement Sobolevo is characterized by greater snow cover than the area of Ust-Bolsheretsk, located to the south. In north-western Kamchatka a decrease in depth of snow cover occurs from south to north.

The area of the settlement Tigil is more distant from the Sea of Okhotsk than the areas of the western littoral described above and differs from them by virtue of its hilly and ridged relief. Snow cover was formed on October 10. Its accumulation occurred during the second ten days of October. In its uniform occurrence a depth of snow cover was observed of

27 cm at a density of 0.11 g/cm^3 . In November, during an increase in depth of snow cover, differences were recorded in its distribution in different areas: in fields, 43 cm, and under cover of hard birch forest, 50 cm. Accumulation of the snow cover in this area concluded in the middle of February.

15. URAL MOUNTAINS.

Khodakov (1967). The Ural Mountains to the north of 60° northern latitude are still largely unsettled, and the hydrometeorology of the region has received little study. However, its natural resources, especially its useful mineral and hydraulic power resources, promise this region an important role in the economics of the country in the near future. National economic development requires a thorough knowledge of the physical and geographical conditions of this region.

The rate of snowstorm transport determines to a considerable extent the structure and properties of the snow. In the bald mountain zone a large portion of its mass is made up of finely- and medium-granulated, dense snow formed from densely packed, fine particles of snowstorm snow. The density of these types of snow during the second half of the winter is usually $0.35\text{-}0.42 \text{ g/cm}^3$. If the depth of the snow is slight, a layer forms in the lower portion of its depth in which it becomes friable. In deep snow masses the density increases continuously with increasing depth, reaching $0.55\text{-}0.58 \text{ g/cm}^2$ at a depth of 10 m. In the forest zone, finely granulated snow with a density of $0.25\text{-}0.30 \text{ g/cm}^3$ is observed only in the upper third of a section. The remainder of the section is occupied by coarsely granulated snow and deep-seated hoarfrost with a density of $0.15\text{-}0.25 \text{ g/cm}^2$. The very dense snow of the unforested Urals has exceptional mechanical strength. The rupture plasticity for dense, finely granulated snow is $120\text{-}250 \text{ g/cm}^3$, and for wind panels (density $0.42\text{-}0.45 \text{ g/cm}^3$), up to 600 g/cm^2 ; shear plasticity is $500\text{-}1500 \text{ g/cm}^2$ and up to 5600 g/cm^2 ; the viscosity coefficient is 1×10^{13} poise and 0.8×10^{12} poise, respectively. The high strength of bald mountain zone snow produces relatively weak development of avalanche activity. Winter avalanches are small in volume, consist of fresh snowstorm snow, and are confined to steep, leeward slopes of mountain plateaus. In the spring, when the snow warms up and melts a little, its strength falls sharply, and the number of avalanches increases.

Our measurements have shown that the coefficient of thermal conductivity of the snow, obtained from observations made on location, corresponds closely to that calculated by the Ab'el's equation. This means that at the snow density 0.2 g/cm^3 characteristic of the forested zone, the thermal conductivity at equal snow strengths and heat flows is one-fourth that at the density 0.4 g/cm^3 characteristic of the bald mountain zone. Since the balance of the surface heat balance varies gradually from south to north, it is evidently possible only by means of a sharp increase in the snow's thermal conductivity to explain the observed sharp increase in strength of permafrost in the unforested zone.

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APPENDIX B
 AVERAGE 5-MONTH (NOVEMBER THROUGH MARCH) WIND SPEED (M/S)
 FOR STATIONS IN THE SOVIET UNION*.

