TEC-SR-7

Handbook for Transformation of Datums, Projections, Grids and Common Coordinate Systems

January 1996

U.S. Army Corps of Engineers
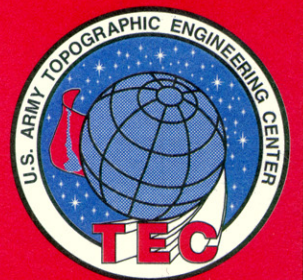Topographic Engineering Center
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13. ABSTRACT *(Maximum 200 words)*
This document provides Army orginizations and agencies with general guidance on selecting the appropriate methods for shifting between local geodetic datums and the World Geodetic System (WGS), and for converting Cartesian and map projection coordinates to and from geodetic coordinates. This guidance is provided to aid the Army community in selecting datum shift algorithms, in developing, selecting, and maintaining software using these algorithms, and in implementing this software to support operational units. Equations are furnished for map projections and datums commonly used within the Army, and references are provided for other, less commonly encountered, map projections and datums.

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EXECUTIVE SUMMARY

This document provides Army organizations and agencies with general guidance on selecting the appropriate methods for shifting between local geodetic datums and the World Geodetic System (WGS), and for converting Cartesian and map projection coordinates to and from geodetic coordinates. This guidance is provided to aid the Army community in selecting datum shift algorithms, in developing, selecting, and maintaining software using these algorithms, and in implementing this software to support operational units. Equations are furnished for map projections and datums commonly used within the Army, and references are provided for other, less commonly encountered, map projections and datums.

Beneficial comments (corrections, recommendations, additions, deletions) or information that may be of use in improving this document should be addressed to the Standards Division of the Digital Concepts & Analysis Center (DCAC) at the Topographic Engineering Center (TEC).

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PREFACE

This report was funded under the U.S. Army Topographic Engineering Center's (TEC) Digital Topographic Data Standards Program.

This report was co-authored at TEC by the following individuals:

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The report is based on a draft of the Military Handbook, Datums, Projections, Grids and Common Coordinate Systems, Transformation of (MIL-HDBK-600008), and was revised during the period July 1994 through July 1995. This Technical Report uses material from MIL-HDBK-600008 co-authored by:

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This work was performed under the previous supervision of Juan Perez, and Richard A Herrmann and the current supervision of John W. Hale, Chief, Standards Division and Regis J. Orsinger, Director, Digital Concepts and Analysis Center.

Mr. Walter E. Boge was Director and COL Richard G. Johnson was Commander and Deputy Director of TEC at the time of publication of this report.

1. SCOPE

1.1 <u>Scope</u>. This Technical Report provides methods and
parameter values to shift positions between approximately 113
geodetic datums and provides methods for converting between
geodetic coordinates, Cartesian coordinates and map projection
coordinates. Guidance is provided on selecting methods
appropriate to the application, and on developing and testing
software to implement these conversions.

1.2 <u>Applicability</u>. The methods provided in this Technical
Report are used for a wide range of Army mapping, charting, and
positioning applications. Equations are furnished for the map
projections commonly used within the Army. References are
provided for other map projections that occasionally may be
encountered. Special high accuracy applications such as
engineering, construction, and real estate boundary surveys are
outside the scope of this Technical Report. Appendix B provides
datum shift parameters for approximately 113 datums that are
currently available from the Defense Mapping Agency (DMA).
However, many more local datums exist. For guidance on datum
shift applications or for datum shift parameters not covered by
this Technical Report, contact TEC [address given in the preface].

1.3 <u>Application guidance</u>. Transformation methods and
parameters should be appropriate to the application. A system
analyst should do a thorough evaluation of the system requirements
before developing detailed specifications. This evaluation should
consider the intended use, data types, formats, and required
accuracy of the output mapping, charting, and geodesy (MC&G) data,
as well as the source and accuracy of the input data. Also, any
hardware constraints must be considered. The methods presented in
this Technical Report can then be examined for how well they can
meet or be adapted to meet the requirements. Where more than one
method will satisfy a basic requirement, give priority to Army-
preferred methods for standardization purposes. Additional
application guidance is furnished in Section 8.

2. REFERENCED DOCUMENTS

2.1 Government documents.

2.1.1 Specifications, standards, and handbooks. The following specifications, standards, and handbooks form a part of this document to the extent specified herein.

SPECIFICATIONS. None

STANDARDS.

MIL-STD-2401   Department of Defense World Geodetic
                        System(WGS) (1994)

STANAG 2211   Geodetic Datums, ellipsoids, grids and
(Fifth Edition)   grid references (1991)

HANDBOOKS. None

2.1.2 Other Government documents, drawings, and publications. The following other Government documents, drawings, and publications form a part of this document to the extent specified herein. Unless otherwise specified, the issues are those cited in the solicitation. The date in parentheses indicates for each document the edition that was used in preparation of this Technical Report.

Document ID/Agency   Title

DMA   Department of Defense Glossary of
                           Mapping, Charting, and Geodetic Terms,
                           Fourth Edition. (1981)

DMA TM 8358.1   Datums, Ellipsoids, Grids, and
                           Grid Reference Systems, (1990)

DMA TM 8358.2   The Universal Grids: Universal
                           Transverse Mercator (UTM) and
                           Universal Polar Stereographic (UPS).
                           (1989)

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DMA Instruction 8000.1 with change 1, 4 Aug 92	Geodetic and Geophysical Sign Conventions and Fundamental Constants. (1988)
Geological Survey Professional Paper 1395	Map Projections - A Working Manual. (Snyder, 1987)
NOAA Manual NOS NGS 5	State Plane Coordinate System of 1983. (Stem, 1989)
NSWC/DL TR-3624	Map Projection Equations. (Pearson, 1977)
U.S. Department of Commerce, Coast and Geodetic Survey, Special Publication No. 251	Conformal Projections in Geodesy and Cartography. (Thomas, 1979)
Engineer Technical Letter No. 1110-1-147	Engineering and Design Conversion to North American Datum of 1983 (1990)
Federal Register v. 55, no. 155, Friday, August 10, 1990 Docket No. 900655-0165	Notice To Adopt Standard Method for Mathematical Horizontal Datum Transformation (NOAA, 1990)

Copies of; DMA Instruction 8000.1 and GS Professional Paper 1395 are
available from USGS, Branch of Distribution, Box 25286, Denver CO
80225. Copies of DoD Glossary of Mapping, Charting, and Geodetic
Terms; DMA TM 8358.1; and DMA TM 8358.2 are available from the
Director, DMA Combat Support Center, ATTN: CCOR, 6001 MacArthur
Boulevard, Bethesda MD 20816-5001. Copies of the NOAA Manual NOS NGS
5 and the Coast and Geodetic Survey Special Publication No. 251 are
available from National Geodetic Information Branch, N/CG174, National
Geodetic Survey, 1315 East-West Highway, Silver Spring MD 20910.
Copies of NSWC/DL TR 3624 (Accession Document Number = AD A037 381 on
paper, or GIDEP #E151-2353 on cartridge) are available on from Defense
Technical Information Center, Bldg. 5, Cameron Station, Alexandria, VA
22304-6145.   Copies of ETL 1110-1-147 are available through the
Department of the Army, U.S. Army Corps of Engineers, Washington, D.C.
203140-1000. Copies of STANAG 2211 and MIL-STD 2401 are available
from the Defense Printing Service, ATTN: DODSSP, 700 Robbins Ave, Bld
4D Philadelphia PA 19111.

2.2 Non-Government publications.

Heiskanen, Weikko A., and Helmut Moritz; Physical Geodesy; Institute of Physical Geodesy, Technical University, Graz, Austria; 1967.

Krakiwsky, Edward J.; Conformal Map Projections in Geodesy; Department of Surveying Engineering, University of New Brunswick, Fredrickton, N.B., Canada; September, 1973.

Rapp, Richard H.; Geometric Geodesy - Part I; Department of Geodetic Science and Surveying, The Ohio State University, Columbus, Ohio; 1984.

Rapp, Richard H.; Geometric Geodesy - Part II; Department of Geodetic Science and Surveying, The Ohio State University, Columbus, Ohio; 1987.

2.3 Order of precedence. MIL-STD-2401 shall take precedence for all matters concerning WGS 84 and datum shift parameters. Nothing in this document, however, supersedes applicable laws and regulations unless a specific exemption has been obtained.

3. DEFINITIONS AND UNITS

3.1 <u>Acronyms</u>. The acronyms used in this Technical Report are defined as follows:

a.	BIH	Bureau International de l'Heure
b.	CONUS	continental United States
c.	CTP	Conventional Terrestrial Pole
d.	DMA	Defense Mapping Agency
e.	DoD	Department of Defense
f.	DTED	Digital Terrain Elevation Data
g.	ECEF	earth-centered, earth-fixed
h.	GPS	Global Positioning System
i.	LGS	local geodetic system
j.	MGRS	Military Grid Reference System
k.	MC&G	mapping, charting and geodesy
l.	MRE	multiple regression equation
m.	MSL	mean sea level
n.	NAD 27	North American Datum 1927
o.	NAD 83	North American Datum 1983
p.	NGS	National Geodetic Survey
q.	NGVD 29	National Geodetic Vertical Datum of 1929
r.	NOAA	National Oceanic and Atmospheric Administration
s.	NOS	National Ocean Survey
t.	NSWC/DL	Naval Surface Weapons Center/Dahlgren Laboratory
u.	TM	Transverse Mercator
v.	UPS	Universal Polar Stereographic
w.	UTM	Universal Transverse Mercator
x.	WGS 72	World Geodetic System 1972
y.	WGS 84	World Geodetic System 1984

3.2 <u>Terms</u>.

3.2.1 <u>Convergence of the meridian</u> (γ). The angle between the projection of the meridian at a point on a map and the projection North axis (typically represented by the y axis of the projection). The convergence of the meridian is positive in the clockwise direction.

3.2.2 <u>Coordinate</u>. Linear or angular quantities that designate the position that a point occupies in a given reference frame or system. Also used as a general term to designate the particular kind of reference frame or system, such as Cartesian coordinates or spherical coordinates.

3.2.3 <u>Datum</u>. Any numerical or geometrical quantity or set of such quantities specifying the reference coordinate system used for geodetic control in the calculation of coordinates of points on the earth. Datums may be either global or local in extent. A local datum defines a coordinate system that is used only over a

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region of limited extent. A global datum specifies the center of the reference ellipsoid to be located at the earth's center of mass and defines a coordinate system used for the entire earth.

3.2.3.1 Horizontal datum. A horizontal datum specifies the coordinate system in which latitude and longitude of points are located. The latitude and longitude of an initial point, the azimuth of a line from that point, and the semi-major axis and flattening of the ellipsoid that approximates the surface of the earth in the region of interest define a horizontal datum.

3.2.3.2 Vertical datum. A vertical datum is the surface to which elevations are referred. A local vertical datum is a continuous surface, usually mean sea level, at which elevations are assumed to be zero throughout the area of interest.

3.2.4 Earth-fixed. Stationary with respect to the earth.

3.2.5 Elevation. Vertical distance measured along the local plumb line from a vertical datum, usually mean sea level or the geoid, to a point on the earth. Often used synonymously with orthometric height.

3.2.6 Ellipsoid. The surface generated by an ellipse rotating about one of its axes. Also called ellipsoid of revolution.

3.2.7 Equator. The line of zero geodetic latitude; the great circle described by the semi-major axis of the reference ellipsoid as it is rotated about the semi-minor axis.

3.2.8 Equipotential surface. A surface with the same potential, usually gravitational potential, at every point; a level surface.

3.2.9 Geocentric Cartesian coordinates. Cartesian coordinates (X, Y, Z) that define the position of a point with respect to the center of mass of the earth.

3.2.10 Geodetic coordinates (geodetic position). The quantities of latitude, longitude, and geodetic height (ϕ, λ ,h) that define the position of a point on the surface of the earth with respect to the reference ellipsoid. Also, imprecisely, called geographic coordinates (See also geographic coordinates).

3.2.11 Geodetic height (ellipsoidal height, h). The height above the reference ellipsoid, measured along the ellipsoidal normal through the point in question. The geodetic height is positive if the point is outside the ellipsoid.

3.2.12 <u>Geodetic latitude (ϕ)</u>. The angle between the plane of the Equator and the normal to the ellipsoid through the computation point. Geodetic latitude is positive north of the equator and negative south of the Equator (See figure 4.2).

3.2.13 <u>Geodetic longitude (λ)</u>. The angle between the plane of a meridian and the plane of the prime meridian. A longitude can be measured from the angle formed between the local and prime meridians at the pole of rotation of the reference ellipsoid, or by the arc along the Equator intercepted by these meridians. See figure 4.2.

3.2.14 <u>Geographic coordinates</u>. The quantities of latitude and longitude that define the position of a point on the surface of the earth. See also geodetic coordinates.

3.2.15 <u>Geoid</u>. The equipotential surface of the earth's gravity field approximated by undisturbed mean sea level of the oceans. The direction of gravity passing through a given point on the geoid is perpendicular to this equipotential surface.

3.2.16 <u>Geoid separation (N)</u>. The distance between the geoid and the mathematical reference ellipsoid as measured along the ellipsoidal normal. This distance is positive outside, or negative inside, the reference ellipsoid. Also called geoidal height; undulation of the geoid. See figure 4.1.

3.2.17 <u>Grid reference system</u>. A plane-rectangular coordinate system usually based on, and mathematically adjusted to, a map projection in order that geodetic positions (latitudes and longitudes) may be readily transformed into plane coordinates and the computations relating to them may be made by the ordinary methods of plane surveying (see also 3.2.21).

3.2.18 <u>Map projection</u>. A function relating coordinates of points on a curved surface (usually an ellipsoid or sphere) to coordinates of points on a plane. A map projection may be established by analytical computation or, less commonly, may be constructed geometrically.

3.2.19 <u>Map scale</u>. The ratio between a distance on a map and the corresponding actual distance on the earth's surface.

3.2.20 <u>Meridian</u>. A north-south reference line, particularly a great circle through the geographical poles of the earth, from which longitudes and azimuths are determined; or the intersection of a plane forming a great circle that contains both geographic poles of the earth, and the ellipsoid.

3.2.21 <u>Military Grid Reference System (MGRS)</u>. MGRS is a system developed by DMA for usage with UTM and UPS projections.

3.2.22 <u>Orthometric height</u>. See elevation.

3.2.23 <u>Parallel</u>. A line on the earth, or a representation thereof, that represents the same latitude at every point.

3.2.24 <u>Prime (initial) meridian</u>. A meridian from which the longitudes of all other meridians are reckoned. This meridian, of longitude 0°, was traditionally chosen to pass through the Greenwich Observatory in Greenwich, England. For new refined coordinate systems, the location of the prime meridian is defined by the International earth Rotation Service, Paris, France.

3.2.25 <u>Reference ellipsoid</u>. An ellipsoid whose dimensions closely approach the dimensions of the geoid; the exact dimensions are determined by various considerations of the section of the earth's surface concerned. Usually a bi-axial ellipsoid of revolution.

3.2.26 <u>Scale factor (projection)</u>. A multiplier for reducing a distance in a map projection to the actual distance on the chosen reference ellipsoid.

3.3 <u>Units</u>. In this Technical Report all distances are expressed in meters. Angles are expressed in degrees, radians or the combination of degrees, minutes, and seconds.

3.4 <u>Sign conventions</u>. The sign conventions used in this Technical Report are defined as follows:

a. <u>Elevation</u>. Positive outside or above the vertical datum, negative inside or below the datum.

b. <u>Geodetic height</u>. Positive outside the reference ellipsoid, negative inside the reference ellipsoid.

c. <u>Geodetic latitude</u>. Positive in the northern hemisphere, negative in the southern hemisphere.

d. <u>Geodetic longitude</u>. Zero at the prime meridian, positive 0° to 180°, eastward from the prime meridian, and negative 0° to 180°, westward from the prime meridian.

e. <u>Geoid separation</u>. Positive outside the reference ellipsoid, negative inside the reference ellipsoid.

f. <u>Cartesian Z</u>. Positive in the northern hemisphere.

3.5 <u>Unit conversion factors</u>. The standard unit of distance measurement is the meter. However, MC&G products in many parts of the world were produced using other units of measurement. The user must ascertain the units used in non-standard products and appropriately scale them for use in the equations given in this

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technical report. Some commonly used conversion factors are listed in Table 3.1.

TABLE 3.1 <u>Conversion Factors</u>.

1 meter	= 3.28083333333 U.S. Survey feet
	= 3.28083989501 International feet
1 International foot	= 0.3048 meter (Exact)
1 U.S. Survey foot	= 1200/3937 meter (Exact)
	= 0.30480060960 meter
1 Int'l Nautical mile	= 1852 meters (Exact)
	= 6076.10333333 U.S. Survey feet
	= 6076.11548556 International feet
1 Int'l Statute mile	= 1609.344 meters (Exact)
	= 5280 International feet (Exact)
1 degree	= 60 minutes
1 degree	= 3600 seconds
1 minute	= 60 seconds
6400 U.S. mils	= 360 degrees

3.5.1 <u>Degrees and radians</u>. A circle contains 360 degrees or 2π radians. If an angle contains d degrees or, equivalently, r radians, then d and r are related by

$$r = \frac{\pi d}{180}$$
(3.1)

3.5.2 <u>Specifying the unit of angular measure</u>. In this Technical Report the unit of angular measure will be specified if it matters. A general rule of thumb is that angles must be in radians in an equation that was derived using calculus and trigonometry. Radians are the usual units in computer evaluation of trigonometric functions.

4. FUNDAMENTAL CONCEPTS

4.1 <u>Introduction</u>. An understanding of geodesy helps one understand how to represent the locations of points on or about the surface of the earth. Terms printed in boldface are defined in Section 3.2.

a. The location of a point is described by a set of **coordinates**. Some types of coordinates, such as Cartesian (see Section 4.3.1), uniquely identify three-dimensional locations of points in space. Their use is computationally efficient in many applications. However, users often have difficulty visualizing the relationship of Cartesian coordinates to the surface of the earth. **Geodetic coordinates** (see Section 4.3.3) make visualization easier by providing horizontal coordinates on the surface of a figure approximating the shape of the earth and a vertical (height) coordinate above or below the reference surface.

b. Geodetic coordinates are defined by a geodetic reference system. A geodetic reference system includes, but is not limited to, a reference surface (usually an ellipsoid) and a set of parameters positioning and orienting the reference surface with respect to the earth. Over the years, many geodetic reference systems have been developed for various localities. These systems are often called local datums and are not rigorously tied to each other.

c. When applications require that two or more points be located in a common reference system, transformation between two coordinate systems may be needed. For example: both weapon and target must be located in the same reference system if the indirect fire computations are to be valid. As the ranges of weapons and target acquisition systems increased and positional data were derived from multiple sources, commanders often were confronted with the problem of using positions on more than one reference system. With the advent of satellite geodesy one can establish global geodetic reference systems and take measurements that allow local datums to be related to the global system. The standard global reference system for DoD is the World Geodetic System 1984 (WGS 84). The process of relating one reference system to another is called a datum shift, or datum transformation.

d. Coordinates generally are represented alpha-
    numerically (e.g. latitude 38° 51' 22.45"N and
    longitude 77° 26' 34.56"W). Position data for an area
    often are represented graphically, as on a map. The
    planar representation on a map provides a distorted
    view of features on the curved earth's surface.
    Geodesists have developed various mathematical methods
    to minimize certain types of map distortions, depending
    on the intended use of the map. The process of
    relating coordinates of points on a curved surface
    (usually an ellipsoid or sphere) to coordinates of
    points on a plane is called **map projection**. The
    preferred map projections for DoD are the Universal
    Transverse Mercator (UTM), used over most of the
    landmass, and Universal Polar Stereographic (UPS), used
    in the polar regions.