<u>Station</u>	<u>Coordinates</u>		<u>Elevation</u> (m)	<u>No. months</u> <u>Temp. < 0°C</u>	<u>Ave. 5-mo.**</u> <u>(Nov-Mar)</u> <u>wind speed</u> <u>(m/s)</u>
	<u>Lat(N)</u>	<u>Long(E)</u>			
Abakan	53° 43'	91° 26'	-	-	2.6
Abramovskiy	66 25	43 15	24	-	7.0
Achinsk	56 18	90 31	235	5	4.8
Adamovka	51 31	59 57	285	5	3.7
Ado Tymova	51 08	142 40	64	5	1.4
Agata	66 55	93 28	244	7	1.9
Aginskoye	55 15	94 55	-	-	2.5
Ak-Baytal	43 09	64 20	232	3	3.6
Aksenovо Zilovskoye	53 04	117 29	707	7	1.6
Ak-Tel	51 10	93 25	-	-	1.0
Aktyubinsk	58 37	57 09	227	5	4.1
Aldan	58 17	125 22	682	7	3.0
Aleksandrov-Gay	50 09	48 33	25	5	5.0
Aleksandrovsk Sakhalinskiy	50 54	142 10	30	6	4.6
Aleksandrovskoye	60 26	77 52	60	7	3.4
Alma-Ata	43 14	76 56	847	3	0.9
Amangel'dy	50 08	65 14	142	5	4.3
Amderma	69 46	61 41	53	7	8.5
Anadyr	64 47	177 34	61	7	7.9
Anavgay	56 06	158 53	-	7	2.0
Anuchino	43 58	133 04	189	5	2.0
Aral'sk	46 47	61 40	56	5	5.2
Arkhangel'sk	64 35	40 30	13	-	4.5
Arkharо	49 25	130 07	142	5	2.7
Arshan	51 54	102 27	-	6	1.2
Artemovsk	54 21	93 26	472	6	1.0
Astrakhan'	46 16	48 02	18	3	4.7
Atbasar	51 49	68 22	308	5	5.5
Ayaguz	47 56	80 23	659	5	3.9
Ayan	56 27	138 09	9	6	2.8
Bakchar	57 01	82 05	-	6	3.8
Bakhta	62 28	89 00	-	-	2.8
Balkhash	46 54	75 00	423	5	5.3
Balakhta	55 23	91 37	-	-	2.9
Barabinsk	55 22	78 24	120	5	5.6
Barouzin	53 37	109 38	486	6	2.1
Barnaul	53 20	83 42	196	5	5.8
Batamay	63 31	129 26	78	7	1.8
Baturino	57 48	85 12	-	6	3.1
Bayevo	53 17	80 46	-	5	4.2
Baykit	61 41	96 25	-	-	0.7
Bektan Ata	47 27	74 49	616	5	5.9
Bel'tsy	47 46	27 56	-	-	3.4
Berezovo	63 56	65 03	20	7	3.4

*See text for various data sources.

**Compilation of a uniform period and/or length of record was not possible. Records dating from 1948 to 1971 were mostly included in the survey, and ranged in length of from 3 to more than 15 years.

Station	Coordinates		Elevation (m)	No. Months Temp. < 0°C	Ave. 5-mo. (Nov-Mar) wind speed (m/s)
	Lat (N)	Long(E)			
Betpak-Dala	46° 01'	70° 12'	-	5	4.4
Big Diomedes Island	65 47	169 05(W)*	39	7	5.7
Birilyussy	57 07	90 32	167	5	2.6
Biysk	52 34	85 15	230	5	4.9
Blagoveshchensk	50 16	127 32	136	5	2.8
Boguchar	49 56	40 34	82	4	3.3
Bogopol'	44 15	135 28	50	5	2.8
Bogotol	56 10	89 35	-	5	4.8
Bogorodskoye	52 23	140 28	31	6	2.2
Boguchany	58 23	97 29	134	6	2.8
Bol'shaya Murta	56 55	93 07	-	-	2.2
Bol'shoy Porog	65 39	90 05	-	-	1.9
Bol'shenarymskoye	49 12	84 31	404	5	1.4
Bolgrad	45 40	28 37	81	2	4.2
Bolotnoye	55 41	84 23	-	5	4.5
Bomnak	54 43	128 50	352	7	2.3
Borovichy	58 24	33 55	-	-	3.5
Borzya	50 23	116 31	684	6	2.6
Bratsk	56 21	101 55	327	7	2.2
Brest	52 07	23 41	143	3	3.9
Bryansk	53 20	34 14	162	5	5.0
Bukhta Pronchishchevoy	75 40	113 11	-	7	4.6
Bukhta Solnechnaya	78 10	102 50	-	-	6.4
Bukhta Ugol'naya	63 03	179 19	-	7	10.6
Buolkalakh	72 56	119 50	-	7	5.8
Buor-Yuryakh	68 11	145 55	-	7	1.2
Burlyu Tobe	46 35	79 06	351	5	3.2
Buynaksk	42 49	47 07	472	2	3.5
Byssa	52 21	131 17	-	6	1.0
Chadan	51 22	91 27	-	6	1.2
Chadopets	58 40	98 51	150	7	2.3
Chara	56 55	118 22	703	7	0.6
Charkov	53 43	90 22	524	5	2.9
Chekunda	50 49	132 10	234	5	1.0
Chelyabinsk	55 09	61 24	230	5	4.2
Chemdal'sk	59 38	103 20	-	-	1.1
Cherdyn'	60 24	56 31	206	6	4.5
Cheremkhovo	53 09	103 05	-	6	2.2
Cherepovets	59 07	37 57	131	5	5.3
Chernovtsy	48 18	25 56	-	3	3.0
Chimbay	42 57	59 49	66	3	2.7
Chirgalandy	50 36	97 15	-	-	0.7
Chita	52 01	113 20	684	-	1.5
Chokurdakh	70 37	147 53	47	7	4.3
Chul'man	56 50	124 52	664	7	1.1
D'elind'e	65 15	135 36	-	7	0.7