    4.2   Reference surfaces. Reference surfaces are used as the
bases for geodetic coordinate systems and elevations.

    4.2.1   Reference ellipsoid. A reference surface called the
reference ellipsoid (mathematically described in Section 4.3.2)
is used for defining geodetic coordinates (Section 4.3.3) because
it is regularly shaped, is mathematically tractable, and can
approximate the surface of the earth. Since the earth is
flattened at the poles, an oblate ellipsoid fits the shape of the
earth better than does a sphere. An ellipsoid of revolution is
the reference surface most commonly used in geodesy and is used
throughout this Technical Report. A geodetic coordinate has a
simply defined relationship with the reference ellipsoid and any
other geodetic coordinate because of the regularity and relative
simplicity of the reference ellipsoid's shape. The parameters for
the commonly used reference ellipsoids are listed in Appendix A.
A reference ellipsoid is often chosen to fit the **geoid**.

    4.2.2   Geoid. The **geoid** is a closed surface of constant
gravity potential approximated by mean sea level (MSL) and the
theoretical extension of MSL through land areas. For purposes of
this technical report, the geoid is considered a reference surface
from which elevations are measured. A reference ellipsoid may fit
the geoid; for example, the maximum difference between the WGS 84
geoid model and the WGS 84 ellipsoid found in Appendix D, Table
D.1 is 102 meters.

    4.2.3   Relationships among topography, the geoid, and the
reference ellipsoid. The relationships among topography, the
geoid, and the reference ellipsoid are illustrated in figure 4.1.
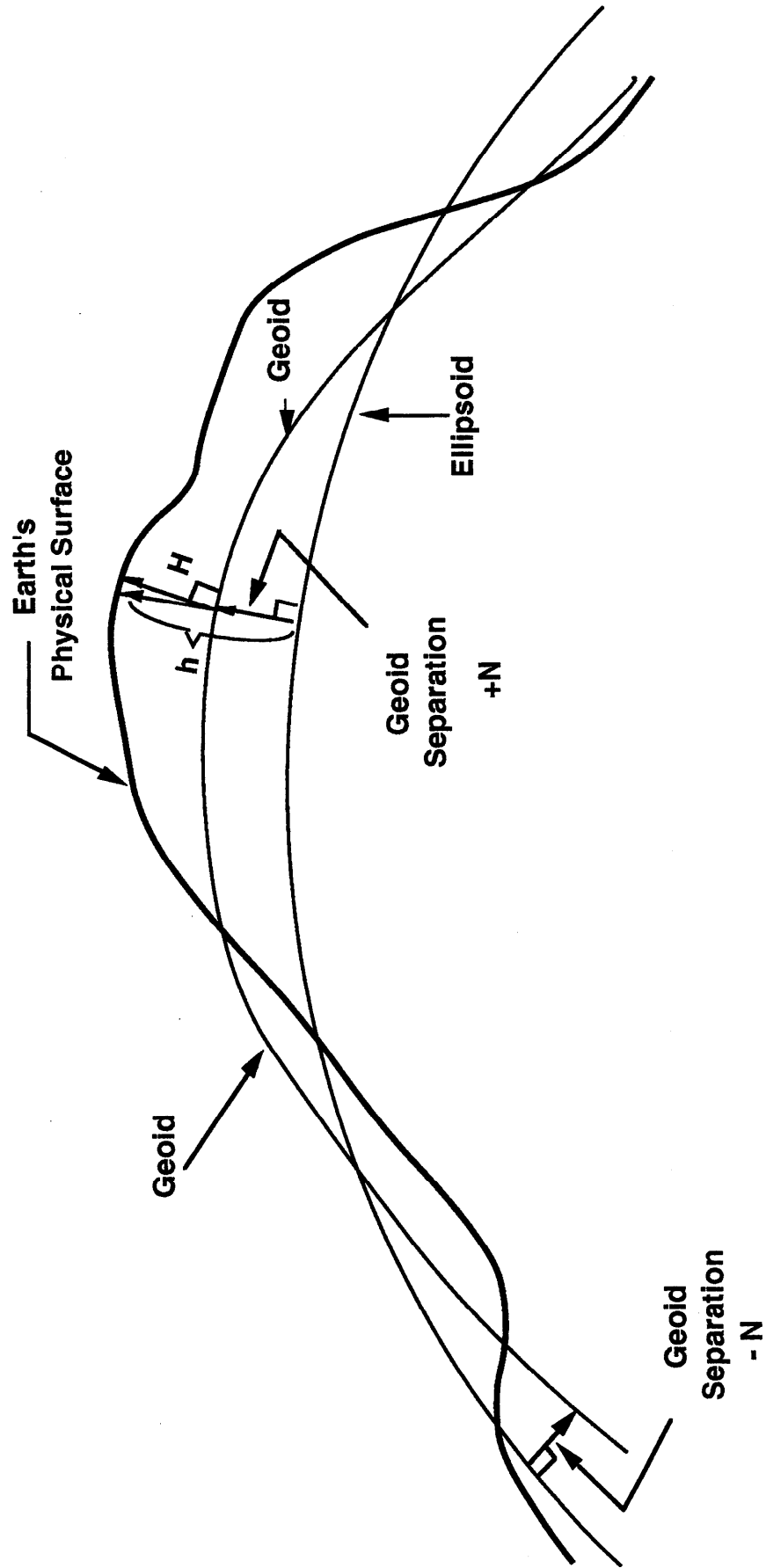
Figure 4.1 The relationship between the reference ellipsoid, the geoid, and the physical surface of the earth. Note relationship of elevation (H), geodetic height (h), and geoid separation (N).

4.3 Earth-fixed coordinate systems. There are many
coordinate systems used to represent positions on or near the
earth. The coordinate systems used in this Technical Report are
all **earth-fixed**. A coordinate system is earth-fixed if the axes
are stationary with respect to the rotating earth. Two primary,
earth-fixed coordinate systems, Cartesian and geodetic, are
discussed in Sections 4.3.1 and 4.3.3, respectively. Conversions
from one to the other are presented in Section 4.3.4.

4.3.1 Cartesian coordinates (X, Y, Z). The Cartesian
coordinate system used in this Technical Report is a right-hand,
rectangular, three-dimensional (X, Y, Z) coordinate system with
an origin at (0,0,0). When the origin is also at the mass center
of the earth, the coordinate system is said to be geocentric.

4.3.2 Cartesian coordinate system/reference ellipsoid
relationship. The relationship of the Cartesian coordinate system
to the reference ellipsoid is shown in figure 4.2. The Z-axis,
the axis of rotation (semi-minor axis) of the ellipsoid, is nearly
parallel to the axis of rotation of the earth. The Z-coordinate
is positive toward the North pole. The X-Y plane lies in the
equatorial plane, the plane swept out by the semi-major axis as
the ellipse is rotated. The X-axis lies along the intersection of
the plane containing the prime (initial) meridian and the
equatorial plane. The X-coordinate is positive toward the
intersection of the prime meridian and equator. The reference
ellipsoid satisfies the equation

$$\frac{X^2}{a^2} + \frac{Y^2}{a^2} + \frac{Z^2}{b^2} = 1 \qquad\qquad (4.1)$$

Here, a and b, respectively, denote the semi-major and semi-minor
axes of the ellipsoid. This surface can be generated by rotating
an ellipse defined by

$$\frac{X^2}{a^2} + \frac{Z^2}{b^2} = 1 \qquad\qquad (4.2)$$

about the Z-axis. Ideally, for global applications, the Z-axis
coincides with the mean axis of rotation of the earth and the
origin is the center of mass of the earth. Often, in practice, an
ellipsoidal surface is fitted to a local area so that the origin
and axes vary slightly from ideal coincidence.

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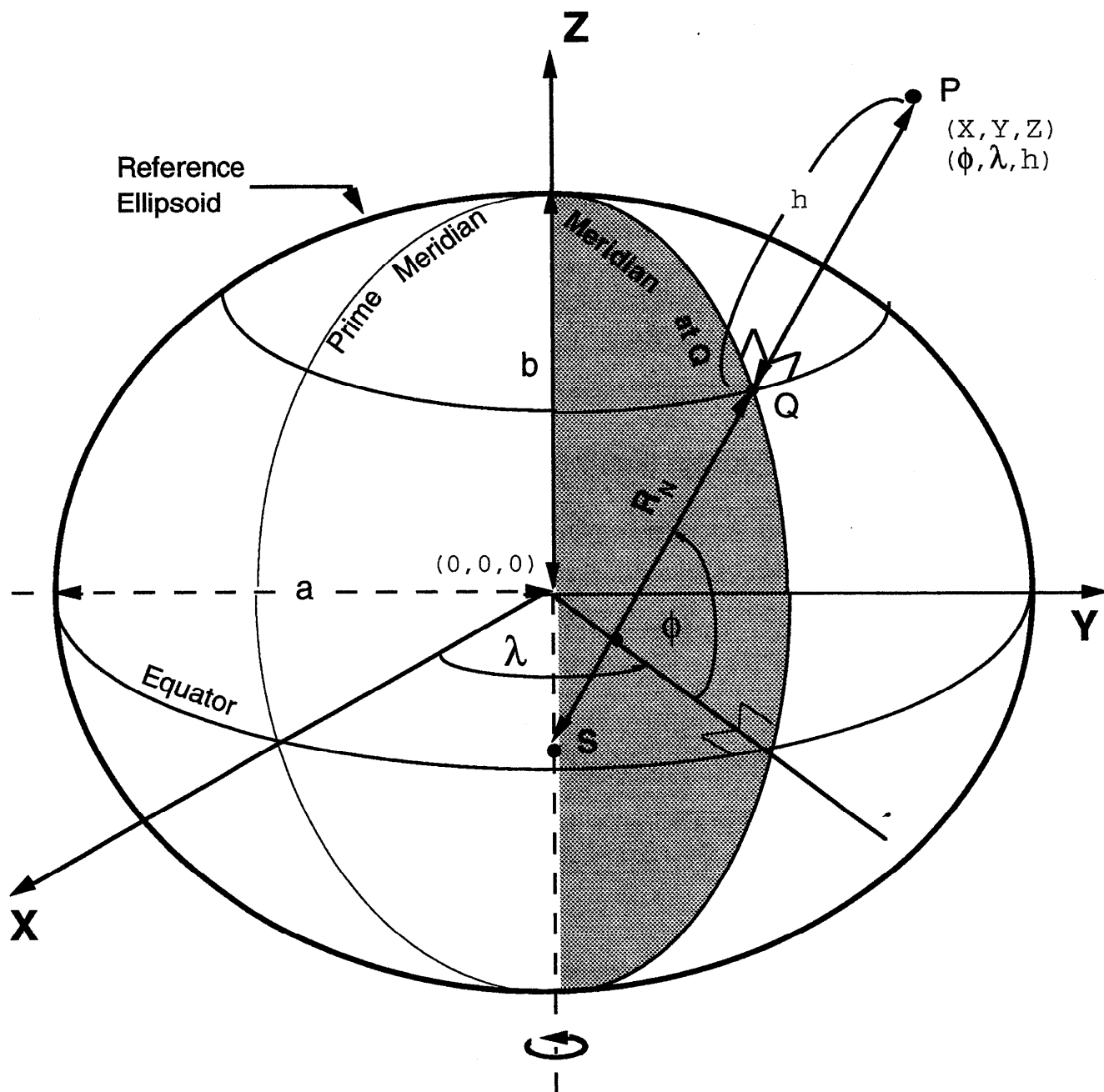
Figure 4.2 The geometric relationship between Cartesian and geodetic coordinates.

4.3.2.1 <u>First eccentricity and flattening</u>. Two important constants are the (first) eccentricity ε and flattening f that are related to a and b by

$$\varepsilon^2 = \frac{a^2 - b^2}{a^2} \qquad (4.3)$$

$$f = \frac{a - b}{a}$$

Frequently a and f are given, rather than a and b, as the defining parameters of a reference ellipsoid

4.3.2.2 <u>Ellipsoid parameters</u>. Appendix A contains a list of reference ellipsoids and their parameters. These can be divided into two categories: those, such as the Bessel Ellipsoid, that are used to approximate the shape of the earth for a local area and those, such as the WGS 84 Ellipsoid, that are used to yield a global approximation to the shape of the earth.

4.3.3 <u>Geodetic coordinates</u> (ϕ, λ, h). Geodetic coordinates consist of geodetic latitude (ϕ), geodetic longitude (λ), and geodetic height (h). Their relationship to the reference ellipsoid is shown in figure 4.2. For an ellipsoid that satisfies equation 4.1, the normal SP intersects the ellipsoid at Q. The angle between the normal SP and the equatorial (X-Y) plane is called the **geodetic latitude** (ϕ) of point P. The meridian plane containing point P is defined as the half-plane containing the Z-axis and point P. The angle between the prime meridian (X-Z) plane and the meridian containing point P is the **geodetic longitude** (λ) of point P. Geodetic longitude is not defined when P lies on the Z-axis. The distance from Q to P is called the **geodetic height** (h).

4.3.3.1 <u>Latitude and longitude limits</u>.

Latitude limits:
-90° to +90°
-90° at the South Pole
0° at the Equator
+90° at the North Pole

Longitude limits:
0° at the Prime meridian
180°W (-180°) to 180°E (+180°)

4.3.4 <u>Coordinate conversion</u>. Two basic earth-fixed coordinate systems, geocentric Cartesian coordinates and geodetic coordinates, have been introduced in Sections 4.3.1 and 4.3.3. The following section explains how coordinates in one system can

be transformed to the other. In Section 6, transformations between coordinates on two different reference ellipsoids are discussed. Map projections are treated in Section 7 and the general conversion process is summarized in Section 8.1.

4.3.4.1 Geodetic to Cartesian coordinate conversion. If geodetic coordinates (ϕ, λ, h) are known, then the Cartesian coordinates are given by

$$X = (R_N + h) \cos \phi \cos \lambda \qquad (4.4)$$
$$Y = (R_N + h) \cos \phi \sin \lambda$$
$$Z = \left(\frac{b^2}{a^2} R_N + h\right) \sin \phi$$

where R_N, the radius of curvature in the prime vertical, is given by

$$R_N = \frac{a^2}{\sqrt{a^2 \cos^2 \phi + b^2 \sin^2 \phi}} \qquad (4.5)$$

$$= \frac{a}{\sqrt{1 - \varepsilon^2 \sin^2 \phi}}$$

See Figure 4.2. Equation sets 4.4 and 4.5 are exact. Calculating h from elevation (H) is discussed on section 6.5.

4.3.4.2   Cartesian to geodetic coordinate conversion.

4.3.4.2.1 Finding λ. In the reverse case, when (X, Y, Z) are known but (ϕ, λ, h) are unknown,

for	the longitude (λ) is given by
X $\neq$ 0, all Y,	λ = arctan $\frac{Y}{X}$ (4.6)
X = 0,	Y > 0, λ = 90° E (+90°)
X = 0,	Y < 0, λ = 90° W (-90° or 270° E)
X = 0,	Y = 0, λ is undefined

Equation set 4.6 is exact. Take care to ensure that the evaluation of the arctan function yields a value of λ lying in the desired quadrant.

4.3.4.2.2 Finding ϕ. When (X, Y) is not (0, 0), the general equations for ϕ and h are more involved. The method presented for finding ϕ is an iterative procedure of Bowring. Following Rapp (1984, pp. 123-124), one finds an initial approximation to a variable β given by

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$$\tan \beta_0 = \frac{a \ Z}{b\sqrt{X^2 + Y^2}} \tag{4.7}$$

Once β_0 is obtained, substitute it for β in the equation

$$\tan \phi = \frac{Z + \varepsilon'^2 \ b \ \sin^3 \beta}{\sqrt{X^2 + Y^2} - a \ \varepsilon^2 \ \cos^3 \beta} \tag{4.8}$$

Here ε^2 and ε'^2 are given by

$$\varepsilon^2 = \frac{a^2 - b^2}{a^2} = 2f - f^2 \tag{4.9}$$

$$\varepsilon'^2 = \frac{a^2 - b^2}{b^2} = \frac{\varepsilon^2}{1 - \varepsilon^2}$$

This approximation to ϕ is substituted into

$$\tan \beta = (1 - f) \tan \phi \tag{4.10}$$

to give an updated approximation to β. This procedure, of using the latest approximation of β (equation 4.10) to produce an updated approximation of ϕ (equation 4.8), and then using this new value of ϕ to update β, can be continued until the updated value of ϕ is sufficiently close to the previous value. For many purposes, one iteration is sufficient when using the Bowring method, so that ϕ can be found directly from equation 4.8, using the original value of β_0. Rapp (1984, p. 124) states that for terrestrial applications, one such iteration yields values accurate to within 0.1 millimeter. In the polar case (X = 0, Y = 0, Z), the latitude (ϕ) is defined to be +90° (90°N) for positive Z values and -90° (90°S) for negative Z values.

   4.3.4.2.3 <u>Calculating h</u>. Once ϕ has been calculated, in non-polar areas h can be found from

$$h = \frac{\sqrt{X^2 + Y^2}}{\cos \phi} - R_N \tag{4.11}$$

In polar regions it is preferable to use

$$h = \frac{Z}{\sin \phi} - R_N + \varepsilon^2 \ R_N \tag{4.12}$$

where R_N is defined in equation 4.5.

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Note: For many applications h is not required and need not be calculated. An exception to this statement is the computation of height or elevation for the Global Positioning System (GPS).

4.4 <u>Representation of Geodetic Coordinates</u> Latitude is given before longitude. Latitude and longitude usually are numerically represented as: decimal degrees (e.g. 44.44° N); integer degrees and decimal minutes (e.g.44° 44.44'S); or integer degrees, integer minutes and decimal seconds (e.g. 44° 44' 44.44" E). For clarity, numeric fields representing degrees, minutes and seconds should be ended by the appropriate symbol, respectively " ° ", " ' ", and " " ". Numeric fields with a value of zero should be indicated by a "00", and not left blank. The hemisphere of both latitude and longitude must be shown, preferably by placing an "N" or "S" at the end of numeric latitude and an "E" or "W" after numeric longitude.

4.5 <u>Height Relationships</u>.

4.5.1 <u>Elevations</u>. Throughout this Technical Report, the term **elevation** (H) is used to denote the distance of a point above the geoid or vertical reference surface as measured along the plumb line. A plumb line follows the direction of gravity and is perpendicular to all equipotential surfaces of the Earth's gravity field that intersect it. Points lying outside the geoid are defined as having positive elevation. For the purposes of this Techncial Report, elevation and **orthometric height** are considered equivalent. In practice, the reference surface for measuring elevations may not exactly coincide with the geoid.

4.5.2 <u>Geoid separation</u>. **Geoid separation** (N) is the distance from a reference ellipsoid to the geoid, measured along the ellipsoidal normal. This is equivalent to the geodetic height (see Section 4.3.3) of a point on the geoid. Geoid separation is positive when the geoid lies outside the ellipsoid.

4.5.3 <u>The relationships among H, h, and N</u>. Geodetic height (h), geoid separation (N), and elevation (H) are depicted in figure 4.1 and are related by equation 4.13.

$$h \approx H + N \qquad\qquad (4.13)$$

4.5.3.1 <u>Notational use of H and h</u>. This Techncial Report uses the classical notation of H for elevation and h for geodetic height.

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5. GEODETIC SYSTEMS AND DATUMS

5.1 <u>Introduction</u>. A geodetic system serves as the framework for determining coordinates of points with respect to the earth. Modern global geodetic systems, such as the World Geodetic System 1984 (WGS 84), (see Section 5.4), have been established using techniques of satellite geodesy. Typically, they are defined by a geocentric Cartesian coordinate system with the Z-axis along the mean rotation axis of the earth and the X-axis through a longitude reference.

a. DoD World Geodetic Systems have an associated reference ellipsoid and a geoid model. The geoid model is sometimes used as a reference surface for elevations. World Geodetic Systems provide a consistent framework for determining geodetic positions.

b. Local geodetic systems are established by a variety of techniques. Typically, they consist of a geodetic datum (see Section 5.2) for horizontal control and a vertical datum (see Section 5.3) for elevations. Any given local geodetic system covers only a small fraction of the earth's surface. Local geodetic datums and vertical datums generally are independent of each other.