<u>Station</u>	<u>Coordinates</u>				<u>Elevation</u> (m)	<u>No. Months</u> <u>Temp. < 0°C</u>	<u>Ave. 5-mo.</u> <u>(Nov-Mar)</u> <u>wind speed</u> <u>(m/s)</u>
	<u>Lat (N)</u>	<u>Long(E)</u>					
Dambuki	54° 21'	127°	38'	270	6	2.0	
Dikson	73 30	80	25	18	7	7.9	
Dnepropetrovsk	48 27	35	03	142	4	4.4	
Dno	57 50	29	59	-	-	3.5	
Dolgiy Most	56 45	96	48	256	7	2.4	
Dolinsk	47 20	142	48	42	5	3.6	
Dudinka	69 24	86	15	28	7	5.3	
Duki	51 43	135	54	131	5	1.2	
Dzerzhinskoye	56 50	95	13	-	5	2.3	
Dzhalal Abad	40 55	72	57	771	1	1.6	
Dzhambeyty	50 15	52	34	34	5	5.8	
Dzhambul	42 51	71	23	641	3	3.1	
Dzhardzhan	68 44	124	00	46	7	3.1	
Dzhezkazgan	47 48	67	43	345	5	4.1	
Egvekinot	66 21	179	07(W)	-	7	6.3	
Ekimchan	53 04	132	56	491	7	0.8	
El'gen	62 48	150	41	303	7	1.1	
El'ton	49 08	46	51	11	4	5.1	
Erzin	50 15	95	10	-	-	1.3	
Fort Shevchenko	44 33	50	15	20	1	6.9	
Gdov	58 46	27	48	-	-	5.0	
Gizhiga	62 03	160	30	10	7	5.5	
Glazov	58 09	52	40	-	5	3.7	
Gol'chikna	71 43	83	36	10	7	6.4	
Gor'kiy	56 20	43	59	-	5	5.8	
Gora Dogdynyn	63 36	104	11	769	7	2.3	
Goryachinsk	52 59	108	18	-	6	4.0	
Goryun	51 16	136	35	78	6	2.2	
Grossevichi	47 58	139	32	36	5	4.3	
Groznyy	43 21	45	41	124	2	2.6	
Gudauri	42 28	44	28	2133	-	3.1	
Guga	52 42	137	32	50	6	0.7	
Gur'yev	47 07	51	51	21	3	5.7	
Gyda	70 56	78	25	4	7	6.2	
Idrinskoye	54 21	92	07	-	-	2.3	
Icha	55 42	155	38	4	5	5.4	
Ichera	58 32	109	47	-	7	1.5	
Idritsa	56 19	28	54	139	5	4.5	
Igarka	67 28	86	34	29	7	4.8	
Ika	59 18	106	24	-	7	1.4	
Il'inskiy	47 59	142	12	18	5	5.2	
Ilrney	67 20	168	14	425	7	1.5	
Irgiz	48 37	61	16	114	5	4.7	
Irkutsk	52 16	104	20	468	-	2.2	
Isit	60 49	125	19	118	7	2.4	
Ivano-Frankovsk	48 54	24	42	276	3	3.7	

<u>Station</u>	<u>Coordinates</u>		<u>Elevation</u> (m)	<u>No. months</u> <u>Temp. < 0°C</u>	<u>Ave. 5-mo.</u> <u>(Nov-Mar)</u> <u>wind speed</u> (m/s)
	<u>Lat(N)</u>	<u>Long(E)</u>			
Ivanovo	57° 10'	40° 58'	138	5	5.1
Ivdel'	60 41	60 26	100	6	1.9
Kachug	53 58	105 52	-	7	1.9
Kagul	45 54	28 11	-	-	4.2
Kaliningrad	54 42	20 37	27	3	4.5
Kalmykovo	49 03	51 52	4	4	4.9
Kamenka	58 33	95 51	-	5	2.4
Kamen Rybolov	44 43	132 04	75	5	3.4
Kamen' na Obi	53 47	81 20	-	5	5.5
Kamenskoye	62 29	166 13	-	7	6.1
Kandalaksha	67 08	32 26	15	6	4.1
Kanevka	67 08	39 40	-	-	2.6
Kanin Nos	68 39	43 18	4	6	9.4
Kansk	56 13	95 41	203	5	3.6
Kapustin Yar	48 35	45 43	9	4	4.3
Kara-Kul'-Bulak	46 47	64 40	-	5	4.4
Karaganda	49 48	73 08	555	5	4.9
Karaginskiy	59 00	163 55	6	6	7.9
Kara-Kem	52 23	92 24	-	-	2.6
Karam	55 09	107 37	-	7	1.3
Kargasok	59 03	80 53	-	7	4.3
Karsakpay	47 50	66 45	488	5	4.1
Karatuz	53 -36	92 53	-	-	1.8
Karaul	70 06	83 08	-	-	7.1
Kaunas	54 53	23 53	75	4	4.4
Kaynar	49 11	77 25	-	5	4.2
Kazach'ye	70 45	136 13	21	7	3.5
Kazachinskoye Exp. Field	57 42	93 17	93	7	2.4
Kazachinskoye	56 42	107 36	-	7	1.3
Kazalinsk	45 46	62 07	68	3	3.7
Kazan'	55 47	49 11	64	5	5.3
Kedon	64 08	159 14	682	7	1.1
Kellog	62 29	86 19	-	7	2.5
Kem'	64 59	34 48	9	6	5.3
Kemerovo	55 23	86 04	153	5	5.0
Kerbo	62 44	101 06	-	-	0.7
Kezhma	58 59	101 09	-	5	2.5
Khabarovsk	48 31	135 10	71	5	4.4
Khadama	53 57	98 51	-	7	1.4
Khanty Mansiysk	60 58	69 04	39	6	4.2
Khar'kov	49 56	36 17	152	4	4.7
Kharlovka	68 47	37 20	12	6	8.8
Khatanga	71 58	102 30	24	7	4.5
Khatyrka	62 03	175 16	152	7	9.0
Khilok	51 21	110 30	801	7	2.2
Khmel'nitskiy	49 24	26 59	195	4	4.1

Station	Coordinates		Elevation (m)	No. months Temp. < 0°C	Ave. 5-mo. (Nov-Mar) wind speed (m/s)
	Lat(N)	Long(E)			
Kholm	57° 09'	31° 11'	-	-	3.4
Khonuu	66 27	143 14	160	7	0.3
Khorog	37 30	71 30	2080	3	0.5
Khosedakhard	67 02	59 24	81	7	4.4
Kingisepp	58 15	22 28	-	4	6.3
Kirensk	57 46	108 08	256	7	2.2
Kirov	58 39	49 37	164	5	4.9
Kirovograd	48 29	32 15	148	3	5.0
Kirovsk	67 37	33 40	-	-	3.4
Kirovskiy	45 06	133 30	-	5	2.0
Kiselevsk	54 00	86 39	-	5	3.8
Kishinev	47 00	28 50	-	-	4.4
Kiyev (Kiev)	50 24	30 27	178	4	4.8
Kochenga	55 55	104 06	-	7	1.0
Kochumdek	64 24	92 48	110	7	1.5
Kokchetav	53 17	69 21	228	5	5.8
Kolpashevo	58 18	82 54	75	7	3.7
Kommunar	54 20	89 18	-	-	4.3
Korenov	73 33	107 10	-	7	5.0
Korf	60 21	166 00	-	7	7.3
Koschagyl	46 48	53 55	-	4	7.2
Kosh Agach	50 01	88 44	1758	7	0.9
Kosistyy	73 38	109 40	-	-	5.5
Kosukhino	71 19	149 23	-	7	4.6
Kovdor	67 34	30 24	-	-	2.9
Kozyrevsk	56 03	159 51	-	5	2.6
Kraskino	42 43	130 47	11	5	4.5
Krasnodar	45 02	39 09	32	1	3.7
Krasnoufimsk	56 37	57 45	220	5	3.3
Krasnoyarsk	56 01	92 53	193	5	3.7
Krasnozerskoye	53 59	79 14	-	5	4.9
Krasnyy Chikoy	50 27	108 45	769	6	1.0
Kreshchenka	55 51	80 02	-	6	3.9
Kumara	51 34	126 43	178	5	1.7
Kunda	59 30	26 32	-	4	5.6
Kupino	54 22	77 18	120	5	5.8
Kuragino	53 53	92 40	-	-	1.7
Kureyka	66 29	87 09	-	-	3.3
Kurgan	55 28	65 24	78	5	4.3
Kursk	51 39	36 11	167	5	4.9
Kustanay	53 13	63 37	171	5	4.3
Kuybyshev	53 15	50 27	43	5	4.4
Kyakhta	50 22	106 27	789	-	1.2
Kuz'movka	62 19	92 02	-	-	1.3
Kuzomen'	66 17	36 54	-	-	6.1
Kyra	49 34	111 58	887	-	1.9