5.2 <u>Geodetic (horizontal) datums</u>. A geodetic datum has a earth-fixed reference ellipsoid that may have been fit, in some manner, to the surface of the earth in the area of interest. Geodetic datums have parameters that define the size and shape of the ellipsoid.

5.2.1 <u>Background</u>. Historically, local geodetic datums have an origin, on the surface of the earth, that relates the geodetic coordinate system to the ellipsoid. For these local datums, the ellipsoid semi-minor axis generally does not coincide with the earth's mean rotation axis. However, the North American Datum 1983 (NAD 83) was developed using the same satellite geodesy techniques used for global geodetic systems. This modern datum has an origin at the center of the reference ellipsoid, and the semi-minor axis coincides with the mean rotation axis of the earth. Without reference to the word vertical, the word datum, as used in this Technical Report, means geodetic (horizontal) datum.

5.2.2 <u>Geodetic datums and WGS 84</u>. Over the years, hundreds of geodetic datums have been developed for various locations. Currently, 113 local geodetic datums have been tied to WGS 84. Coordinate transformations from local datums to WGS 84 are discussed in Section 6.

5.3 Vertical datums and elevations. By definition, a **vertical datum** is a surface of zero elevation. Elevations are measured (positive upward) from the vertical datum. Ideally, a vertical datum would closely approximate the geoid.

5.3.1 Background. As a practical matter, it is impossible to access the geoid surface directly for use as a vertical datum. Historically, tide gage measurements were averaged over many years to establish the local mean sea level (MSL) references for vertical datums. The National Geodetic Vertical Datum of 1929 (NGVD 29), which is based on tide gage measurements and precise leveling surveys, is estimated to be within a few meters of the geoid. There is greater uncertainty in the relationships between the geoid and other local vertical datums.

5.3.2 Vertical datums and mean sea level (MSL). The exact relationships between different vertical datums are unknown. Therefore, when shifting datums, all elevations are considered referenced to MSL regardless of the vertical datum used to establish them. For Army applications, two exceptions to this rule are:

a. Conversion of elevations derived from WGS 72 geodetic height measurements to WGS 84 coordinates (see Section 6.7.1).

b. Civil works, engineering and construction applications are outside the scope of this Technical Report.

5.3.3 Relationship between local vertical and horizontal datums. Generally, vertical datums and horizontal datums are independently defined. For example, NGVD 29 is the major vertical datum, while the North American Datum 1927 (NAD 27) and NAD 83 are the major horizontal datums used in North America.

5.4 World Geodetic System 1984 (WGS 84). A world geodetic system provides the basic reference frame, geometric figure and gravimetric model for the earth, and provides the means for relating positions on various local geodetic systems to an earth-centered, earth-fixed (ECEF) coordinate system.

a. WGS 84 is the current standard DoD geodetic system. WGS 84 is the latest in a series of DoD-developed world geodetic systems. It is a replacement for WGS 72, the previous DoD standard ECEF world geodetic system.

b. The origin of the WGS 84 coordinate system is the center of mass of the earth. The WGS 84 Z-axis is parallel to the direction of the Conventional Terrestrial Pole (CTP), as defined by the Bureau International de l'Heure (BIH). The WGS 84 prime meridian is parallel to the BIH Zero meridian. The X-axis is the intersection of the plane of the WGS 84 reference meridian and the CTP equatorial

plane. The epoch is 1984.0. The Y-axis completes a right-handed, earth-fixed Cartesian coordinate system.

c. The WGS 84 reference ellipsoid is a geocentric ellipsoid of revolution. The WGS 84 coordinate system origin and axes also serve as the geometric center and the X, Y, and Z axes of the WGS 84 ellipsoid. WGS 84 ellipsoid parameters are given in Appendix A.

d. The WGS 84 geoid model can be used to convert between elevations and geodetic heights (see Section 5.5).

e. Datum shift techniques and parameters needed to relate positions on many local datums to WGS 84 were also developed as part of WGS 84 (see Section 6 and Appendices B and E).

f. Only those portions of WGS 84 related to the purpose of this Technical Report were described above. A comprehensive view of WGS 84 can be obtained from DMA TRs 8350.2 (1991), 8350.2-A (1987), and 8350.2-B (1987).

5.5 <u>WGS 84 referenced elevations</u>. Sometimes elevations must be generated from geodetic heights, without reference to a local vertical datum. For example, satellite positioning systems, like the Global Positioning System (GPS), establish coordinates geometrically and cannot directly measure elevations. An elevation can be obtained from a WGS 84 geodetic height using:

$$H \approx h_{WGS\ 84} - N_{WGS\ 84} \qquad (5.1)$$

where $N_{WGS\ 84}$ is the value of WGS 84 geoid separation for the WGS 84 geodetic position of the point. Appendix D contains a method for interpolating $N_{WGS\ 84}$ from a table of grided values. Table D.1 is a list of $N_{WGS\ 84}$ on a 10° x 10° grid. It is useful for moderate accuracy mapping and charting applications. TEC will furnish qualified users with either digital 30' x 30' or 1° x 1° $N_{WGS\ 84}$ tables or spherical harmonic expansion tables for survey and high accuracy applications. Requests should be sent to the address in the preface.

6.   DATUM SHIFTS

6.1   Introduction.   Several methods are available to shift coordinates from one geodetic datum to another. To develop these shifts, coordinates of one or more physical locations must be known on both datums. Typically, for the local datum to WGS 84 shifts presented in this Technical Report, the WGS 84 coordinates of local datum survey control points were determined using Doppler satellite observations. Two classes of datum shift methods are geometric transformation and polynomial fitting techniques. Only geometric transformations are discussed here.

a.   The generalized geometric transformation model assumes the origins of the two coordinate systems are offset from each other; the axes are not parallel; and there is a scale difference between the two datums. This **seven-parameter model** is discussed in Section 6.2. Data from at least three well-spaced positions are needed to derive a seven-parameter geometric transformation.

b.   The number of parameters in a geometric transformation model can be reduced. The commonly used **three-parameter model** neglects rotations between coordinate systems axes and scale differences between datums. Only the origins are offset. The three-parameter model is discussed in Section 6.3. Data from one or more positions are required to derive a three-parameter geometric transformation.

c.   The WGS 72 to WGS 84 transformation, given in Section 6.4, is a special five-parameter geometric transformation. When WGS 84 superseded WGS 72 as the DoD-preferred reference system, this transformation was created specifically to transform WGS 72 coordinates to WGS 84 coordinates.

d.   Heights used for datum shifting are treated in section 6.5. Shifting of vertical datums is discussed in Section 6.7.

6.1.1   Shifting between two local datums.   The datum shift methods given in this Technical Report are for shifting from a local datum to WGS 84 and from WGS 84 to a local datum. When shifts are needed between two local datums, shift the coordinates on the first local datum to WGS 84, then shift the WGS 84 coordinates to the second local datum.

6.1.1.1 Shifting from NAD 27 to NAD 83.
The software package NADCON has been recommended as the standard method for transformations between the North American Datum of 1927 (NAD27) and the North American Datum of 1983 (NAD 83) by The

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Federal Geodetic Control Committee (Federal Register.) Corpscon, a package incorporating NADCON, is mandated for the Army Corps of Engineers (ETL 1110-1-147.) Because NADCON is a survey tool it is beyond the scope of this Technical Report. Questions can be addressed to the Survey Division of TEC at the address in the preface, (703)355-2798, or to the National Geodetic Survey at (301)713-3178.

6.2 <u>Seven-parameter geometric transformation</u>. The seven-parameter transformation, the most general transformation between local and global Cartesian coordinates, is considered in DMA TR 8350.2-A (1987). The underlying assumption for this seven-parameter model is that the local and WGS 84 Cartesian axes differ by seven parameters: three rotational parameters $(\omega, \varepsilon, \psi)$, a scale change ΔS, and three origin shift parameters $(\Delta X, \Delta Y, \Delta Z)$. See Figure 6.1.
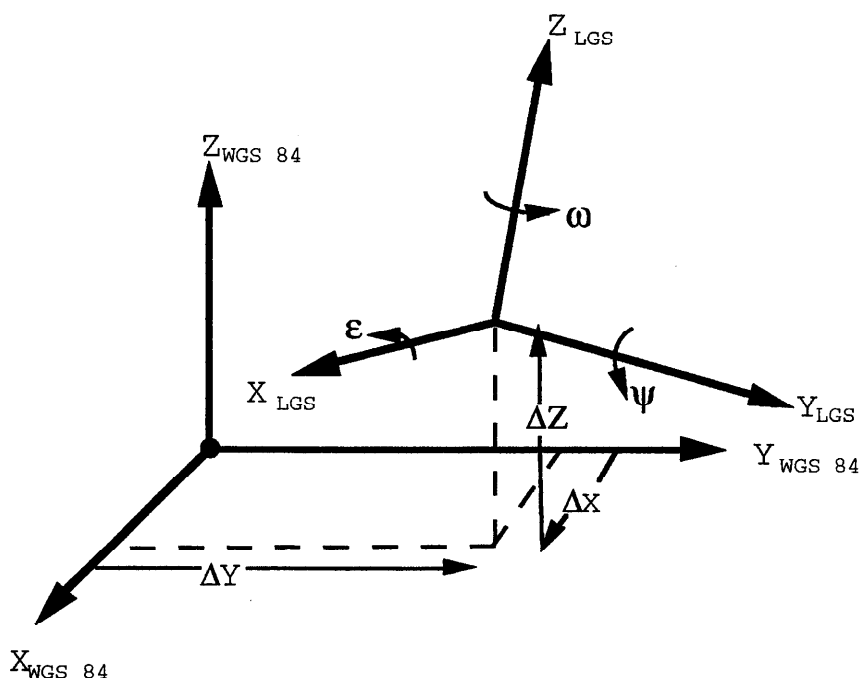
FIGURE 6.1. Relationship between coordinate axes in a seven-parameter geometric transformation.

6.2.1 <u>Transformation to WGS 84 Cartesian coordinates</u>. For the seven-parameter model, the local and WGS 84 Cartesian coordinates are related by

$$
\begin{aligned}
X_{WGS\ 84} &= X_{LGS} + \Delta X + \omega Y_{LGS} - \psi Z_{LGS} + \Delta S X_{LGS} \\
Y_{WGS\ 84} &= Y_{LGS} + \Delta Y - \omega X_{LGS} + \varepsilon Z_{LGS} + \Delta S Y_{LGS} \\
Z_{WGS\ 84} &= Z_{LGS} + \Delta Z + \psi X_{LGS} - \varepsilon Y_{LGS} + \Delta S Z_{LGS}
\end{aligned}
\tag{6.1}
$$

The subscript LGS designates Local Geodetic System coordinates and the subscript WGS 84 designates World Geodetic System 1984 coordinates.

6.2.2 <u>Transformation to LGS Cartesian coordinates</u>. The equations for converting WGS 84 geocentric Cartesian coordinates to local Cartesian coordinates are

$$X_{LGS} = X_{WGS\ 84} - \Delta X - \omega Y_{WGS\ 84} + \psi Z_{WGS\ 84} - \Delta S X_{WGS\ 84} \qquad (6.2)$$
$$Y_{LGS} = Y_{WGS\ 84} - \Delta Y + \omega X_{WGS\ 84} - \varepsilon Z_{WGS\ 84} - \Delta S Y_{WGS\ 84}$$
$$Z_{LGS} = Z_{WGS\ 84} - \Delta Z - \psi X_{WGS\ 84} + \varepsilon Y_{WGS\ 84} - \Delta S Z_{WGS\ 84}$$

6.2.3 <u>Parameter values</u>. The parameters ΔX, ΔY, ΔZ, ω, ε, ψ, and ΔS, for the European 1950 and Ordinance Survey of Great Britain 1936 datums are tabulated in Appendix E. Use of seven-parameter datum shifts for some applications in Europe and Great Britain is proscribed by STANAG 2211. Since the values of the rotational parameters ω, ε, and ψ are given in the table in arc seconds, these values must be converted to radians before using equation sets 6.1 and 6.2.

6.2.4 <u>Accuracy of seven-parameter transformation</u>. The RMS differences between WGS 84 surveyed and transformed coordinates using the seven-parameter datum transformations are given in Appendix E. For some datums, the accuracies for the seven-parameter transformation are not an improvement over those for the three-parameter transformation. Equation sets 6.1 and 6.2 provide good approximations to a rigorous orthogonal transformation when ω, ε, ψ, and ΔS are small.

6.3 <u>Three-parameter(ΔX, ΔY, ΔZ) geometric transformation</u>. Frequently, it is not practical to determine the parameters for a seven-parameter geometric datum shift or implement such a shift method. Often, only an origin shift, three-parameter model is used as shown in figure 6.2. The three-parameter datum transformation satisfies the requirements of most mapping and charting applications. The three-parameter model can be applied directly to Cartesian coordinates (see Sections 6.3.1 and 6.3.2) or used to compute shifts in geodetic coordinates via the Molodensky equations (see Section 6.3.3)

    a.   The datum shift parameters (ΔX, ΔY, ΔZ) are the coordinates of the origin of the local reference ellipsoid in the WGS 84 Cartesian coordinate system.

    b.   Datum shift parameters for 113 local geodetic systems are listed in Appendix B, Table B.1. The area covered by the datum shift parameters, the local datum ellipsoid, and

the estimates of the errors in the datum shift parameters also are given. Derivation of ΔX, ΔY, and ΔZ datum shift parameters is discussed in DMA TR 8350.2-A (1987).
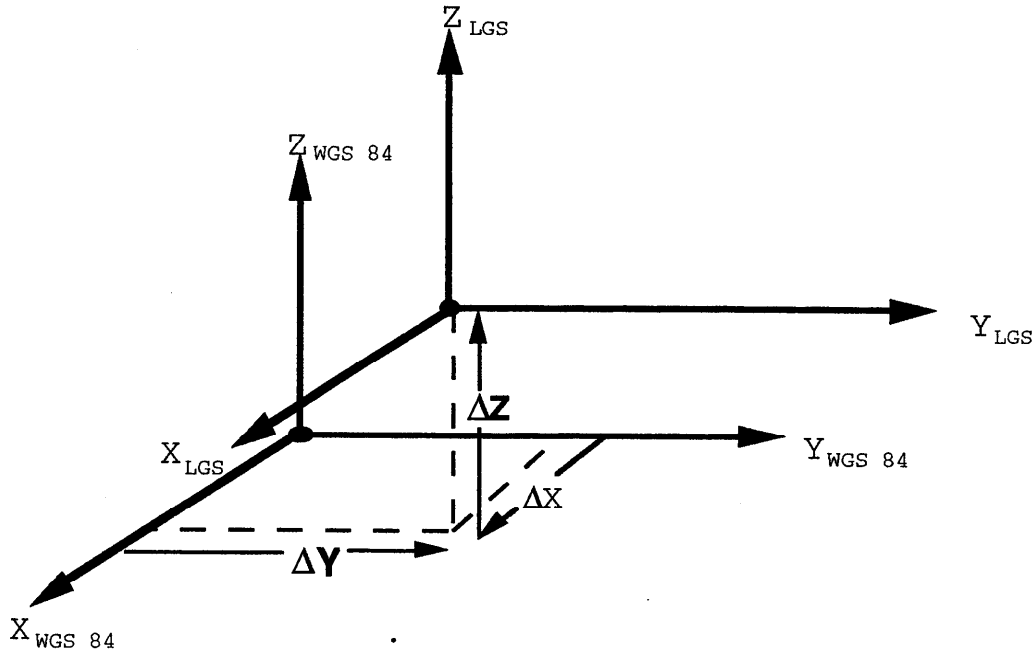
FIGURE 6.2. Relationship between coordinate axes in a three-parameter geometric transformation.

6.3.1 <u>Transformation to WGS 84 Cartesian coordinates</u>. In certain cases, ΔX, ΔY, and ΔZ datum shift parameters can be applied directly to convert local geodetic system coordinates to WGS 84 Cartesian coordinates as follows:

$$X_{WGS\ 84} = X_{LGS} + \Delta X$$
$$Y_{WGS\ 84} = Y_{LGS} + \Delta Y \qquad (6.3)$$
$$Z_{WGS\ 84} = Z_{LGS} + \Delta Z$$

Equation set 6.3 is a subset of equation set 6.1 with ω, ψ, ε, and ΔS set to zero.

6.3.1.1 <u>Three-step method: Transformation to WGS 84 geodetic coordinates</u>. Local geodetic coordinates can be shifted to WGS 84 coordinates in three steps:

    a. Convert local geodetic coordinates to local geocentric Cartesian coordinates using equation set 4.4. Ensure that local ellipsoid parameters are used.

    b. Shift the local geocentric Cartesian coordinates to WGS 84 geocentric Cartesian coordinates using equation set 6.3.

c. Convert WGS 84 geocentric Cartesian coordinates to WGS 84 geodetic coordinates using the methods discussed in Section 4.3.4.2. Ensure that WGS 1984 ellipsoid parameters are used.

6.3.2 <u>Transformation to local geodetic coordinates</u>. To convert WGS 84 Cartesian coordinates to the local geodetic system, apply ΔX, ΔY, and ΔZ datum shifts as follows

$$X_{LGS} = X_{WGS\ 84} - \Delta X \qquad (6.4)$$
$$Y_{LGS} = Y_{WGS\ 84} - \Delta Y$$
$$Z_{LGS} = Z_{WGS\ 84} - \Delta Z$$

Equation set 6.4 is a subset of equation set 6.2 with ω, ψ, ε, and ΔS set to zero.

6.3.2.1 <u>Three-step method: Transformation to local geodetic coordinates</u>. To shift WGS 84 geodetic coordinates to local geodetic coordinates

a. Convert WGS 84 geodetic coordinates to WGS 84 geocentric Cartesian coordinates using equation set 4.4. Insure that WGS 84 ellipsoid parameters are used.
b. Shift the WGS 84 geocentric Cartesian coordinates to local geocentric Cartesian coordinates using equation set 6.4.
c. Convert the local geocentric Cartesian coordinates to local geodetic coordinates using a method of Section 4.3.4.2. Insure that local ellipsoid parameters are used.

6.3.3 <u>Molodensky shifts</u>. The standard Molodensky method is an approximation to the three-step transformation methods of Sections 6.3.1.1 and 6.3.2.1. To use the Molodensky methods to transform local geodetic coordinates to WGS 84 geodetic coordinates, calculate $\Delta\phi$, $\Delta\lambda$, and Δh shifts using the standard Molodensky formulas Rapp, 1987.

$$\Delta\phi = [-\sin\phi\cos\lambda\,\Delta X - \sin\phi\sin\lambda\,\Delta Y + \cos\phi\,\Delta Z \qquad (6.5)$$
$$+ \frac{\varepsilon^2\sin\phi\cos\phi}{W}\Delta a$$
$$+ \sin\phi\cos\phi(2N + \varepsilon'^2 M\sin^2\phi)(1-f)\Delta f]/[M + h]$$

$$\Delta\lambda = \frac{[-\sin\lambda\,\Delta X + \cos\lambda\,\Delta Y]}{[(N + h)\cos\phi]}$$

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$$\Delta h = \cos \phi \cos \lambda \, \Delta X + \cos \phi \sin \lambda \, \Delta Y + \sin \phi \, \Delta Z$$
$$- W \Delta a + \frac{a(1-f)}{W} \sin^2 \phi \, (\Delta f)$$

in which

$$W^2 = 1 - \varepsilon^2 \sin^2 \phi$$
$$M = \frac{a(1 - \varepsilon^2)}{W^3}$$
$$N = \frac{a}{W}$$
$$\phi = \phi_{LGS}$$
$$\lambda = \lambda_{LGS}$$
$$\varepsilon^2 = \frac{a^2 - b^2}{a^2} = 2f - f^2$$

$$\varepsilon'^2 = \frac{a^2 - b^2}{b^2} = \frac{\varepsilon^2}{1 - \varepsilon^2}$$

The units of $\Delta \phi$ and $\Delta \lambda$ are radians. Note that $N = R_N$ of equation set 4.5.

Note: Most applications do not require shifting geodetic heights, so Δh does not need to be computed or applied. Geodetic heights, which are used to compute datum shift parameters, are a special construct that may not be well related to the local geodetic system. Do not use Δh to shift local geodetic heights. If this is a requirement, seek competent geodetic council.

a. Apply the $\Delta \phi$, $\Delta \lambda$, and Δh shifts as follows:

$$\phi_{WGS\ 84} = \phi_{LGS} + \Delta \phi_{LGS} \qquad\qquad (6.6)$$
$$\lambda_{WGS\ 84} = \lambda_{LGS} + \Delta \lambda_{LGS}$$
$$h_{WGS\ 84} = h_{LGS} + \Delta h_{LGS}$$

where $\Delta \phi_{LGS}$, $\Delta \lambda_{LGS}$, Δh_{LGS} are computed from equation set 6.5 using ϕ_{LGS}, λ_{LGS}, h_{LGS}, semi-major axis and flattening for the local geodetic system. ΔX, ΔY, and ΔZ are found in Appendix B. Δa and Δf are calculated as follows;

$$\Delta a = a_{WGS\ 84} - a_{LGS}$$
$$\Delta f = f_{WGS\ 84} - f_{LGS}$$

Note: See section 6.5 for use of h_{LGS}.

b. The Molodensky formulas can also be used to convert from WGS 84 to the local geodetic system:

$$\phi_{LGS} = \phi_{WGS\ 84} + \Delta\phi_{WGS\ 84}$$
$$\lambda_{LGS} = \lambda_{WGS\ 84} + \Delta\lambda_{WGS\ 84} \qquad\qquad (6.7)$$
$$h_{LGS} = h_{WGS\ 84} + \Delta h_{WGS\ 84}$$

where $\Delta\phi_{WGS\ 84}$, $\Delta\lambda_{WGS\ 84}$, and $\Delta h_{WGS\ 84}$ are computed using $\phi_{WGS\ 84}$, $\lambda_{WGS\ 84}$, $h_{WGS\ 84}$, WGS 84 semi-major axis and flattening, and the signs of ΔX, ΔY, and ΔZ, as given in Appendix B, are reversed. Δa and Δf are calculated as follows

$$\Delta a = a_{LGS} - a_{WGS\ 84}$$
$$\Delta f = f_{LGS} - f_{WGS\ 84}$$

Note that ellipsoid parameters a and f are found in Table A.1, Appendix A.

6.3.4 <u>ΔX, ΔY, and ΔZ shift values</u>. Appendix B (Table B.1) contains the ΔX, ΔY, and ΔZ values for the three-parameter methods of converting local geodetic systems to WGS 84. The accuracies, in terms of error estimates for ΔX, ΔY, and ΔZ, are given in Appendix B. A further discussion of the accuracy of these parameter shifts will be given in Section 6.3.5.

6.3.5 <u>Accuracy of ΔX, ΔY, and ΔZ shift parameters</u>. Error estimates for ΔX, ΔY, and ΔZ are tabulated in Appendix B. These estimates include the errors associated with the Doppler stations as well as the residual differences between transformed local system coordinates and the reference WGS 84 coordinates used to develop the datum shifts. (A Doppler station refers to a position where the WGS 84 coordinates were obtained by observing and processing of TRANSIT satellite data.)

a. For 38 datums, datum shifts were developed using a single Doppler station. For these datums, no measure of internal consistency of the datum shifts is calculable, and DMA has assigned an accuracy value of 25 meters in each coordinate. These error statistics do not reflect

any errors in the coordinates used to compute the datum shift.

b. Datum shifts for eight datums were calculated from sources other than Doppler stations for WGS 84 geodetic coordinates. No error statistics are provided for these datums.

c. Geodetic coordinate difference error estimates can be approximated from rectangular coordinate errors using the following spherical relationships:

$$\sigma_{\Delta\phi} = \sqrt{(\sigma_{\Delta X} \sin\phi \cos\lambda)^2 + (\sigma_{\Delta Y} \sin\phi \sin\lambda)^2 + (\sigma_{\Delta Z} \cos \phi)^2}$$

$$\sigma_{\Delta\lambda} = \sqrt{(\sigma_{\Delta X} \sin\lambda)^2 + (\sigma_{\Delta Y} \cos\lambda)^2} \tag{6.8}$$

$$\sigma_{\Delta h} = \sqrt{(\sigma_{\Delta X} \cos\phi \cos\lambda)^2 + (\sigma_{\Delta Y} \cos\phi \sin\lambda)^2 + (\sigma_{\Delta Z} \sin\phi)^2}$$

Note that these relationships neglect any correlations that may exist between the error estimates for ΔX, ΔY and ΔZ.

6.3.6   Local datum distortion. For some local datums, such as the Australian Geodetic datums, one set of ΔX, ΔY, and ΔZ datum shift parameters provides a consistent measure of the datum shifts throughout the datum. For other datums, ΔX, ΔY, and ΔZ shift parameters for the whole local datum, called mean-value datum shift parameters, poorly fit one or more regions within the datum. Regional variation in the datum shift parameters may result from using different, poorly connected surveys to establish the geodetic control on the local datum; from rotation of the local datum; and from scale differences. An extreme example of regional datum shift variation is the ΔY component for the ARC 1950 datum. The mean-value ΔY is -90 meters. However, ΔY ranges from -5 to -108 meters for different countries covered by the ARC 1950 datum. Appendix B (Table B.1) contains regional ΔX, ΔY, and ΔZ datum shift parameters for 17 local datums. For these datums, one should base the selection of mean or regional ΔX, ΔY, and ΔZ shift values on the accuracy required for the application. Only mean value three-parameter shifts are available for the remaining datums listed in Table B.1.

6.4   WGS 72 to WGS 84 transformation. WGS 84 has replaced WGS 72 as the accepted geodetic reference system for most DoD applications. The WGS 84 system, developed through a more extensive set of satellite-derived and surface data than was available at the start of WGS 72, is an improved geometric and gravitational model of the earth. Whenever possible, WGS 72 coordinates should be converted to the WGS 84 reference system.

6.4.1 <u>Selecting a conversion method</u>. Before selecting a WGS 72 to WGS 84 conversion method, find the source of the WGS 72 coordinates. Select the conversion method using the guidance in Table 6.1.

TABLE 6.1. Selecting a WGS 72 to WGS 84 Transformation Method.

<u>Source of WGS 72 Coordinates</u>	<u>Transformation Method</u>
1. Doppler satellite station.	Direct (6.4.2)
2. Local datum to WGS 72 transformation. Local datum coordinates known.	Local datum to WGS 84 transformation (6.2, 6.3, 6.6)
3. Local datum to WGS 72 transformation. Local datum coordinates unknown.	Two-Step (6.4.3) Shift method known
4. Local datum to WGS 72 transformation using localized datum shift parameters derived for nearby WGS 72 Doppler station. Local datum coordinates unknown.	Direct (6.4.2)
5. Unknown.	Direct (6.4.2) Accuracy unknown

6.4.2 <u>Direct WGS 72 to WGS 84 transformation</u>. The direct WGS 72 to WGS 84 transformation reflects fundamental changes between the WGS 72 and WGS 84 systems. Changes include a shift in coordinate system origin, a shift in longitudinal reference (initial meridian), a change in system scale, and changes in ellipsoidal parameters. For WGS 72 coordinates not originating in the WGS 72 system (i.e., non-Doppler-derived coordinates), take care when employing the direct WGS 72 to WGS 84 transformation, as any inaccuracies and uncertainties inherent in the WGS 72 coordinates will be directly transferred into the derived WGS 84 coordinates. These inaccuracies and uncertainties are generally a result of local geodetic system to WGS 72 coordinate transformations using mean rather than localized datum shifts. Thus, every effort should be made to find the source of all WGS 72 coordinates before doing a direct WGS 72 to WGS 84 transformation.

6.4.2.1 <u>Conversion equations</u>. It is important to note that the relationship between the WGS 72 coordinates and the WGS 84 coordinates cannot be described by a three-parameter model (see Section 6.3).

a. The equations for the direct WGS 72 to WGS 84 geodetic coordinate transformation are

$$\phi_{WGS\ 84} = \phi_{WGS\ 72} + \Delta\phi \qquad\qquad (6.9)$$

$$\lambda_{WGS\ 84} = \lambda_{WGS\ 72} + \Delta\lambda$$

$$h_{WGS\ 84} = h_{WGS\ 72} + \Delta h$$

where

$$\Delta\phi = \frac{4.5\ \cos\phi_{WGS\ 72}}{a_{WGS\ 72}\ Q} + \frac{\Delta f\ \sin 2\phi_{WGS\ 72}}{Q} \quad \text{(arc seconds)}$$

$$\Delta\lambda = 0.554 \qquad \text{(arc seconds)}$$

$$\Delta h = 4.5\ \sin\phi_{WGS\ 72} + a_{WGS\ 72}\ \Delta f\ \sin^2\phi_{WGS\ 72} - \Delta a + \Delta r \quad \text{(meters)}$$

$$\Delta f = 0.3121057 \times 10^{-7}$$

$$a_{WGS\ 72} = 6378135 \quad \text{(meters)}$$

$$\Delta a = 2.0 \qquad \text{(meters)}$$

$$\Delta r = 1.4 \qquad \text{(meters)}$$

$$Q = \frac{\pi}{180 \cdot 3600}$$

b. The equations for the direct WGS 84 to WGS 72 geodetic coordinate transformation are:

$$\phi_{WGS\ 72} = \phi_{WGS\ 84} + \Delta\phi$$
$$\lambda_{WGS\ 72} = \lambda_{WGS\ 84} + \Delta\lambda \qquad\qquad (6.10)$$
$$h_{WGS\ 72} = h_{WGS\ 84} + \Delta h$$

where

$$\Delta\phi = \frac{-4.5\ \cos\phi_{WGS\ 84}}{a_{WGS\ 84}\ Q} - \frac{\Delta f\ \sin 2\phi_{WGS\ 84}}{Q} \quad \text{(arc seconds)}$$

$$\Delta\lambda = -0.554 \qquad \text{(arc seconds)}$$

$$\Delta h = -4.5\ \sin\phi - a_{WGS\ 84}\ \Delta f\ \sin^2\phi_{WGS\ 84} + \Delta a - \Delta r \quad \text{(meters)}$$

$$\Delta f = 0.3121057 \times 10^{-7}$$

$$a_{WGS\ 84} = 6378137 \quad \text{(meters)}$$

$$\Delta a = 2.0 \quad \text{(meters)}$$

$$\Delta r = 1.4 \quad \text{(meters)}$$

$$Q = \frac{\pi}{180 \cdot 3600}$$

Elevation based on the WGS 84 geoid model will be different from elevations based on the WGS 72 geoid model due to differences in the geoid models (see Section 6.7.1 for conversion).

6.4.2.2 <u>Estimated errors $(\sigma_{\Delta\phi}, \sigma_{\Delta\lambda}, \sigma_{\Delta h})$</u>. The estimated errors in a direct WGS 72 to WGS 84 transformation are

$$\sigma_{\Delta\phi} = \sigma_{\Delta\lambda} = 3 \text{ m} \qquad (6.11)$$

$$\sigma_{\Delta h} = 3 \text{ m to 4 m}$$

These error estimates do not include errors in the original WGS 72 coordinates. The accuracy estimate for $\sigma_{\Delta h}$ reflects discrepancies between WGS 72 and WGS 84 geoid models.

6.4.3 <u>Two-step WGS 72 to WGS 84 transformation</u>. If WGS 72 coordinates are known to have been derived from local geodetic system coordinates using WGS 72 ΔX, ΔY, and ΔZ datum shift parameters, the preferred method for transforming them to the WGS 84 system is to do a two-step conversion. First, do a WGS 72 to local transformation using the inverse of the transformation originally used to shift local datum coordinates to WGS 72. This step will eliminate the inaccuracies introduced using local to WGS 72 datum shifts. Next, do a local geodetic system to WGS 84 conversion to produce coordinates within the WGS 84 system (see Sections 6.2, 6.3, or 6.6). Seppelin (1974) provides mean value local datum to WGS 72 datum shift parameters.

6.5 <u>Approximating geodetic heights for datum Transformation</u>. When doing the Molodensky approximation (Section 6.3.3), ellipsoidal height (h) is often unavailable. MSL height (H) can be substituted for h without introducing errors in horizontal position significant to mapping and charting applications. The inpact of this substitution is so small that zero can be used in place of h, if H is unavailable.

NOTE: The generation and usage of local datum ellipsoidal heights are outside the scope of this Technical Report. For such applications, consult TEC at the address in the foreward.

6.6   Multiple regression equations (MRE). The use of MREs for datum shifting is not recommended.

6.7   Vertical datum shifts. Because exact relationships between vertical datums are unknown, shifting elevations between different vertical datums should not be attempted, except for WGS 72 elevations computed from measured WGS 72 geodetic heights (see Section 6.7.1.) Otherwise, the following are assumed to be the same:

a. Elevations on different local vertical datums.
b. Elevations computed from measured WGS 84 geodetic heights and WGS 84 geoid separations.

It is estimated that errors resulting from this assumption will be one to two meters. Should the accuracy be insufficient, contact TEC at the address listed in the foreword for assistance.

6.7.1   Satellite-derived WGS 72 elevations based on measured WGS 72 geodetic heights. The WGS 84 geoid model is a closer approximation to the geoid than the older WGS 72 geoid model. Elevations obtained from WGS 72 geodetic height measurements and geoid separations may be improved by converting to WGS 84 using

$$H = H_{WGS\ 72} + N_{WGS\ 72} + \Delta h - N_{WGS\ 84} \qquad (6.12)$$

where $H_{WGS\ 72}$ is the elevation computed using the WGS 72 geoid, $N_{WGS\ 72}$ is the WGS 72 geoid separation, and Δh is computed using equation set (6.9).

6.7.2   Survey-derived elevations. Elevations obtained by conventional leveling surveys, for points with WGS 72 geodetic horizontal coordinates, should not be converted to WGS 84 elevations. Simplified procedures to convert WGS 72-derived elevations to WGS 84 are not recommended.

6.8   Unlisted datums. This Technical Report provides datum shift parameters between 113 local datums and WGS 84. Regional shift parameters are also given for portions of many of these datums. Contact TEC at the address given in the foreword for any datums and their respective shift parameters that are not listed in this Technical Report.

7. MAP PROJECTIONS

7.1 <u>Introduction</u>. Map projections are used to represent the earth's features or points on a map or in a plane coordinate system. A map projection is a systematic representation of a part of the earth's surface on a plane.

7.1.1 <u>The map projection process</u>. The map projection process can be done in three steps:

(1) The positions of points on the surface of the earth (ϕ, λ, h) are reduced to positions on the reference ellipsoid (ϕ, λ, h = 0).
(2) Next, a map projection is chosen and used to transform the feature's geodetic coordinates (ϕ, λ, h = 0) into planar map projection coordinates (x, y).
(3) Finally, the map projection coordinates of the features may be plotted, on a map, or stored for future use.

7.1.2 <u>Properties of projections</u>. Once the earth's surface is projected on a plane, one can observe such properties as distance, angle, direction, shape, and size. None of the projections described in this Technical Report faithfully preserves all these characteristics. Some type of distortion is always introduced by a map projection. However, a given map projection can preserve some of these properties, and if a sufficiently small part of the earth's surface is displayed, then other properties can be represented approximately. Map projections presented in this Technical Report have areas of coverage beyond which the projection should not be used due to increased distortion.

7.1.3 <u>Mapping equations</u>. The map projections discussed in this Technical Report have mapping equations that are defined on a reference ellipsoid and are functions of geodetic coordinates (ϕ, λ). Geodetic height, h, does not enter into the projection equations since h = 0 on a reference ellipsoid. The equations that express projection coordinates, x and y, in terms of (ϕ, λ) are sometimes called **mapping equations**.

7.1.4 <u>Conformal projection</u>. Each projection presented in this Technical Report has the important property of being conformal. A **conformal projection** is angle preserving. If two lines on the ellipsoid meet at an angle of θ, then in a conformal projection the image of these lines on the map meet at the same angle, θ.

7.1.5 <u>Scale factor</u>. Two quantities associated with map projections, which are often confused, are the scale factor of the projection and the map scale. The **scale factor** of the projection

is used in precise direction and distance calculations using a map
projection's coordinates. The scale factor of a projection is the
ratio of arc length along a differentially small line in the plane
of the projection to arc length on the ellipsoid. This number
depends on both the location of the point and on the direction of
the line along which arc length is being measured. However, for
conformal projections, the scale factor of the projection is
independent of the direction of the line and depends only upon the
location of the point. Since all projections treated in this
Technical Report are conformal, the scale of the projection will
be referred to as the point scale factor (k).

7.1.6 <u>Map scale</u>. The **map scale** is the approximate number
for converting map distances to terrestrial distances and is
usually published on the map. Technically, the map scale is the
ratio of a distance on the map to the corresponding distance on
the ellipsoid. Although the map scale changes over the area of a
map sheet, it is customary to assign a single value to denote the
map's scale. This value is the true map scale along at least one
line on the map. Usually, the map scale is given by a ratio 1:n,
so that the larger the map scale, the smaller the value of n. For
military maps of scale greater than 1:100,000 (n < 100,000), the
map scale changes only slightly over a map sheet, and the map
scale approximates the ratio of distance on the map to distance on
the ellipsoid.

7.1.7 <u>Convergence of the meridian</u>. The **convergence of the
meridian** (γ), often called grid convergence, is used to convert
between direction in the map projection (grid azimuth) and
geodetic azimuth. The convergence of the meridian at a point on a
map is the angle between the projection of the meridian at that
point and the projection North axis (typically represented by the
y axis of the projection). The convergence of the meridian is
positive in the clockwise direction. Convergence of the meridian
should not be confused with declination, which is the difference
between geodetic and magnetic north at the point.

7.1.8 <u>Information sources</u>. Information on map projections
can be found in Krakiwsky (1973); NSWC/DL TR-3624 (1977); U.S.
Department of Commerce, Coast and Geodetic Survey, Special
Publication No. 251 (1979); and Geological Survey Professional
Paper 1395 (1987).

7.2 <u>Mercator projection</u>. The Mercator projection is a
conformal projection for which the point scale factor is one along
the equator. The equator lies on the line y = 0. This projection
is not defined at the poles.

7.2.1 <u>Meridians and parallels</u>. Meridians and parallels
provide the framework for the Mercator projection. Meridians are
projected as parallel straight lines that satisfy the equation x =
constant. Evenly spaced meridians on the ellipsoid project to

evenly spaced straight lines on the projection. Parallels are
projected as parallel straight lines perpendicular to meridians
and satisfy the equation y = constant. Evenly spaced parallels on
the ellipsoid project to unevenly spaced parallels on the
projection. The spacing between projected parallels increases
with distance from the equator. See figure 7.1.
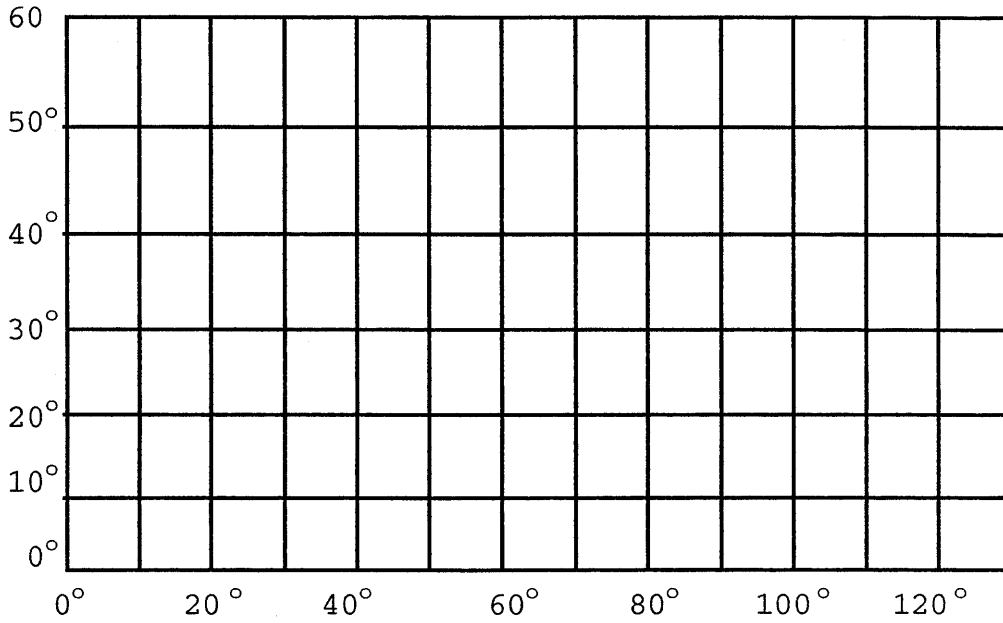
FIGURE 7.1. Meridians (longitude) and parallels (latitude)
in the Mercator projection.