<u>Station</u>	<u>Coordinates</u>		<u>Elevation</u> (m)	<u>No. Months</u> <u>Temp. < 0°C</u>	<u>Ave. 5-mo.</u> <u>(Nov-Mar)</u> <u>wind speed</u> (m/s)
	<u>Lat(N)</u>	<u>Long(E)</u>			
Kyusyur	70° 41'	127° 24'	29	7	3.3
Kyzyl	51 42	94 27	640	7	0.9
L'vov	49 49	23 57	324	3	3.8
Leninakan	40 47	43 50	1528	4	1.2
Leningrad	59 56	30 16	6	5	3.2
Leushi	59 37	65 47	82	6	3.8
Lodeynoye Pole	60 44	33 33	-	-	4.5
Lokshak	54 44	130 77	84	7	1.0
Lomonosovo	75 37	91 20	-	-	7.1
Losinoborskaya	58 27	89 28	-	-	2.5
Loukhi	66 04	33 04	-	6	3.2
Lovozero	68 00	35 04	-	-	3.3
Lubny	50 01	33 02	164	-	3.9
Lukoyanov	55 02	44 30	206	5	4.0
Magadan	59 35	150 47	118	7	5.7
Magnitogorsk	53 21	59 05	381	5	4.8
Maksimkin Yar	58 41	86 49	104	7	3.1
Malinovo	45 25	134 16	-	5	2.7
Mama	58 18	112 54	-	7	2.8
Mamakan	57 48	114 01	-	7	1.7
Markovo	64 41	170 25	32	7	3.1
Marresale	69 43	66 49	11	7	6.8
Maslyanino	54 20	84 13	-	-	3.2
Maysk	57 47	77 17	114	6	3.3
Melitopol'	46 50	35 22	39	2	5.6
Mezen'	65 52	44 13	14	6	4.8
Minsk	53 52	27 32	234	4	4.6
Mogocha	53 44	119 47	619	7	1.3
Monchegorsk	67 56	32 58	-	-	4.4
Mondy	51 40	100 59	1304	7	2.5
Moscow	55 45	37 34	156	5	4.0
Motygino	58 11	94 40	-	-	2.3
Mozdok	43 44	44 40	132	3	2.3
Murmansk	68 58	33 03	46	6	6.1
Mutoray	61 20	100 30	-	-	2.0
Mys Chelyuskin	77 43	104 17	13	7	6.9
Mys Kamenny	68 28	73 36	14	7	6.9
Mys Kronotskiy	54 45	162 08	-	5	4.4
Mys Leskina	72 17	79 42	11	7	7.2
Mys Navarin	62 16	179 08	10	7	9.7
Mys Osten-Saken	76 14	98 50	-	7	6.2
Mys Ozernoy	57 43	163 19	-	6	5.3
Mys Shmidta	68 55	179 29(W)	7	7	5.7
Mys Sosunova	46 31	138 16	-	5	6.9
Mys Sterlegova	75 25	88 54	7	7	7.3
Mys Syurkum	50 06	140 41	-	5	6.2

Station	Coordinates		Elevation (m)	No. months Temp. < 0°C	Ave. 5-mo. (Nov-Mar) wind speed (m/s)
	Lat(N)	Long(E)			
Mys Tigrovyy	43° 56'	58° 44'	-	2	4.6
Mys Uelen	66 10	169 50(W)	7	7	6.6
Mys Yelizavety	54 25	142 43	77	6	6.3
Nakanno	62 52	108 26	-	7	1.3
Nakhichevan	39 12	45 25	875	3	1.4
Nal'chik	43 30	43 37	441	2	1.9
Napas	59 56	81 59	89	7	2.7
Nar'yan Mar	67 39	53 01	7	7	5.3
Narva	59 23	28 12	-	5	4.6
Naryn	41 26	76 00	2048	6	1.5
Nayakhan	61 55	159 00	22	7	7.9
Nazimovo	59 30	90 58	-	-	3.0
Nelyaty	56 30	115 40	474	7	1.5
Nerchinskiy Zavod	51 19	119 37	626	-	0.9
Nevel'sk	46 40	141 52	6	4	6.0
Nevinnomyssk	44 38	41 38	333	3	4.5
Nevon	58 02	102 44	231	7	1.6
Nikol'skoye	55 12	165 59	6	5	9.3
Nikolayev	46 58	32 00	19	3	3.8
Nikolayevsk na Amure	53 09	140 42	42	6	3.4
Nizhneangarsk	55 47	109 33	-	7	2.1
Nizhneilimsk	57 11	103 16	-	7	1.7
Nizhnekolymsk	68 32	160 56	8	7	2.4
Nizhneudinsk	54 53	99 02	410	5	1.8
Noginskiy	64 32	91 10	-	-	2.3
Nogliki	51 50	143 08	10	6	4.9
Norsk	52 21	129 55	207	5	1.2
Noril'sk	69 20	88 06	-	6	6.1
Novgorod	58 31	31 17	-	-	5.4
Novosibirsk	55 02	82 54	162	5	4.8
Novo-Yerudinskiy	59 45	93 32	-	-	1.8
Novyy Port	67 40	72 52	5	7	7.3
Odessa	46 29	30 38	64	2	5.4
Odinnadtsaty	55 54	119 36	1079	7	1.9
Okhotsk	59 22	143 12	6	7	5.1
Okhotskiy Perevoz	61 53	135 33	-	7	0.7
Oktyabr'skoye	62 27	66 03	38	7	2.8
Ol'khon	53 03	106 54	468	6	3.3
Oleminsk	60 24	120 55	125	7	2.2
Olenek	68 30	112 26	127	7	1.3
Olonets	60 59	32 58	15	5	4.4
Opochka	56 42	28 43	-	-	3.7
Oymyakon	63 28	142 49	-	6	0.4
Onguday	50 45	86 09	-	5	0.3
Onor	50 14	142 35	180	6	2.4
Ordzhonikidze	43 03	44 39	670	2	1.2