7.2.2 Mercator mapping equations.

7.2.2.1 Finding (x,y). The following discussion of the
Mercator mapping equations can be found in Geological Survey
Professional Paper 1395 (1987). First, choose a central meridian
(λ_o) that represents the zero point on the x axis. Given a point
(ϕ, λ), in radians, on the ellipsoid, the corresponding point
(x, y) in the plane is

$$x = a(\lambda - \lambda_o) \tag{7.1}$$

$$y = a \ln \left[\tan\left(\frac{\pi}{4} + \frac{\phi}{2}\right)\left(\frac{1 - \varepsilon \sin \phi}{1 + \varepsilon \sin \phi}\right)^{\frac{\varepsilon}{2}} \right]$$

or

$$y = \left(\frac{a}{2}\right) \ln \left[\left(\frac{1 + \sin \phi}{1 - \sin \phi}\right)\left(\frac{1 - \varepsilon \sin \phi}{1 + \varepsilon \sin \phi}\right)^{\varepsilon} \right] \tag{7.2}$$

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where   a = semi-major axis of the ellipsoid
         ε = eccentricity
         λ_o = longitude of the central meridian

ϕ, λ and λ_o are expressed in radians.

7.2.2.2 <u>Point scale factor and convergence of the meridian</u>.
The equations for the point scale factor (k) and the convergence
of the meridian (γ) are

$$k = \frac{a}{N \cos \phi} \qquad (7.3)$$

$$\gamma = 0$$

where $$N = \frac{a}{\sqrt{1 - \varepsilon^2 \sin^2 \phi}}$$

7.2.2.3 <u>Finding (ϕ, λ)</u>. Next, consider the inverse problem
of finding (ϕ, λ) when (x, y) is known. The longitude can be
found from

$$\lambda = \frac{x}{a} + \lambda_o \qquad (7.4)$$

where both λ and λ_o are expressed in radians. The value of ϕ can
be found iteratively. Let an intermediate parameter t be defined

$$t = e^{(\frac{-y}{a})} \qquad (7.5)$$

After an initial value of

$$\phi_o = \frac{\pi}{2} - 2 \arctan t \qquad (7.6)$$

iterate the following equation for ϕ:

$$\phi_{n+1} = \frac{\pi}{2} - 2 \arctan \left(t \left[\frac{(1 - \varepsilon \sin \phi_n)}{(1 + \varepsilon \sin \phi_n)} \right]^{\frac{\varepsilon}{2}} \right) \qquad (7.7)$$

Substitute the initial value (ϕ_o) into equation 7.7 to find the next candidate ϕ_1. Similarly, substitute ϕ_1 into equation 7.7 to obtain the updated candidate, ϕ_2. Continue this process until the difference between successive values of ϕ_n is sufficiently small. Then set $\phi = \phi_n$. All angles are measured in radians.

7.2.3 Accuracy. The Mercator equations for x, y, k, γ, and λ are exact. The iterative equation for ϕ can be updated until any desired accuracy is obtained.

7.2.4 Area of coverage. In the Mercator projection, as the latitude (ϕ) approaches the poles, the y coordinate approaches infinity. Area and length distortion increases with distance from the equator. For example, the point scale factor is approximately 2 at 60° latitude and 5.7 at 80° latitude.

7.3 Transverse Mercator (TM) projection. The Transverse Mercator projection is a conformal projection for which the point scale factor equals one along the central meridian. The line y = 0 is the projection of the equator, and the line x = 0 is the projection of the central meridian. The Universal Transverse Mercator (UTM) projection (see Section 7.4) is a modification of the Transverse Mercator projection.

7.3.1 Meridians and parallels. Both the central meridian and the equator are represented as straight lines. No other meridian or parallel is projected onto a straight line. Since the point scale factor is one along the central meridian, this projection is most useful near the central meridian. Scale distortion increases away from this meridian. (See Figure 7.2.)
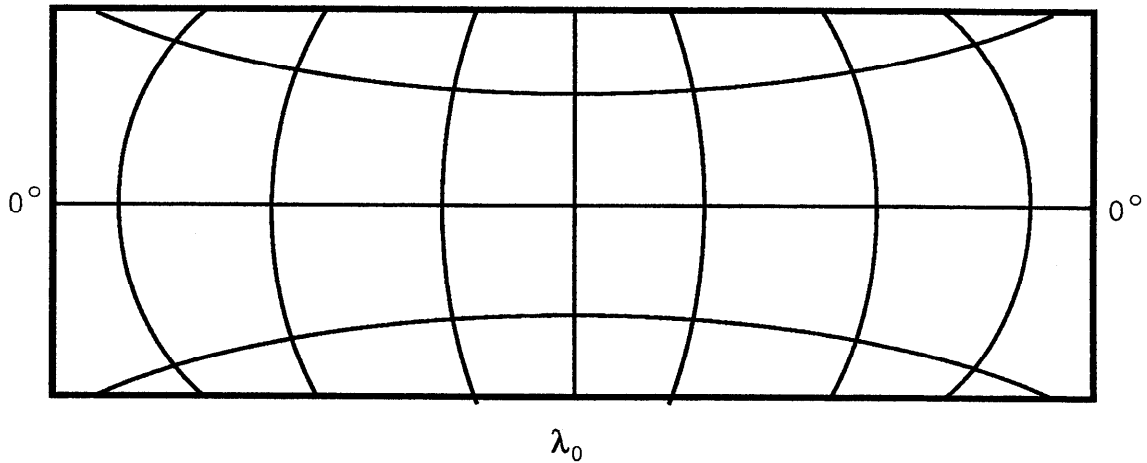
FIGURE 7.2.   Meridians and parallels in the Transverse
              Mercator Projection. λ_0 is the central
              meridian.

7.3.2 Transverse Mercator mapping equations.

7.3.2.1 **Finding (x,y)**. The presentation here follows U.S.
Department of Commerce, Coast and Geodetic Survey, Special
Publication No. 251 (1979). A point with geodetic coordinates (ϕ,
λ) in radians has Transverse Mercator coordinates (x,y) that are
given by

$$x = N \Lambda \cos\phi + \frac{N \Lambda^3 \cos^3 \phi}{6} (1 - t^2 + \eta^2) \qquad (7.8)$$

$$+ \frac{N \Lambda^5 \cos^5\phi}{120} (5 - 18t^2 + t^4 + 14\eta^2 - 58t^2\eta^2)$$

$$y = S_\phi + \frac{N \Lambda^2}{2} \sin \phi \cos \phi$$

$$+ \frac{N \Lambda^4}{24} \sin \phi \cos^3 \phi (5 - t^2 + 9\eta^2 + 4\eta^4)$$

$$+ \frac{N \Lambda^6}{720} \sin \phi \cos^5 \phi (61 - 58t^2 + t^4 + 270\eta^2 - 330t^2\eta^2)$$

where

$$N = \frac{a}{\sqrt{1 - \varepsilon^2 \sin^2 \phi}}$$

(7.9)

$$t = \tan \phi$$

$$\eta = \varepsilon' \cos \phi$$

$$\varepsilon'^2 = \frac{\varepsilon^2}{(1 - \varepsilon^2)}$$

$$\varepsilon^2 = \frac{a^2 - b^2}{a^2}$$

$$S_\phi = a[A_o\phi - A_2 \sin 2\phi + A_4 \sin 4\phi - A_6 \sin 6\phi + A_8 \sin 8\phi]$$

$$A_o = 1 - \frac{1}{4}\varepsilon^2 - \frac{3}{64}\varepsilon^4 - \frac{5}{256}\varepsilon^6 - \frac{175}{16384}\varepsilon^8$$

$$A_2 = \frac{3}{8}(\varepsilon^2 + \frac{1}{4}\varepsilon^4 + \frac{15}{128}\varepsilon^6 - \frac{455}{4096}\varepsilon^8)$$

$$A_4 = \frac{15}{256}(\varepsilon^4 + \frac{3}{4}\varepsilon^6 - \frac{77}{128}\varepsilon^8)$$

$$A_6 = \frac{35}{3072}(\varepsilon^6 - \frac{41}{32}\varepsilon^8)$$

$$A_8 = - \frac{315}{131072} \varepsilon^8$$

$$\Lambda = \lambda - \lambda_o \quad \text{in radians}$$

$$\lambda_o = \text{central meridian in radians}$$

At the poles, where $\phi = +$ or $- \pi/2$, the Transverse Mercator coordinates are $x = 0$ and $y = S_\phi$, a quadrant arc on the ellipsoid.

7.3.2.2 Finding (ϕ, λ).

7.3.2.2.1 Finding the footpoint latitude (ϕ_1). To find (ϕ, λ), given x and y, first compute the footpoint latitude (ϕ_1) using the following iterative method (See Krakiwsky 1973). The equation to be iterated to yield the footpoint latitude (ϕ_1) is

$$\phi' = \phi - \frac{S_\phi - y}{S'_\phi}$$

(7.10)

where

$$S'_\phi = a[A_o - 2A_2\cos2\phi + 4A_4\cos4\phi - 6A_6\cos6\phi + 8A_8\cos8\phi]$$

(7.11)

and S_ϕ, A_0, A_2, A_4, A_6, and A_8 are defined in equation set 7.9.

Let the first candidate for ϕ_1 be $\phi = y/a$. When this value of ϕ is substituted into equation 7.10, the updated candidate for ϕ_1 is ϕ'. Similarly, the substitution of ϕ' for ϕ in equation 7.10

gives the next updated value of ϕ'. When successive values of ϕ' are sufficiently close, then the footpoint latitude is found by setting $\phi_1 = \phi'$. A non-iterative method for finding ϕ_1 can be found in NOAA Report NOS NGS 5 (1989).

7.3.2.2.2 <u>Finding (ϕ, λ) from ϕ_1</u>. Once ϕ_1 is determined, (ϕ, λ) can be found from the approximations

$$\phi = \phi_1 - \frac{t_1 N_1}{R_1} \left[\frac{1}{2} \left(\frac{x}{N_1}\right)^2 - \frac{B_4}{24} \left(\frac{x}{N_1}\right)^4 + \frac{B_6}{720} \left(\frac{x}{N_1}\right)^6 \right] \qquad (7.12)$$

$$\lambda = \lambda_o + \sec \phi_1 \left[\frac{x}{N_1} - \frac{B_3}{6} \left(\frac{x}{N_1}\right)^3 + \frac{B_5}{120} \left(\frac{x}{N_1}\right)^5 \right]$$

where

$$B_3 = 1 + 2t_1^2 + \eta_1^2 \qquad (7.13)$$

$$B_4 = 5 + 3t_1^2 + \eta_1^2 - 4\eta_1^4 - 9t_1^2\eta_1^2$$

$$B_5 = 5 + 28t_1^2 + 24t_1^4 + 6\eta_1^2 + 8t_1^2\eta_1^2$$

$$B_6 = 61 + 90t_1^2 + 46\eta_1^2 + 45t_1^4 - 252t_1^2\eta_1^2$$

$$- 3\eta_1^4 - 66t_1^2\eta_1^4 - 90t_1^4\eta_1^2 + 225t_1^4 \eta_1^4$$

$$\varepsilon'^2 = \frac{\varepsilon^2}{(1 - \varepsilon^2)}$$

$$\eta_1^2 = \varepsilon'^2 \cos^2 \phi_1$$

$$t_1 = \tan \phi_1$$

$$N_1 = \frac{a}{\sqrt{1 - \varepsilon^2 \sin^2 \phi_1}}$$

$$R_1 = \frac{a(1 - \varepsilon^2)}{\left(\sqrt{1 - \varepsilon^2 \sin^2 \phi_1}\right)^3}$$

If $\phi_1 = \pm \frac{\pi}{2}$ then $\phi = \pm 90°$ (a pole) and λ is undefined.

7.3.2.3 Point scale factor as a function of x and ϕ_1. As a function of x and ϕ_1, the point scale factor is given by

$$k = 1 + \frac{C_2}{2} \left(\frac{x}{N_1}\right)^2 + \frac{C_4}{24} \left(\frac{x}{N_1}\right)^4 + \frac{1}{720} \left(\frac{x}{N_1}\right)^6 \tag{7.14}$$

where

$$C_2 = 1 + \eta_1^2 \tag{7.15}$$

$$C_4 = 1 + 6\eta_1^2 + 9\eta_1^4 - 24t_1^2 \eta_1^2$$

and t_1 , η_1, and N_1 are given by equation set 7.13.

7.3.2.4 Point scale factor as a function of (ϕ,λ). As a function of (ϕ, λ), the point scale factor is given by

$$k = 1 + \frac{\Lambda^2 \cos^2 \phi}{2} D_2 + \frac{\Lambda^4 \cos^4 \phi}{24} D_4 + \frac{\Lambda^6 \cos^6 \phi}{720} D_6 \tag{7.16}$$

in which

$$D_2 = 1 + \eta^2 \tag{7.17}$$
$$D_4 = 5 - 4t^2 + 14\eta^2 + 13\eta^4 - 28t^2\eta^2 - 48t^2\eta^4$$
$$D_6 = 61 - 148t^2 + 16t^4$$

and Λ, t, and η are given in equation set 7.9.

7.3.2.5 Convergence of the meridian in terms of x and ϕ_1.
The convergence of the meridian, in terms of x and ϕ_1, is

$$\gamma = \frac{t_1 x}{N_1} \left[1 - \frac{E_2}{3} \left(\frac{x}{N_1}\right)^2 + \frac{E_4}{15} \left(\frac{x}{N_1}\right)^4 - \frac{E_6}{315} \left(\frac{x}{N_1}\right)^6 \right] \tag{7.18}$$

where

$$E_2 = 1 + t_1^2 - \eta_1^2 - 2\eta_1^4 \tag{7.19}$$

$$E_4 = 2 + 5t_1^2 + 2\eta_1^2 + 3t_1^4 + t_1^2\eta_1^2 + 9\eta_1^4 - 7t_1^2\eta_1^4$$

$$E_6 = 17 + 77t_1^2 + 105t_1^4 + 45t_1^6$$

and t_1, η_1, and N_1 are given by equation set 7-13.

7.3.2.6 Convergence of the meridian in terms of (ϕ, λ). The convergence of the meridian, in terms of (ϕ, λ), is

$$\gamma = \Lambda \sin \phi \left(1 + F_2 \frac{\Lambda^2 \cos^2 \phi}{3} + F_4 \frac{\Lambda^4 \cos^4 \phi}{15} + F_6 \frac{\Lambda^6 \cos^6 \phi}{315} \right) \quad (7.20)$$

in which

$$F_2 = 1 + 3\eta^2 + 2\eta^4 \qquad\qquad (7.21)$$
$$F_4 = 2 - t^2 + 15\eta^2 + 35\eta^4 - 15t^2\eta^2 - 50t^2\eta^4$$
$$F_6 = 17 - 26t^2 + 2t^4$$

and Λ, t, and η are given in equation set 7.9.

7.3.3 **Accuracy**. The Transverse Mercator equations for x, y, ϕ, λ, S_ϕ, ϕ_1, k, and γ are approximations. Within 4° of the central meridian, the equations for x, y, ϕ, and λ have an error of less than 1 centimeter. (See Geological Survey Professional Paper 1395, 1987.)

7.3.4 **Area of coverage**. The equations given here for the Transverse Mercator projection can be used within 4° of the central meridian. If these equations are used with UTM coordinates, then consult Section 7.4.1 for guidance on the area of coverage for UTM coordinates.

7.4 **Universal Transverse Mercator (UTM) projection**. The UTM projection is a family of projections that differ from the Transverse Mercator projection in these ways (see Figure 7.3):

   a. The longitudes of the central meridians (in degrees) are 3 + 6n, n = 0,...59.
   b. The point scale factor along the central meridian is 0.9996.
   c. The y value, called the northing, has an origin of 0 meters in the northern hemisphere at the equator. The y value, called the southing, has an origin of 10,000,000 meters at the equator and decreases toward the pole in the southern hemisphere.
   d. The x value, called the easting, has an origin of 500,000 meters at the central meridian.

7.4.1 **UTM zones**. For the UTM grid system, the ellipsoid is divided into 60 longitudinal zones of 6 degrees each. Zone number one lies between 180° E and 186° E. The zone numbers increase consecutively in the eastward direction. It is intended that a UTM projection should be used only in one of the 6-degree zones, plus the overlap area. In each zone, the UTM projection extends for a 40 km overlap into the two adjacent zones. Moreover, UTM coordinates are used only between 80° S and 84° N plus the overlap region. The UTM projection extends to 84° 30' N and 80° 30' S to overlap with the Universal Polar Stereographic (UPS) projections.

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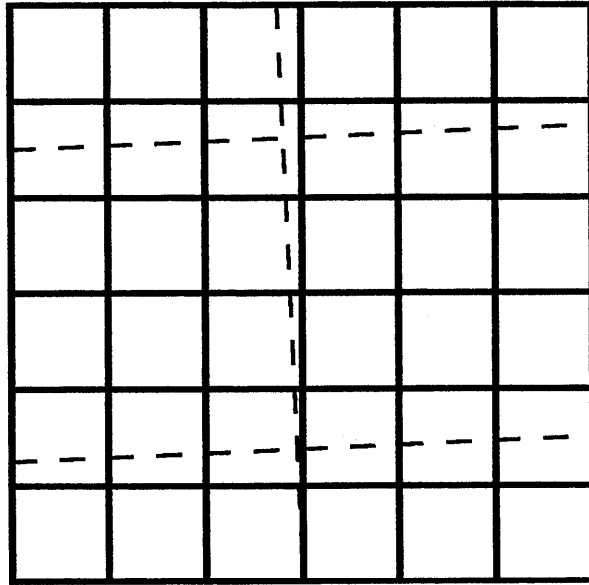
FIGURE 7.3. Meridians and parallels (dashed) and a
Universal Transverse Mercator Grid.

    7.4.1.1 <u>Finding the UTM zone and central meridian</u>. To
find UTM coordinates from geodetic coordinates, first compute the
UTM zone. A standard zone can be computed directly from geodetic
longitude. If the geodetic longitude (λ) is expressed in positive
radians, then the zone (z) is

$$z = \text{Greatest Integer} \leq \left(31 + \frac{180\lambda}{6\pi}\right) \quad \text{for } 0 \leq \lambda < \pi \,, \qquad (7.22)$$

$$z = \text{Greatest Integer} \leq \left(\frac{180\lambda}{6\pi} - 29\right) \quad \text{for } \pi \leq \lambda < 2\pi$$

Once the zone (z) is known, the central meridian is given by
equation set 7.23. If a longitude lies on a UTM zone boundary,
then the user must choose the desired UTM zone. If a map is being
used, then the UTM zone appears on the map. The central meridian
of zone (z) is

$$\lambda_o = (6z - 183)\frac{\pi}{180} \qquad \text{for } (z \geq 31) \qquad (7.23)$$

$$\lambda_o = (6z + 177)\frac{\pi}{180} \qquad \text{for } (z \leq 30)$$

with λ_o expressed in radians.