<u>Station</u>	<u>Coordinates</u>		<u>Elevation</u> (m)	<u>No. months</u> <u>Temp. < 0°C</u>	<u>Ave. 5-mo.</u> <u>(Nov-Mar)</u> <u>wind speed</u> (m/s)
	<u>Lat(N)</u>	<u>Long(E)</u>			
Orel	52° 55'	36° 05'	-	5	6.0
Orenburg	51 45	55 06	109	5	4.3
Orlik	52 30	99 49	1407	7	1.5
Ostrov Andreya	76 53	111 25	-	-	5.6
Ostrov Domashniy	79 30	91 08	3	7	5.8
Ostrova Izvestiy Tsik	75 55	82 30	-	-	6.2
Ostrova Krasoflotskiye	79 40	98 45	-	-	5.5
Ostrov Malyy Taymyr	78 08	107 12	-	-	5.6
Ostrov Morzhovets	66 44	42 30	-	6	7.7
Ostrov Prebrazheniya	74 40	112 56	6	7	5.1
Ostrov Russkiy	77 10	96 25	10	7	6.4
Ostrov Uyedineniya	77 30	82 12	10	7	6.6
Ostrov Vize	79 30	76 59	18	7	6.5
Oymyakon	63 28	142 49	-	7	0.4
Ozernovskiy	51 29	156 29	6	6	8.9
Ozero Taymyr	74 30	102 30	-	7	5.5
Padany	63 17	33 25	-	6	3.9
Palana	59 06	159 59	-	7	3.9
Panfilov	44 10	80 04	640	3	1.9
Parnu	58 24	24 32	-	4	5.3
Pavlodar	52 17	76 57	146	5	5.3
Penza	53 08	45 01	174	5	5.4
Perm	58 01	56 18	160	5	5.0
Petropavlovsk	54 50	69 09	135	5	5.2
Petropavlovsk Kamchatskiy	52 58	158 45	7	5	6.6
Petrozavodsk	61 49	34 20	-	5	4.8
Pikhtovka	56 00	82 42	-	6	3.3
Pil'vo	50 03	142 10	54	5	5.2
Pinsk	52 07	26 08	143	4	4.7
Pkhusun	43 22	134 48	7	5	4.0
Podkamennaya Tunguska	61 36	90 09	60	7	3.1
Pogranichnyy	44 24	131 23	-	5	3.8
Poligus	62 00	94 41	-	-	0.6
Poliny Osipenko	52 25	136 28	-	-	1.4
Poronaysk	49 13	143 06	4	5	3.2
Potapovo	68 40	86 20	-	-	4.6
Preobrazhenka	60 02	108 05	-	7	2.3
Prikumsk	44 48	44 10	118	2	3.7
Priozersk	61 02	30 07	-	-	3.1
Proliv Yugorskiy Shar	69 49	60 45	13	7	8.9
Provideniya	64 26	173 14(W)	3	7	4.9
Pskov	57 50	28 21	42	5	4.0
Pudozh	61 48	36 32	-	5	3.6
Pytalovo	57 04	27 54	-	-	4.3
Rang-kul'	38 28	74 22	-	5	2.6
Reboly	63 49	30 47	-	6	3.5

<u>Station</u>	<u>Coordinates</u>		<u>Elevation</u> (m)	<u>No. months</u> <u>Temp. < 0°C</u>	<u>Ave. 5-mo.</u> <u>(Nov-Mar)</u> <u>wind speed</u> (m/s)
	<u>Lat(N)</u>	<u>Long(E)</u>			
Riga	56° 58'	24° 04'	3	4	5.4
Romny	44 45	134 28	216	5	1.7
Rostov na Donu	47 15	39 49	77	3	5.4
Rubtsovsk	51 30	81 13	214	5	5.8
Ryazan'	54 37	39 43	135	5	5.5
Salekhard	66 32	66 32	35	7	4.1
Sam-Pervyy	45 29	56 06	82	3	4.8
Saratov	51 34	46 02	156	5	5.3
Sarny	51 12	26 37	153	3	3.4
Saskylakh	71 58	114 05	24	7	3.4
Semipalatinsk	50 21	80 15	206	5	3.1
Serov	59 30	60 32	132	5	2.9
Severnoye	56 21	78 21	123	5	2.7
Severo-Yeniseyskiy	60 22	93 01	-	-	3.5
Seymchan	62 55	152 25	207	7	1.0
Shenkursk	62 06	42 54	50	5	4.1
Shepetkovo	66 33	159 25	125	7	1.2
Shepetovka	50 10	27 03	260	4	4.6
Shimanovsk	51 59	127 39	276	5	2.3
Shologontsy	66 15	114 17	235	7	0.8
Simferopol'	45 01	33 59	205	1	4.8
Skovorodino	54 00	123 58	399	7	2.0
Skuratovo	53 34	37 03	248	5	5.9
Slautnoye	63 11	167 52	304	7	4.6
Slavgorod	52 58	78 39	125	5	5.3
Smolensk	54 45	32 04	241	5	4.8
Soroki	48 09	28 18	-	-	4.9
Sortavala	61 42	30 41	-	5	3.9
Srednekolymsk	67 27	153 41	27	7	1.7
Stanovishche (Camp na Ser)	70 35	72 37	15	7	6.5
Stanovishche Tiuteyykha	71 23	67 33	45	7	7.6
Staraya Russa	58 00	31 23	-	-	4.5
Stravropol'	45 03	41 59	574	3	4.7
Strelka Chunya	61 45	102 48	-	-	2.0
Sukhanovka	51 21	139 06	20	5	4.1
Sukhaya Tunguska	65 10	87 55	-	-	3.5
Sukhobuzimskoye	56 30	93 16	-	5	3.4
Suntar	62 09	117 39	124	7	1.9
Sura	63 35	45 38	50	6	3.5
Surgut	61 15	73 30	43	6	4.7
Sutungu	46 45	136 01	-	5	2.6
Sverdlovsk	58 48	60 38	237	5	3.9
Svetlyy	56 26	115 55	-	7	1.6
Syktyvkar	61 40	50 51	96	-	4.3
Sym	60 20	88 23	-	-	2.3
Tallinn	59 25	24 48	43	4	5.6