7.4.1.2 Non-standard width UTM zones. Non-standard width UTM zones are used in North Atlantic regions and Norway as delineated in Table 7.1 (see DMA TM 8358.1 (1990). These non-standard zones may also be applied in the 30' UPS overlap region.

TABLE 7.1. Non-standard zone limits.

Zone	Latitude		Longitude		Central Meridian
	Lower	Upper	West	East	
31	56° N	64° N	0° E	3° E	3° E
32	56	64	3	12	9
31	72	84	0	9	3
32	72	84		not used	
33	72	84	9	21	15
34	72	84		not used	
35	72	84	21	33	27
36	72	84		not used	
37	72	84	33	42	39

7.4.2 Reference source. The standard references for the UTM projection are DMA TM 8358.1 (1990) and DMA TM 8358.2 (1989).

7.4.3 UTM equations.

7.4.3.1 Finding UTM coordinates (x_{UTM}, y_{UTM}). Having found the central meridian, next compute Transverse Mercator coordinates (x, y), using equation sets 7.8 and 7.9. Then, the UTM coordinates can be found from:

$$x_{UTM} = 0.9996 \ x_{TM} + 500,000 \qquad\qquad (7.24)$$
$$y_{UTM} = 0.9996 \ y_{TM} \qquad\qquad \text{(Northern hemisphere)}$$
$$y_{UTM} = 0.9996 \ y_{TM} + 10,000,000 \quad \text{(Southern hemisphere)}$$
$$k_{UTM} = 0.9996 \ k_{TM}$$
$$\gamma_{UTM} = \gamma_{TM}$$

7.4.3.2 Computing geodetic coordinates. To compute geodetic coordinates from UTM coordinates, a zone, and a hemisphere, first solve equation set 7.25 for (x_{TM}, y_{TM}), equation set 7.23 for λ_o, and then compute (ϕ, λ) from equation set 7.12.

$$x_{TM} = \frac{1}{0.9996} \ (x_{UTM} - 500,000) \qquad\qquad (7.25)$$
$$y_{TM} = \frac{1}{0.9996} \ (y_{UTM}) \qquad\qquad \text{(Northern hemisphere)}$$
$$y_{TM} = \frac{1}{0.9996} \ (y_{UTM} - 10,000,000) \qquad \text{(Southern hemisphere)}$$
$$k_{TM} = \frac{1}{0.9996} \ k_{UTM}$$
$$\gamma_{TM} = \gamma_{UTM}$$

7.4.3.3 <u>Discontinuity</u>. Notice that UTM coordinates are discontinuous in y at the equator.

7.4.4 <u>Accuracy</u>. The accuracy of the UTM equations is based on the accuracy of the Transverse Mercator equations. Within the area of coverage, the equations for x, y, ϕ, and λ have an error of less than one centimeter. (See Geological Survey Professional Paper 1395, 1987.)

7.4.5 <u>Area of coverage</u>. The area of coverage for UTM coordinates is defined by zone limits, latitude limits, and overlap.

Zone limits:	6° zones, extending 3° to each side of the central meridian. See Table 7.1 for exceptions to this general rule.
Latitude limits:	North: 84° South: 80°
Zone overlap:	40 km on either side of the zone limits.
Polar overlap:	30' toward the poles North: 84° 30' South: 80° 30'

7.5 <u>Lambert Conformal Conic projection</u>. The Lambert conformal conic projection is a conformal projection in which the projected parallels are arcs of concentric circles centered at the pole. The projected meridians are radii of concentric circles that meet at the pole. Moreover, there are one or two parallels, called standard parallels, along which the point scale factor is one. (See Figure 7.4.)
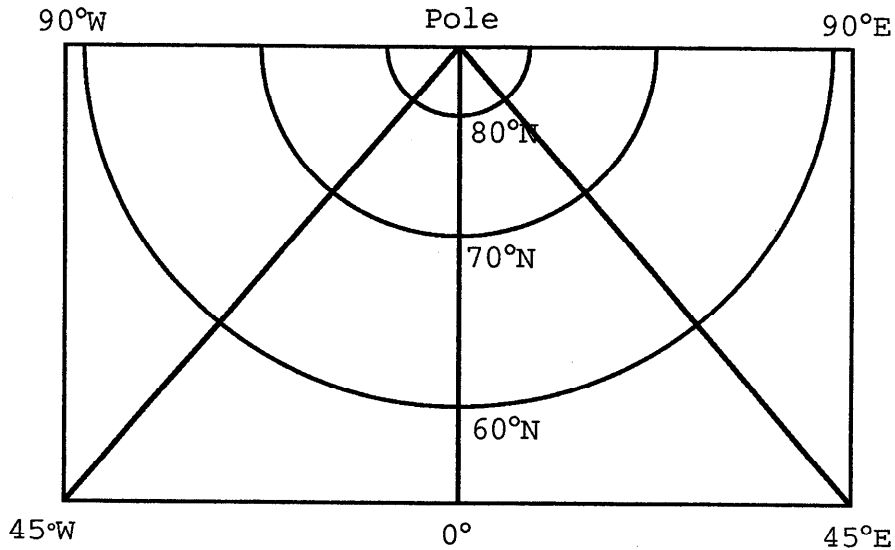
FIGURE 7.4. Meridians and parallels in the Lambert Conformal Conic Projection. (Not drawn to scale.)

7.5.1 Lambert Conformal Conic with two standard parallels.
This presentation of the mapping equations follows Krakiwsky
(1973). The first case to be considered is the case of two
standard parallels, ϕ_1 and ϕ_2. The point scale factor is one
along both standard parallels. Choose a fixed central meridian,
λ_o.

7.5.1.1 Finding (x,y). Begin by fixing a latitude (ϕ_o)
that is below the area of interest. For a given (ϕ, λ) the
mapping equations are

$$x = r \sin L\Lambda$$
$$y = r_o - r \cos L\Lambda$$
(7.26)

where

$$\Lambda = \lambda - \lambda_o$$
$$r = K e^{-Lq}$$
(7.27)
$$r_o = K e^{-Lq_o}$$
$$K = \frac{N_1 \cos \phi_1}{L e^{-Lq_1}} = \frac{N_2 \cos \phi_2}{L e^{-Lq_2}}$$

e = base of natural logarithms (ln e = 1)
ϕ_1 = first standard parallel
ϕ_2 = second standard parallel
$$N = \frac{a}{\sqrt{1 - \varepsilon^2 \sin^2 \phi}}$$

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$$q = \ln \left[\tan \left(\frac{\pi}{4} + \frac{\phi}{2} \right) \left(\frac{1 - \varepsilon \sin \phi}{1 + \varepsilon \sin \phi} \right)^{\frac{\varepsilon}{2}} \right]$$

$$\varepsilon^2 = \frac{a^2 - b^2}{a^2}$$

$$L = \frac{\ln(N_1 \cos \phi_1) - \ln(N_2 \cos \phi_2)}{q_2 - q_1}$$

$q_j = q$ evaluated at ϕ_j , $j = 0,1,2$

$N_j = N$ evaluated at ϕ_j , $j = 1,2$

7.5.1.2 <u>Finding (ϕ, λ) with two standard parallels</u>. For the reverse case, when (x, y) is given along with ϕ_0, ϕ_1, ϕ_2, and λ_o, first compute L, r_o, K, N_1, N_2, q_o, q_1, and q_2 from equation set 7.27. Then compute λ and q from

$$\lambda = \frac{\theta}{L} + \lambda_o \qquad\qquad (7.28)$$

$$q = \frac{\ln \frac{K}{r}}{L}$$

in which

$$r = \frac{r_o - y}{\cos \theta} \qquad\qquad (7.29)$$

$$\theta = \arctan \frac{x}{r_o - y}$$

Note: When evaluating the arc tangent, the user must take care to insure that the resulting angle lies in the desired quadrant.

Once q is found, then ϕ can be determined using the Newton - Raphson method, which is explained next. Define functions f and f', and an initial approximation ϕ' by

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$$f(\phi) = -q + \frac{1}{2} \ln \left[\left(\frac{1 + \sin \phi}{1 - \sin \phi} \right) \left(\frac{1 - \varepsilon \sin \phi}{1 + \varepsilon \sin \phi} \right)^{\varepsilon} \right] \qquad (7.30)$$

$$f'(\phi) = \frac{1 - \varepsilon^2}{(1 - \varepsilon^2 \sin^2 \phi) \cos \phi}$$

$$\phi' = 2 \arctan (e^q) - \pi/2$$

Then successive iterations for ϕ are found from the equation

$$\phi'' = \phi' - \frac{f(\phi')}{f'(\phi')} \qquad (7.31)$$

In particular, substitute ϕ' into this equation to obtain ϕ'', the first approximation to ϕ. Then set $\phi' = \phi''$, substitute ϕ' into the equation to obtain ϕ'', the second approximation to ϕ. Continue this procedure until the differences between successive candidates for ϕ are sufficiently small.

7.5.2 <u>Lambert Conformal Conic with one standard parallel</u>. In the second case, the point scale factor is one along one standard parallel, ϕ_o. Fix a central meridian, λ_o.

7.5.2.1 <u>Finding (x,y)</u>. The mapping equations are

$$x = r \sin L\Lambda \qquad (7.32)$$
$$y = r_o - r \cos L\Lambda$$

where

$$\Lambda = \lambda - \lambda_o$$

$$L = \sin \phi_o \qquad (7.33)$$

$$K = N_o \cot \phi_o \, e^{q_o \sin\phi_o}$$

$$r = K \, e^{-Lq}$$

$$r_o = K \, e^{-Lq_o}$$

e = base of natural logarithms ($\ln e = 1$),

$$N = \frac{a}{\sqrt{1 - \varepsilon^2 \sin^2 \phi}}$$

$$q = \ln \left[\tan \left(\frac{\pi}{4} + \frac{\phi}{2} \right) \left(\frac{1 - \varepsilon \sin \phi}{1 + \varepsilon \sin \phi} \right)^{\frac{\varepsilon}{2}} \right]$$

q_o = q evaluated at ϕ_o

N_o = N evaluated at ϕ_o

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7.5.2.2 <u>Finding (φ,λ) with one standard parallel</u>. To transform in the other direction, suppose that x, y, ϕ_o, and λ_o are given. First, compute L, K, N_o, q_o, r, and r_o from equation set 7.33. Then, as in the previous case, (φ, λ) can be found from

$$\lambda = \frac{\theta}{L} + \lambda_o \qquad (7.34)$$

$$q = \frac{\ln \dfrac{K}{r}}{L}$$

in which

$$r = \frac{r_o - y}{\cos \theta} \qquad (7.35)$$

$$\theta = \arctan \frac{x}{r_o - y}$$

Again, φ can be obtained from q by using the Newton-Raphson method (see equation sets 7.30 and 7.31).

7.5.3 <u>Scale factor and convergence</u>. The scale factor and convergence of the meridian are given by

$$k = \frac{rL}{N \cos \phi} \qquad (7.36)$$

$$\gamma = L\lambda$$

In the case of two standard parallels, the values of r, L, and N are given in equation set 7.27. For one standard parallel, the values of r, L, and N are found in equation set 7.33.

7.5.4 <u>Accuracy</u>. For the Lambert conformal conic projection with either one or two standard parallels, the equations for x, y, λ, q, k, and γ are exact. The equation for φ can be iterated until the desired accuracy is obtained.

7.5.5 <u>Area of coverage</u>. The family of Lambert conformal conic projections and their limiting cases is used worldwide. For a given member of this family, the point scale factor increases as a point moves away from the standard parallel(s).

7.6   Polar Stereographic projection. The Polar Stereographic projection is a limiting case of the Lambert conformal conic projection when the one standard parallel approaches a pole. In this conformal projection meridians are straight lines, parallels are concentric circles, and the point scale factor is one at the pole (see Figure 7.5).
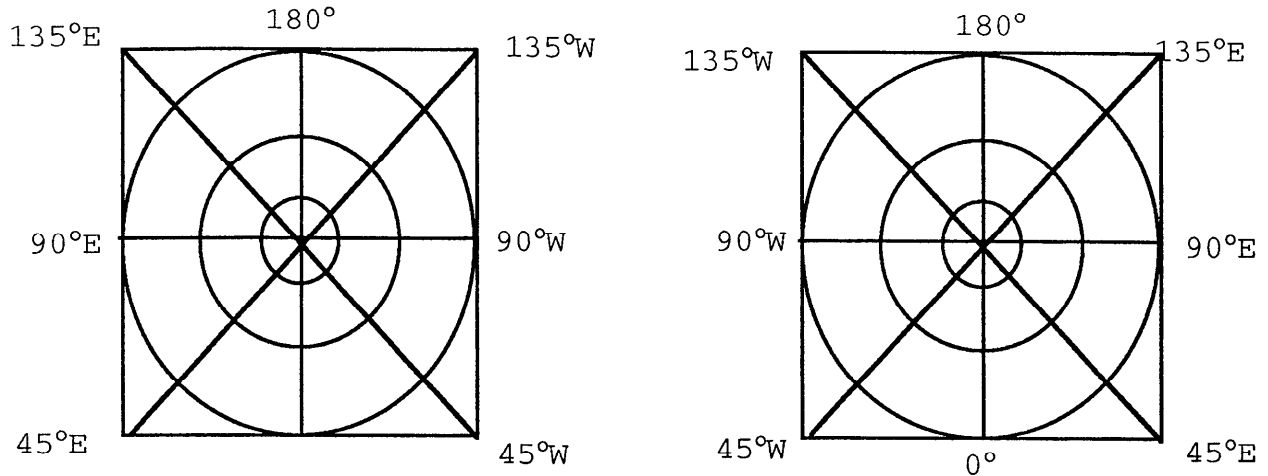
Figure 7.5.   Meridians and parallels in the Universal Polar Stereographic Projection. Left: South Zone, Right: North Zone   (Not drawn to scale.)

7.6.1   Polar Stereographic mapping equations. There is considerable diversity in the treatment that various authors give to the polar Stereographic projection. Some authors interchange the x and y axis; others change the sign convention. The discussion here is largely taken from NSWC/DL TR-3624 (1977) since it most closely conforms to DMA TM 8358.2 (1989).

7.6.1.1   Finding (x,y).   Given (ϕ, λ), the Polar Stereographic coordinates are

$$x = r \sin \lambda \qquad\qquad (7.37)$$
$$y = -r \cos \lambda \qquad \text{(Northern hemisphere)}$$
$$y = r \cos \lambda \qquad \text{(Southern hemisphere)}$$

in which

$$r = K \tan \left(\frac{\pi}{4} - \frac{|\phi|}{2} \right) \left(\frac{1 + \varepsilon \sin |\phi|}{1 - \varepsilon \sin |\phi|} \right)^{\frac{\varepsilon}{2}} \qquad (7.38)$$

$$K = \frac{2a^2}{b} \left(\frac{1 - \varepsilon}{1 + \varepsilon} \right)^{\frac{\varepsilon}{2}}$$

$$\varepsilon^2 = \frac{a^2 - b^2}{a^2}$$

$|\phi|$ denotes the absolute value of ϕ.

(ϕ, λ) is expressed in radians.

At the poles, λ is not defined. Nevertheless, these equations can be used at the poles since r is zero there. At $\phi = \pm \pi/2$, pick an arbitrary value for λ. Then, equation sets 7.37 and 7.38 yield Polar Stereographic coordinates of $(0, 0)$.

7.6.1.2 <u>Finding (ϕ,λ)</u>. The discussion of the conversion from (x, y) to (ϕ, λ) follows Krakiwsky (1973). If $(x, y) = (0, 0)$, then $\phi = \pm 90°$ and λ is undefined. Otherwise, the longitude is given by

$$\lambda = - \arctan \frac{x}{y} \quad \text{(Northern hemisphere)} \quad (7.39)$$

$$\lambda = \arctan \frac{x}{y} \quad \text{(Southern hemisphere)}$$

If $(x, y) \neq (0, 0)$, then x, y or both are non-zero. Use one of these non-zero values to find q from the equations

$$e^{-q} = \frac{|y|}{K \, |\cos \lambda|} \quad \text{when } |x| > |y| \quad (7.40)$$

$$e^{-q} = \frac{x}{K \sin \lambda} \quad \text{when } |y| \geq |x|$$

where

$$K = \frac{2a^2}{b} \left(\frac{1 - \varepsilon}{1 + \varepsilon}\right)^{\frac{\varepsilon}{2}} \quad (7.41)$$

As in the Lambert conformal conic projection, ϕ can be found from q using the Newton-Raphson method. Define functions f and f', and an initial approximation ϕ_o by

$$f(\phi) = -q + \frac{1}{2} \ln \left[\left(\frac{1 + \sin \phi}{1 - \sin \phi}\right)\left(\frac{1 - \varepsilon \sin \phi}{1 + \varepsilon \sin \phi}\right)^{\varepsilon}\right] \quad (7.42)$$

$$f'(\phi) = \frac{1 - \varepsilon^2}{(1- \varepsilon^2 \sin^2 \phi)\cos \phi}$$

$$\phi_o = 2 \arctan e^{q} - \frac{\pi}{2}$$

Then, successive iterations for ϕ are found from the equation

$$\phi_n = \phi_{n-1} - \frac{f(\phi_{n-1})}{f'(\phi_{n-1})} \tag{7.43}$$

In particular, substitute ϕ_0 into this equation to obtain ϕ_1, the first approximation to ϕ. Then substitute ϕ_1 into the equation to obtain ϕ_2, the second approximation to ϕ. Continue this procedure until the differences between successive candidates for ϕ are sufficiently small. Even though q approaches infinity at the poles, this method is satisfactory for many computers. This procedure yields a positive value of ϕ. In the southern hemisphere, replace this positive value of ϕ by $-\phi$.

7.6.1.3 <u>Alternate method for finding ϕ</u>. Alternatively, an approximate equation for ϕ is given in Geological Survey Professional Paper 1395 (1987).

$$\phi = \chi + A\ \sin 2\chi + B\ \sin 4\chi + C\ \sin 6\chi + D\ \sin 8\chi \tag{7.44}$$

where

$$\chi = \frac{\pi}{2} - z \qquad \text{(in radians)} \tag{7.45}$$

$$\tan \frac{z}{2} = \frac{x}{K \sin \lambda} \qquad x \neq 0$$

$$\tan \frac{z}{2} = \left| \frac{y}{K \cos \lambda} \right| \qquad y \neq 0$$

$$K = \frac{2a^2}{b} \left(\frac{1 - \varepsilon}{1 + \varepsilon} \right)^{\frac{\varepsilon}{2}}$$

$$A = \frac{1}{2}\varepsilon^2 + \frac{5}{24}\varepsilon^4 + \frac{1}{12}\varepsilon^6 + \frac{13}{360}\varepsilon^8$$

$$B = \frac{7}{48}\varepsilon^4 + \frac{29}{240}\varepsilon^6 + \frac{811}{11520}\varepsilon^8$$

$$C = \frac{7}{120}\varepsilon^6 + \frac{81}{1120}\varepsilon^8$$

$$D = \frac{4279}{161280}\varepsilon^8$$

$$\varepsilon^2 = \frac{a^2 - b^2}{a^2}$$

Since this procedure yields a positive value of ϕ, replace ϕ by $-\phi$ in the southern hemisphere.

7.6.1.4 <u>Finding the point scale factor</u>. The point scale factor is given by

$$k = 1 \quad \text{at the poles} \tag{7.46}$$

$$k = \frac{r}{N \cos \phi} \quad \text{elsewhere}$$

where

$$r = K \tan \left(\frac{\pi}{4} - \frac{|\phi|}{2} \right) \left(\frac{1 + \varepsilon \sin |\phi|}{1 - \varepsilon \sin |\phi|} \right)^{\frac{\varepsilon}{2}} \tag{7.47}$$

$$K = \frac{2a^2}{b} \left(\frac{1 - \varepsilon}{1 + \varepsilon} \right)^{\frac{\varepsilon}{2}}$$

$$N = \frac{a}{\sqrt{1 - \varepsilon^2 \sin^2 \phi}}$$

7.6.1.5 <u>Finding the convergence of the meridian</u>. The convergence of the meridian is

$$\gamma = \lambda \qquad \text{(Northern hemisphere)} \tag{7.48}$$

$$\gamma = -\lambda \qquad \text{(Southern hemisphere)}$$

7.6.2 <u>Accuracy</u>. For the Polar Stereographic projection, the equations for x, y, λ, and q are exact. The equation for ϕ can be iterated until the desired accuracy is obtained.

7.6.3 <u>Area of coverage</u>. There is no general agreement on the area of coverage for the Polar Stereographic projection. Section 7.7.5 gives the area of coverage for UPS coordinates.

7.7 <u>Universal Polar Stereographic (UPS) projection</u>. The Universal Polar Stereographic (UPS) projection is the standard military grid used in polar regions (see Figure 7.6). The main reference for the UPS grid is DMA TM 8358.2 (1989). The UPS grid is a family of two projections that differ from the Polar Stereographic projection in these ways:

a. Both the x and y values (see Figure 7.6), called easting and northing, respectively, have an origin of 2,000,000 meters.
b. The scale factor at the origin is 0.994.
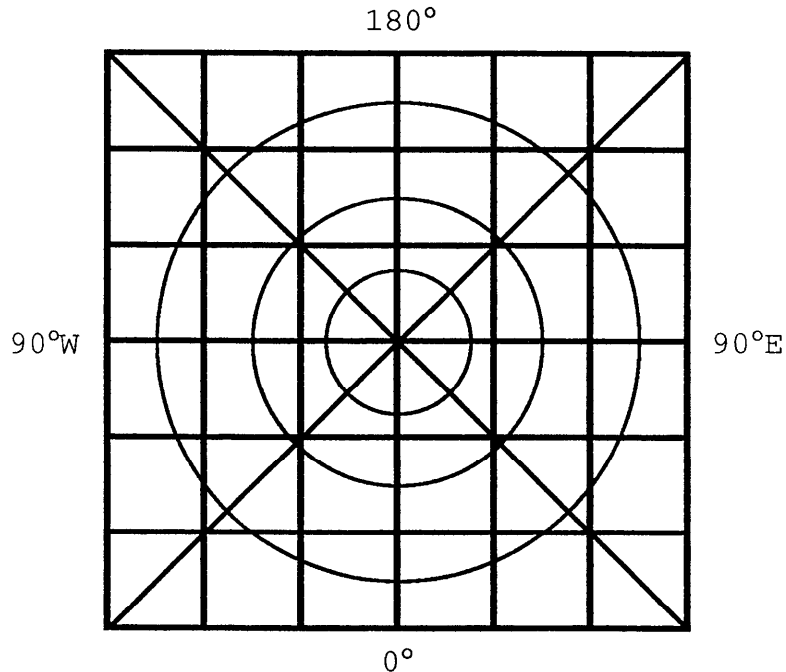
Figure 7.6.   Meridians and parallels imposed on a Universal
            Polar Stereographic grid (north zone).

The limits of the system are north of 84° N and south of 80° S. In
order to provide a 30-minute overlap with the UTM grid, the UPS
grid contains an overlap that extends to 83° 30'N and 79° 30'S.
Although the UPS grid originally was defined to use the
International Ellipsoid, in practice several reference ellipsoids
have been used.   The ellipsoid appropriate to the selected datum
should be used.

    7.7.1   <u>Universal Polar Stereographic (UPS) mapping</u>
<u>equations</u>.   For known geodetic coordinates, the UPS easting is
given by

$$x_{UPS} = 2,000,000 + 0.994r \sin \lambda \tag{7.49}$$

and the UPS northing is given by

$$y_{UPS} = 2,000,000 - 0.994r \cos \lambda \quad \text{(Northern hemisphere)} \tag{7.50}$$
$$y_{UPS} = 2,000,000 + 0.994r \cos \lambda \quad \text{(Southern hemisphere)}$$

The parameter r is computed using equation set 7.38.

    7.7.2   <u>Finding (ϕ, λ)</u>.   If a hemisphere and UPS coordinates
(x, y) are known, convert the UPS values to Polar Stereographic
coordinates using equation set 7.51, and then transform to
geodetic coordinates as described in Sections 7.6.1.2 and 7.6.1.3.

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$$x_{PS} = \frac{1}{0.994} (x_{UPS} - 2,000,000) \qquad (7.51)$$

$$y_{PS} = \frac{1}{0.994} (y_{UPS} - 2,000,000)$$

In particular, at the poles where $x = 2,000,000$ and $y = 2,000,000$ then, $\phi = \pm 90°$ and λ is undefined. Otherwise longitude and latitude can be found from

$$\lambda = - \arctan \frac{(x_{UPS} - 2,000,000)}{(y_{UPS} - 2,000,000)} \quad \text{(Northern hemisphere) (7.52)}$$

$$\lambda = + \arctan \frac{(x_{UPS} - 2,000,000)}{(y_{UPS} - 2,000,000)} \quad \text{(Southern hemisphere)}$$

$$\phi = \chi + A \sin2\chi + B \sin4\chi + C \sin6\chi + D \sin8\chi$$

in which

$$\chi = \frac{\pi}{2} - z \qquad (7.53)$$

$$\tan \frac{z}{2} = \frac{(x_{UPS} - 2,000,000)}{0.994K \sin \lambda} \qquad x_{UPS} \neq 2,000,000$$

$$\tan \frac{z}{2} = - \frac{(y_{UPS} - 2,000,000)}{0.994K \cos \lambda}$$

$$y_{UPS} \neq 2,000,000 \quad \text{(Northern hemisphere)}$$

$$\tan \frac{z}{2} = + \frac{(y_{UPS} - 2,000,000)}{0.994K \cos \lambda}$$

$$y_{UPS} \neq 2,000,000 \quad \text{(Southern hemisphere)}$$

$$K = \frac{2a^2}{b} \left(\frac{1 - \epsilon}{1 + \epsilon} \right)^{\frac{\epsilon}{2}}$$

$$\epsilon^2 = \frac{a^2 - b^2}{a^2}$$

The values of A, B, C, and D are found using equation set 7.45.

In the southern hemisphere, replace the computed value of ϕ with $-\phi$.

7.7.3 <u>Finding point scale factor and convergence of the meridian</u>. The point scale factor and convergence of the meridian are

$$k = 0.994 \text{ at the poles} \qquad (7.54)$$

$$k = 0.994 \frac{r}{N \cos \phi} \text{ elsewhere}$$

$$\gamma = \lambda \qquad \text{(Northern hemisphere)}$$
$$\gamma = -\lambda \qquad \text{(Southern hemisphere)}$$

where r and N are given in equation set 7.47.

7.7.4 <u>Accuracy</u>. The accuracy of the UPS equations is based upon the accuracy of their Polar Stereographic counterparts. For the Polar Stereographic projection, the equations for x, y, λ, and q are exact. The equation for ϕ can be iterated until the desired accuracy is obtained.

7.7.5 <u>Area of coverage</u>. The area of coverage for UPS coordinates is defined by zone limits and overlap.

North zone: The north polar area 84° - 90°

South zone: The south polar area 80° - 90°

UTM overlap: 30' overlap,
                    North: 83° 30'
                    South: 79° 30'

7.8 <u>The U.S. Military Grid Reference System (MGRS)</u>. The MGRS is an alpha-numeric system, based upon the UTM and UPS map projections, for identifying positions. An MGRS coordinate consists of a zone designation, alphabetic 100,000-meter grid square designator, and numeric coordinates within the 100,000 meter grid square. MGRS coordinates are defined for the primary UTM and UPS areas, but not for the overlap areas.

7.8.1 <u>MGRS coordinates in the UTM area</u>. In the UTM area, MGRS coordinates are based on the ellipsoid, geodetic latitude, UTM zone, easting and northing as follows:

a. The first two characters of the MGRS coordinate are the two digits of the numeric UTM zone (see sections 7.4.1.1 and 7.4.1.2.) Leading zeros must be included.

b. The third character of the MGRS coordinate is a letter representing a band of geodetic latitude. Beginning at 80° South and proceeding northward, 20 bands are lettered C through X,

omitting I and O. The bands are all 8° high except band X (72°N to 84°N), which is 12° high.

c. The fourth and fifth characters of the MGRS coordinate are a pair of letters representing the 100,000-meter grid square. The letter pair can be selected from figure 7.7 and 7.8 (from DMA TM 8358.1.) First, reduce the UTM northing by multiples of 2,000,000 until the resulting value is in the range of 0 to 1,999,999 meters. Second, reduce the UTM northing and easting values to the nearest 100,000 meters. Third, locate the corresponding 100,000 meter northing and easting grid lines for the UTM zone number on the figure. The grid square identifier will be immediately above and to the right of the intersection of the easting and northing grid lines on the figure.

Note: The alphabetic method of designating 100,000-meter grid squares has changed over time. Older products may have different grid square designations than those described above. Also, some software was programmed to compute MGRS coordinates based on geographic position and a map of preferred ellipsoids. The boundaries between preferred ellipsoids have changed with time. Therefore, MGRS coordinates computed with such software may not agree with current map products. MGRS computation software should always require input of the ellipsoid. If MGRS coordinates are to be used in conjuction with a map product they should be checked against the map product to verify compatibility.

d. The remainder of the MGRS coordinate consists of the numeric easting and northing values within the 100,000-meter grid square. The left half of the digits is the easting grid value, which is read to the right of the 100,000-meter easting grid line established in step c. above. The right half of the digits is the northing grid value, which is read northward of the 100,000-meter northing grid line established in step c. above. Both the easting and northing grid values are within the range of 0 to 100,000 meters (i.e., zone and grid square designators, only) to 1 meter (i.e., zone and grid designators followed by 10 digits, five for easting grid value and five for northing grid value). Both easting and northing value must have the same resolution and must include leading zeros.

7.8.2 <u>MGRS coordinates in the UPS area</u>. In the UPS area, MGRS coordinates are based on the ellipsoid, geodetic latitude and longitude, easting, and northing as follows:

a. The first character of the MGRS coordinate is
A in the Southern and Western Hemispheres,
B in the Southern and Eastern Hemispheres,
Y in the Northern and Western Hemispheres, and
Z in the Northern and Eastern Hemispheres as shown in figures 7.9 and 7.10.

FIGURE 7.7

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ZONES

SET 1
1 7 13 19 25 31
37 43 49 55

SET 2
2 8 14 20 26 32
38 44 50 56

SET 3
3 9 15 21 27 33
39 45 51 57

SET 4
4 10 16 22 28 34
40 46 52 58

SET 5
5 11 17 23 29 35
41 47 53 59

SET 6
6 12 18 24 30 36
42 48 54 60

2,000,000 m
1,500,000 m
1,000,000 m
500,000 m
0

200,000 m 300,000 400,000 500,000 600,000 700,000 800,000 E

USE THIS DIAGRAM FOR:

GEODETIC REFERENCE SYSTEM 1980 ELLIPSOID
INTERNATIONAL ELLIPSOID
WORLD GEODETIC SYSTEM 1984 ELLIPSOID
WORLD GEODETIC SYSTEM 1972 ELLIPSOID
AUSTRALIAN NATIONAL ELLIPSOID
CLARKE 1866 ELLIPSOID (UTM ZONES 47-50 ONLY)
EVEREST ELLIPSOID
SOUTH AMERICAN 1969 ELLIPSOID (GRS 67)

GRID ZONE DESIGNATIONS
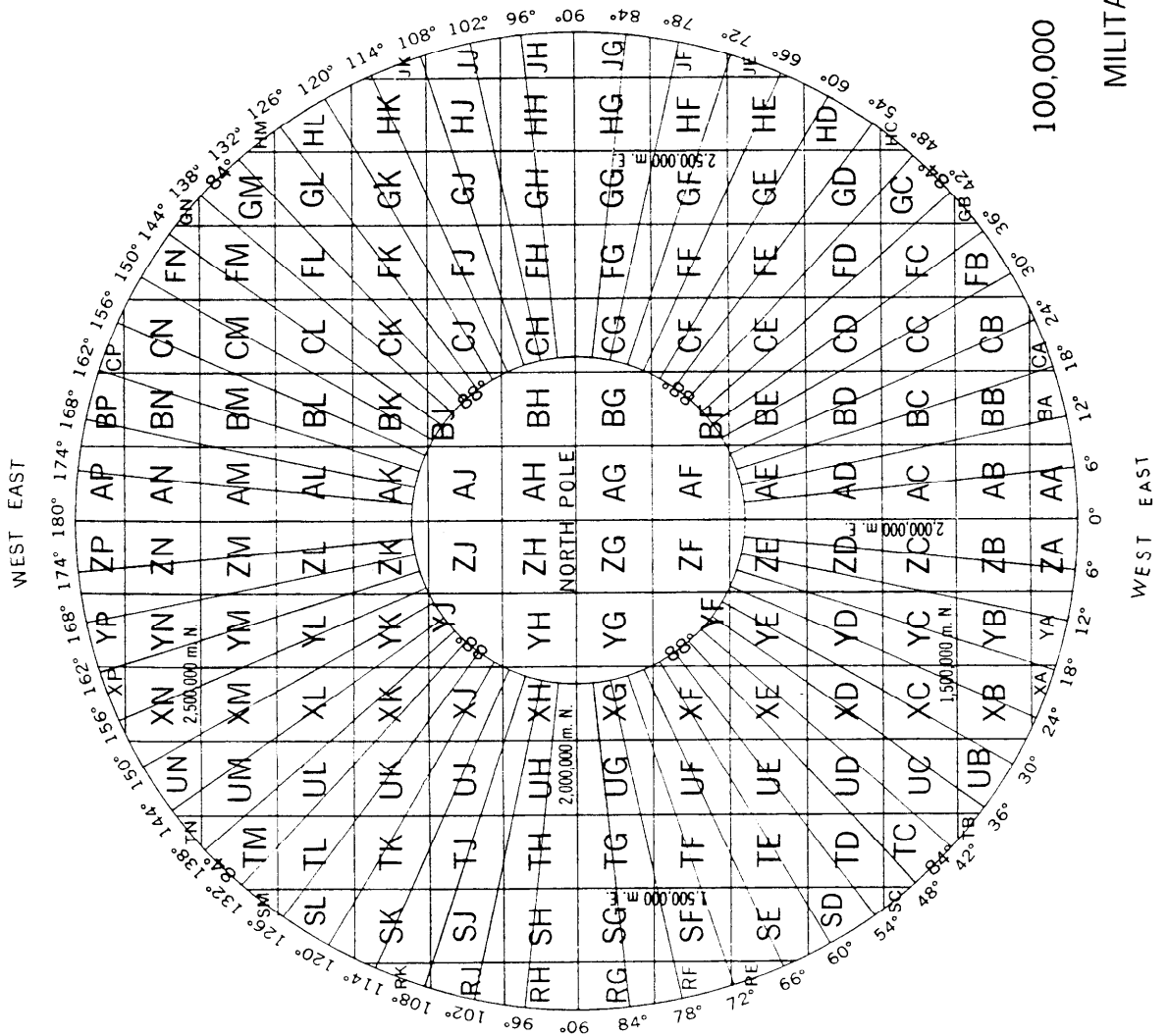
FIGURE 7.9

100,000 METER SQUARE IDENTIFICATIONS
FOR THE
MILITARY GRID REFERENCE SYSTEM

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GRID ZONE DESIGNATIONS
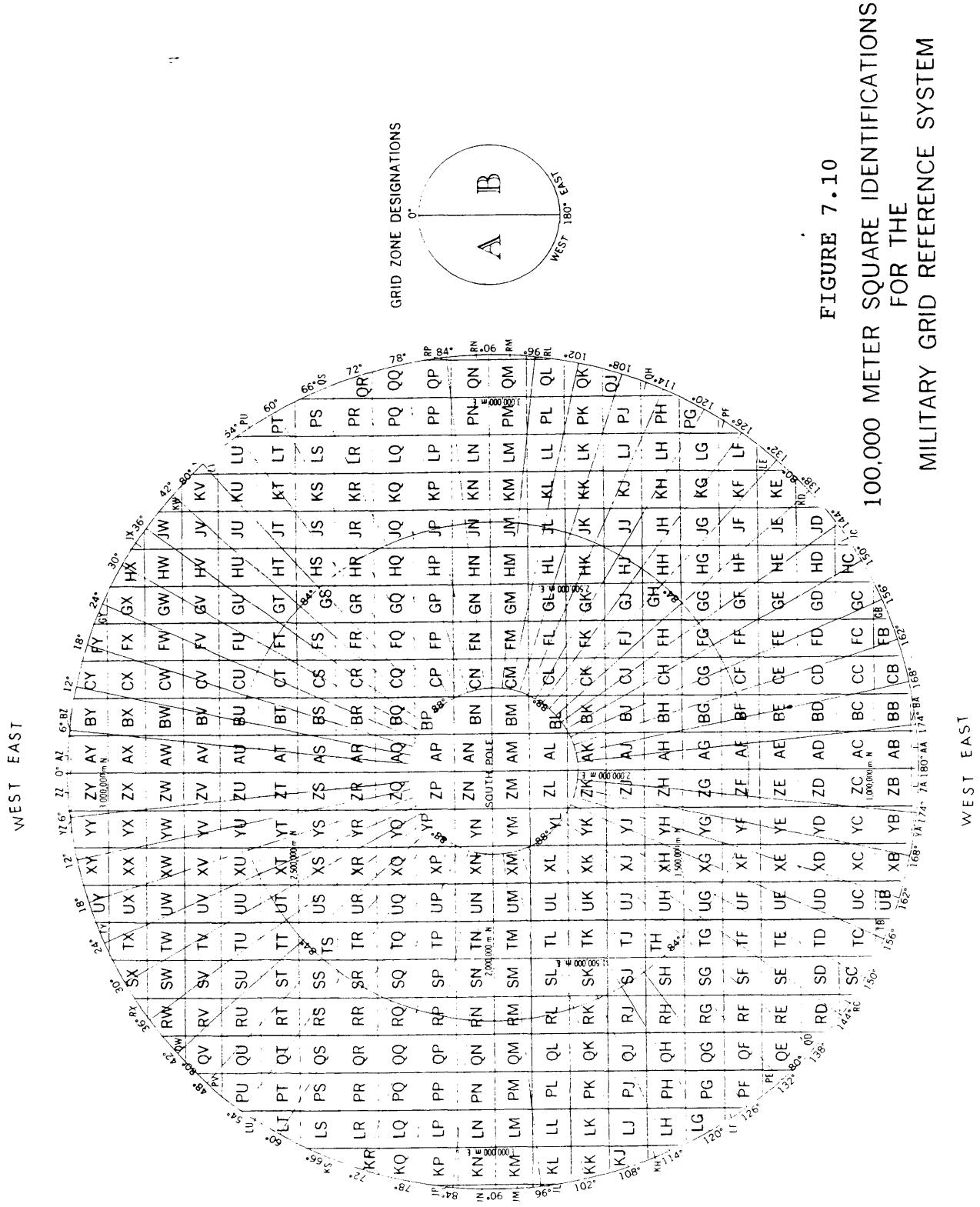
FIGURE 7.10

100,000 METER SQUARE IDENTIFICATIONS
FOR THE
MILITARY GRID REFERENCE SYSTEM

b.   The second and third characters of the MGRS coordinate
are letters representing a 100,000 meter grid square. The letter
pair can be selected from figure 7.9 or 7.10 (from DMA TM 8358.1.)
Reduce the UPS easting and northing values to the nearest 100,000
meters. Locate the corresponding 100,000 meter northing and
easting grid lines for the UTM zone number on the figure. The
grid square identifier will be immediatly above and to the right
of the intersection of the easting and northing grid lines on the
figure.

c.   The method for calculating the remainder of the MGRS
coordinate is the same as in paragraph 7.8.1 d.

7.9   The World Geographic Reference System (GEOREF). The
GEOREF is an alpha-numeric system for reporting positions for air
defense and strategic air operations. It is based upon geodetic
coordinates. GEOREF is described in DMA TM 8358.1 (1990).