<u>Station</u>	<u>Coordinates</u>		<u>Elevation</u> (m)	<u>No. months</u> <u>Temp. < 0°C</u>	<u>Ave. 5-mo.</u> <u>(Nov-Mar)</u> <u>wind speed</u> (m/s)
	<u>Lat(N)</u>	<u>Long(E)</u>			
Tambey	71° 30'	71° 50'	8	7	7.1
Tambov	52 43	41 27	-	5	4.3
Tanguy	55 23	100 58	-	7	1.5
Tara	56 54	74 23	74	5	4.0
Tarasily-yar	73 02	88 10	-	7	5.4
Tarko Sale	64 55	77 49	27	7	3.2
Tartu	58 23	26 43	-	5	4.0
Tashtyp	52 48	85 54	455	5	2.9
Taimba	60 18	98 58	168	7	1.1
Tayshet	55 57	98 00	-	5	3.4
Tazovskiy	67 28	78 44	4	7	6.4
Teli	51 02	90 14	-	-	0.5
Tembenchi	64 57	98 51	-	-	0.8
Temir	49 09	57 07	235	5	5.2
Terney	45 02	136 40	11	5	6.6
Tikhvin	59 39	33 31	-	-	4.2
Tiksi	71 35	128 55	7	7	5.8
Tisul'	55 45	88 19	189	5	4.7
Tobol'sk	58 09	68 11	43	5	4.1
Tokma	58 16	105 54	-	7	1.5
Tokubay Mogila	45 57	77 14	-	5	2.1
Tomsk	56 26	84 58	121	5	5.5
Toora-Khem	52 28	96 09	-	-	0.6
Trofimovsk	72 36	127 02	-	7	8.1
Troitsk	57 18	94 58	163	6	1.8
Troitsko-Pechorskoye	62 42	56 12	107	7	3.3
Troitskoye	49 27	136 34	29	5	4.0
Tselinograd	51 08	71 22	348	5	6.2
Tsyp-Navolok	69 43	33 05	8	6	9.8
Tula	54 12	37 27	165	5	5.3
Tulun	54 35	100 33	-	7	0.9
Tungokochen	53 34	115 34	812	7	1.1
Tuoy Khaya	62 32	111 14	239	7	2.3
Tura	64 17	100 15	139	7	1.6
Turan	52 08	93 55	-	-	1.2
Turkestan	43 16	68 13	209	2	2.6
Turochak	52 16	87 10	-	5	0.8
Turukhansk	65 49	87 59	32	7	3.9
Tymlat	59 31	163 07	304	7	4.4
Tyrma	50 04	132 08	313	5	1.0
Tyumen'	57 09	65 30	103	5	4.7
Ucharal	46 10	80 56	388	4	3.3
Ufa	54 45	56 00	196	5	4.6
Ulegorsk	49 05	142 04	53	5	5.5
Uman'	48 46	30 14	216	4	4.8
Umba	66 41	34 18	-	-	5.0

<u>Station</u>	<u>Coordinates</u>		<u>Elevation</u> (m)	<u>No. months</u> <u>Temp. < 0°C</u>	<u>Ave. 5-mo.</u> <u>(Nov-Mar)</u> <u>wind speed</u>
	<u>Lat(N)</u>	<u>Long(E)</u>			<u>(m/s)</u>
Ust Maya	60° 23'	134° 27'	175	7	1.1
Ust' Bol'sheretsk	52 40	156 14	7	6	6.6
Ust' Kamchatsk	56 13	162 28	6	6	5.8
Ust' Kamo	60 43	97 29	-	7	1.2
Ust' Kut	56 46	105 40	-	7	1.1
Ust' Port	69 39	84 24	27	7	5.9
Ust' Shchugor	64 16	57 37	75	-	3.4
Ust' Tsil'ma	65 27	52 10	27	7	4.7
Ust' Voyampolka	58 30	159 10	7	6	6.2
Ust' Yudoma	59 11	135 09	211	7	1.0
Uyaly	44 35	61 09	55	3	4.3
Uzhgorod	48 38	22 16	118	3	2.5
Uzhur	55 18	89 50	-	-	3.8
Vanavara	60 22	102 16	260	7	1.6
Vanzhil'kynak	60 22	84 06	-	7	2.1
Velikaya Kema	45 28	137 15	-	4	5.6
Velikiye	56 21	30 31	103	5	3.1
Vel'mo Pervoye	61 11	93 05	-	-	1.4
Vengerovo	55 41	76 45	-	5	4.0
Ventspils	57 22	21 33	4	2	6.6
Vereshchagino	64 14	87 37	-	-	3.3
Verkhne Imbatskoye	63 11	87 58	39	7	3.6
Verkhnyaya Mishikha	51 38	105 35	1280	7	3.0
Verkhoyansk	67 33	133 23	-	7	0.6
Verkhniy-Baskunchak	48 13	46 44	35	4	4.3
Vikulovo	56 49	70 37	70	5	3.3
Vilyuysk	63 46	121 37	107	7	1.9
Vitim	59 28	112 34	193	7	2.4
Vivi	63 54	97 50	-	-	1.6
Vladivostok	43 70	131 54	138	5	7.1
Volkhov	59 49	32 22	-	-	3.9
Volgograd	48 41	44 21	145	4	6.6
Volnovakha	47 36	37 30	-	-	5.2
Volochanka	71 00	94 28	-	7	3.5
Vologda	59 17	39 52	118	5	6.1
Vorenzha	63 53	35 15	-	6	3.3
Vorkuta	67 29	64 01	180	7	6.3
Vorogovo	61 02	89 35	64	7	3.8
Voronezh	51 47	39 10	164	5	5.2
Vostochnyy	48 17	142 38	6	5	4.2
Vozhega	60 28	40 12	178	5	4.1
Vyborg	60 42	28 45	-	-	3.8
Vyshniy	57 35	34 34	160	5	3.2
Yakutsk	62 05	129 45	103	7	2.0
Yaral'in	67 00	110 00	230	7	1.8
Yartsevo	60 14	90 12	57	7	3.8

<u>Station</u>	<u>Coordinates</u>		<u>Elevation</u> (m)	<u>No. months</u> <u>Temp. < 0°C</u>	<u>Ave. 5-mo.</u> <u>(Nov-Mar)</u> <u>wind speed</u> (m/s)
	<u>Lat(N)</u>	<u>Long(E)</u>			
Yashkul'	46° 11'	45° 21'	-	4	5.6
Yekaterino Nikol'sk	47 44	130 58	74	5	3.5
Yelabuga	55 46	52 04	89	5	3.5
Yelets	52 38	38 31	130	5	5.5
Yelizovo	53 10	158 24	-	6	3.0
Yemtsa	63 04	40 21	107	5	3.1
Yeniseysk	58 27	92 10	78	7	2.4
Yerbogachen	61 16	108 01	278	7	1.6
Yerevan	40 08	44 28	907	1	0.6
Yermakovskoye	53 16	92 24	301	-	2.0
Yerofey Pavlovich	53 58	121 56	-	7	1.2
Yesil'	51 53	66 20	221	5	4.2
Yessey	68 29	102 10	199	7	2.2
Yur'yevets	57 20	43 07	132	5	4.5
Yushkozero	64 45	32 07	-	6	3.0
Yuzhno-Sakhalinsk	46 55	142 44	-	5	4.5
Zaporozh'ye	47 48	35 12	86	3	5.6
Zavitinsk	50 07	129 28	242	5	2.8
Zaysan	47 28	84 55	602	5	1.8
Zemetchino	53 30	42 37	129	5	4.5
Zeya	53 54	127 14	232	6	2.1
Zherdevka	51 51	41 48	-	5	4.1
Zhigalovo	54 48	105 08	-	7	1.4
Zhigansk	66 46	123 24	57	7	3.2
Zhitkovichi	52 14	27 52	-	4	4.3
Zima	53 55	102 40	-	6	2.0
Zimnegorskyy	65 28	39 44	81	6	6.3
Zlatoust	55 10	59 41	457	5	4.6
Zmeinogorsk	51 10	82 13	387	5	3.9
Znamenka	51 30	95 36	-	-	0.8
Zyryanka	65 44	150 54	39	7	2.0