7.10   Non-standard grids. The specifications and diagrams of
non-standard grids, used as the primary or secondary grid on DMA
produced maps, are contained in DMA TM 8358.1 (1990).

7.11   Modifications to map projections. Providing equations
for transforming between geodetic coordinates and every local map
projection is beyond the scope of this Technical Report. However,
many local map projections are minor modifications of the map
projections given in this Technical Report. Typically, these
modifications include use of false Eastings (E) and Northings (N)
at the projection origin, a non-unity scale factor (k_o) along the
principal line of the projection, and units other than meters.
Sometimes the local projection origin is offset from the origin
used in this Technical Report. Chapter four of DMA TM 8352.1
(1991) gives parameters for several local projections.

If the local projection is based on one of the projections
given in this Technical Report convert geodetic latitude (ϕ) and
longitude (λ) to coordinates in the local projection by using the
following projection;

1.   Make sure (ϕ, λ) are on the ellipsoid for the local
     projection and determine the ellipsoid's parameters, a
     and either f or b.

2.   Convert (ϕ, λ) to map projection coordinates (X_p, Y_p)
     using the appropriate map projection equations given in
     this Technical Report. Note that the units for X_p and Y_p
     are in meters.

3.   If the local projection has false coordinates, E, N, at
     the origin, convert them to units of meters if necessary.

4. The map projection coordinates (X_p, Y_p) and scale factor (k_p) can be converted to local projection coordinates using equations 7.55 if the origin of the local projection is the same as the origin of that projection as used in this technical report. Those origins are central meridian (λ_o) and equator for Transverse Mercator and Mercator projections, and standard parallel(s) and equator for the Lambert Conformal projection.

$$X = k_o X_p + E$$

$$Y = k_o Y_p + N \tag{7.55}$$

$$k = k_o k_p$$

$$\gamma = \gamma_p$$

where

E = false X-value (in meters) at the origin

N = false Y-value (in meters) at the origin

ko = scale factor along the appropriate line.
(Along the equator for the Mercator projection, along the standard parallel(s) for the Lambert Conformal Conic projection, and along the central meridian for the Transverse Mercator projection.)

5. Convert (X, Y) to the units that are locally used. It is the user's responsibility to determine if this procedure is valid. This can be done by comparing computed mapping coordinates to test points where the relationship between (ϕ, λ) and (X, Y) is known, independently.

The above procedure is intended to be used only when the origin of the local map projection coincides with the origin used in this Technical Report. Some Transverse Mercator projections have origins of (ϕ_o, λ_o) instead of (0, λ_o) on the equator. For some of these projections, agreement with test points can be obtained by changing the computation of Y in equation set 7.55 to

$$Y = k_o Y_{TM} + N - k_o S\phi_o \qquad\qquad (7.56)$$

where

k_o = scale factor on the central meridian, λ_o

Y_{TM} = Transverse Mercator Y value computed from equation set 7.8

N = false Northing in meters of the origin

$S\phi_o$ = value computed using equation set 7.9
(arc length in meters on the ellipsoid from the equator to the latitude ϕ_o.)

8. SELECTION OF TECHNIQUES, SOFTWARE DEVELOPMENT, AND TESTING

8.1 The general conversion process. The general coordinate conversion process can be divided into three basic parts:

1. Conversion between map projection coordinates and geodetic coordinates (see Figure 8.1).

2. Coordinate conversion between geodetic coordinates and Cartesian coordinates within a datum (see Figure 8.2).

3. Coordinate transformation between datums (datum shifts) (see Figure 8.3).

The general conversion process which includes all methods presented in this Technical Report, is described graphically in Figure 8.4.

a. Figure 8.1 represents the map projection coordinate (x,y) to the geodetic coordinate (ϕ,λ) conversion described in Section 7. As noted in Section 7.1, map projection coordinates do not include the height information necessary for finding the geodetic height (H) or elevation (h). For conversions from map projection to geodetic coordinates, the user may have to provide this information. Elevations can often be inferred from contour lines on the map or they can be obtained from DTED or direct measurement. Geodetic height may be measured by using GPS or by calculating from an elevation and geoid separation, or for some applications, may be assumed to be 0 (zero).

$$x,y \quad \overset{\text{A}}{\longleftrightarrow} \quad \phi, \lambda$$
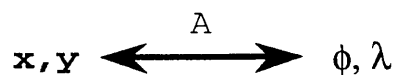
Figure 8.1. Map projection to geodetic coordinate conversion. Height information for H and h is not included in the conversion.

where

x,y = map projection coordinates
ϕ = geodetic latitude
λ = geodetic longitude
A = map projection coordinates to/from latitude, longitude conversion

b. Figure 8.2 represents coordinate conversions within a datum as described in Section 4.3.4 and 6.5.

$$X, Y, Z$$

$$C \updownarrow$$

$$\Phi, \lambda, h$$

$$B \updownarrow$$

$$\Phi, \lambda, H$$
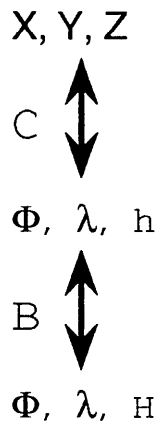
FIGURE 8.2. Coordinate conversions within a datum,

where

$$X, Y, Z = \text{Cartesian coordinates}$$
$$\phi, \lambda, h = \text{geodetic coordinates (latitude, longitude, geodetic height)}$$
$$H = \text{elevation} \cdot$$
$$B = \text{geodetic height to/from elevation conversion}$$
$$C = \text{geodetic coordinate to/from Cartesian coordinate conversion}$$

Note that a conversion between Cartesian coordinates (X,Y,Z) and local datum coordinates (ϕ,λ,H) is not a direct conversion and requires going through the geodetic coordinates (ϕ,λ,h).

c. Figure 8.3 represents the datum shift portion of the general conversion process as described in Section 6. Two attributes of figure 8.3 are important to note.

   1. Direct local datum to local datum shifts are not performed.

   2. Elevations do not change when changing datums. Exceptions to the second statement are explained in Section 6.7.

d. In general, a direct local datum to local datum shift should not be done. Instead, the local datum to local datum shift should be done in two steps.

   1. An (old) local datum to WGS 84 shift.

   2. A WGS 84 to (new) local datum shift.

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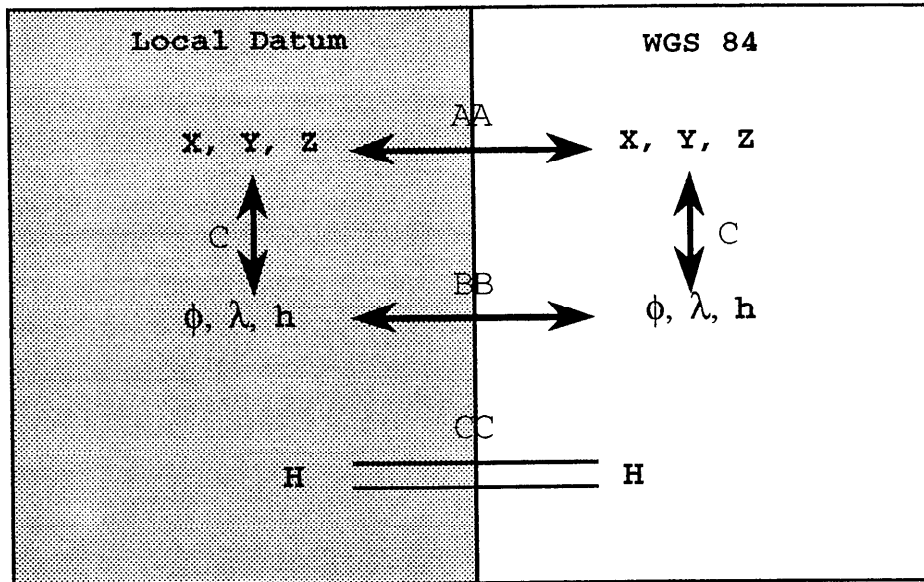
Figure 8.3. Datum shifts.
Note elevation values (H) do not generally change when
    changing datums,

where

    X, Y, Z = Cartesian coordinates
    ϕ, λ, h = geodetic coordinates (latitude, longitude,
              geodetic height)
        H = elevation
        C = geodetic coordinate to/from Cartesian coordinate
            conversion
       AA = Cartesian coordinate-based datum shifts
       BB = geodetic coordinate-based datum shifts
       CC = elevations are unchanged

e. The general conversion process, which includes all
    methods presented in this Technical Report, is described
    graphically in figure 8.4. Figure 8.4 is the combination
    of figures 8.1 through 8.3,
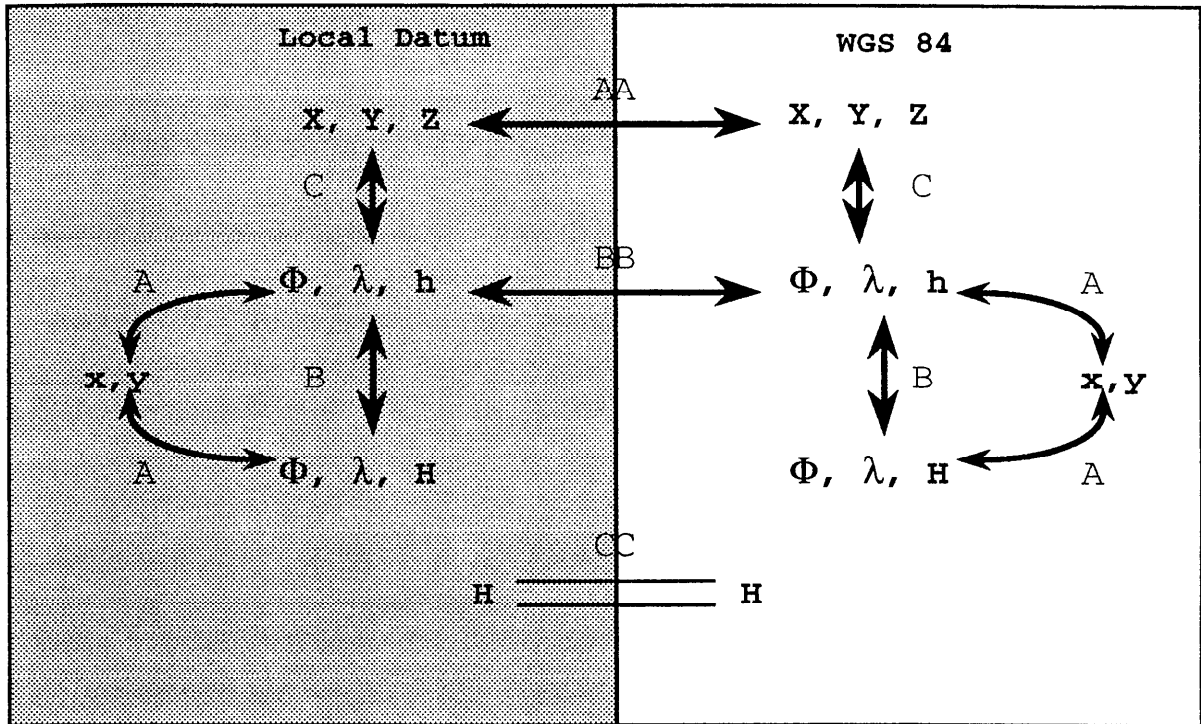
Figure 8.4. The general conversion process,

where

           x, y = map projection coordinates
       X, Y, Z = Cartesian coordinates
     ϕ, λ, h = geodetic coordinates (latitude, longitude,
                  geodetic height)
              H = elevation
              A = map projection coordinates to/from latitude,
                  longitude conversion
              B = geodetic height to/from elevation conversion
              C = geodetic coordinate to/from Cartesian coordinate
                  conversion
            AA = Cartesian coordinate-based datum shifts
            BB = geodetic coordinate-based datum shifts
            CC = elevations are unchanged

f. In figures 8.1 through 8.4, coordinate conversions are
    denoted as a single letter (A, B, C). They are described
    in Section 7, Section 6.5 and Section 4.3.4,
    respectively. Datum shifts are denoted as double letters
    (AA, BB, CC) and are described in Section 6.

g. The Cartesian coordinate-based datum shifts (AA) include:

Seven-parameter geometric transformation (6.2)
$\Delta X, \Delta Y, \Delta Z$ direct application   (6.3.1 & 6.3.2)

Both of these datum shift methods are geometric transformations.

h. The geodetic coordinate based datum shifts (BB) include:

Molodensky shifts   (6.3.3)
$\Delta X, \Delta Y, \Delta Z$ three-step method   (6.3.1.1 & 6.3.2.1)
Direct WGS 72 to WGS 84 transformation (6.4.2)

i. For the purposes of this Technical Report, the elevation to elevation shift (CC) is the equality function (=); that is, the elevation for the same location on different datums is considered to be the same (H = H). An exception to this rule is the WGS 72-derived elevations to WGS 84 elevations shift. Section 6.7 describes the elevation to elevation shift process.

8.1.1 <u>Application of the general conversion process.</u> The general conversion process described above can be used to solve coordinate conversion/datum shift problems in five steps:

a. Locate beginning point and ending point of the problem on figure 8.4.

b. Determine valid paths between beginning and end points.

c. Evaluate alternatives and select a path.

d. Locate the needed equations and parameters in this Technical Report.

e. Implement the selected path.

8.1.2 <u>Procedural examples.</u>

8.1.2.1 <u>Example 8.1.</u> Horizontal geodetic coordinate-based datum shifts $((\phi, \lambda, h)_{LGS}$ to $(\phi, \lambda)_{WGS\ 84})$ (Local datum geodetic heights (h_{LGS}) typically are not available. It is not recommended they be transformed to WGS 84).

This Technical Report has discussed two methods of doing datum shifts. These methods are the three-step method, and Molodensky shifts. Either the three-parameter $\Delta X, \Delta Y, \Delta Z$ shift or the seven-parameter shift may be used in the three-step method. This example corresponds to a "BB" shift of figures 8.3 and 8.4.

a. <u>Approach 1 - Three-step method using a seven-parameter shift.</u>

   1. Compute $(X,Y,Z)_{LGS}$ from $(\phi,\lambda,h)_{LGS}$ using equation sets 4.4 and 4.5, and the reference ellipsoid parameters from Table A.1.

   2. Compute $(X,Y,Z)_{WGS\ 84}$ from $(X,Y,Z)_{LGS}$ using the seven-parameter shift equation set 6.1 and shift parameters from Table E.1.

   3. Compute $(\phi,\lambda)_{WGS\ 84}$ from $(X,Y,Z)_{WGS\ 84}$ using equation sets 4.6 through 4.10.

b. <u>Approach 2 - Three-step method using ΔX, ΔY, ΔZ shifts</u>

   1. Compute $(X,Y,Z)_{LGS}$ from $(\phi,\lambda,h)_{LGS}$ using equation sets 4.4 and 4.5. Reference ellipsoid parameters are found in Table A.1.

   2. Compute $(X,Y,Z)_{WGS\ 84}$ from $(X,Y,Z)_{LGS}$ using equation set 6.3, and shift parameters from Table B.1

   3. Compute $(\phi,\lambda)_{WGS\ 84}$ from $(X,Y,Z)_{WGS\ 84}$ and equation sets 4.6 through 4.10.

c. <u>Approach 3 - Using Molodensky shift</u>

   1. Compute $\Delta\phi$ and $\Delta\lambda$ from (ϕ,λ,h_{LGS}) using the Molodensky equation set 6.5, shift parameters from Table B.1, and the local datum's reference ellipsoid parameters given in Table A.1.

   2. Compute $(\phi,\lambda)_{WGS\ 84}$ from $(\phi,\lambda)_{LGS}$ using equation set 6.6.

8.1.2.2 <u>Example 8.2.</u> WGS 84 Cartesian coordinate to local latitude, longitude, elevation.

Several paths for converting $(X,Y,Z)_{WGS\ 84}$ to $(\phi_{LGS},\lambda_{LGS},H)$ can be traced in figure 8.4. The steps of three approaches are provided below.

a. <u>Approach 1 - Using a seven-parameter datum shift</u>

   1. Compute $(X,Y,Z)_{LGS}$ from $(X,Y,Z)_{WGS\ 84}$ using the seven-parameter shift equation set 6.2 and parameters from Table E.1.

2. Compute $(\phi, \lambda)_{LGS}$ from $(X, Y, Z)_{LGS}$ using equation sets 4.6 through 4.10 with the parameters of the local ellipsoid.

3. Compute $h_{WGS\ 84}$ from $(X, Y, Z)_{WGS\ 84}$ using equations 4.7 through 4.11 (Use equation 4.12 in polar regions).

4. Compute H from $h_{WGS\ 84}$ and $N_{WGS\ 84}$ using equation 5.1, resulting in ϕ_{LGS}, λ_{LGS}, H.

b. <u>Approach 2 - Using a direct ΔX, ΔY, ΔZ datum shift</u>

1. Compute $(X, Y, Z)_{LGS}$ from $(X, Y, Z)_{WGS\ 84}$ using the ΔX, ΔY, ΔZ shift equation set 6.4 and shift parameters from Table B.1.

2. Compute $(\phi, \lambda)_{LGS}$ from $(X, Y, Z)_{LGS}$ using equation sets 4.6 through 4.11 with the parameters of the local ellipsoid.

3. Compute $h_{WGS\ 84}$ from $(X, Y, Z)_{WGS\ 84}$ using equations 4.7 through 4.11. (Use equation 4.12 in polar regions).

4. Compute H from $h_{WGS\ 84}$ and $N_{WGS\ 84}$ using equation 5.1, resulting in ϕ_{LGS}, λ_{LGS}, H.

c. <u>Approach 3 - Using a Molodensky datum shift</u>

1. Compute $(\phi, \lambda, h)_{WGS\ 84}$ from $(X, Y, Z)_{WGS\ 84}$ using equation sets 4.6 through 4.11. (In polar regions, equation set 4.11 should be replaced by equation set 4.12).

2. Compute $(\phi, \lambda)_{LGS}$ from $(\phi, \lambda)_{WGS\ 84}$ using the Molodensky equation set 6.5, equation set 6.7 and shift parameters from Table B.1.

3. Compute H from $h_{WGS\ 84}$ and $N_{WGS\ 84}$ using equation 5.1, resulting in ϕ_{LGS}, λ_{LGS}, H.

8.1.2.3 <u>Example 8.3.</u> Local latitude, longitude, elevation $(\phi_{LGS}$, λ_{LGS}, H) to WGS 84 latitude, longitude, elevation $(\phi_{WGS\ 84}$, $\lambda_{WGS\ 84}$, H).

Example 8.3 is a extension of example 8.1. Example 8.3 shows the recommended methods for determining geodetic height used in each of the geometric datum shift methods. The steps of three alternate approaches are provided as follows.

a. <u>Approach 1 - Using a seven-parameter datum shift</u>

   1. Compute $(X,Y,Z)_{LGS}$ from $(\phi,\lambda,H)_{LGS}$ using equation sets 4.4 and 4.5.

   2. Compute $(X,Y,Z)_{WGS\ 84}$ from $(X,Y,Z)_{LGS}$ using the seven parameter shift equation set 6.1 and shift parameters from Table E.1.

   3. Compute $(\phi,\lambda)_{WGS\ 84}$ from $(X,Y,Z)_{WGS\ 84}$ using equation sets 4.6 through 4.10.

   4. Take H from the input coordinates resulting in $\phi_{WGS\ 84}$, $\lambda_{WGS\ 84}$, H.

b. <u>Approach 2 - Using a ΔX, ΔY, ΔZ datum shift</u>

   1. Compute $(X,Y,Z)_{LGS}$ from $(\phi,\lambda,H)_{LGS}$ using equation sets 4.4 and 4.5.

   2. Compute $(X,Y,Z)_{WGS\ 84}$ from $(X,Y,Z)_{LGS}$ using ΔX, ΔY, ΔZ shift equation set 6.3, and shift parameters from Table B.1.   .

   3. Compute $(\phi,\lambda)_{WGS\ 84}$ from $(X,Y,Z)_{WGS\ 84}$ using equation sets 4.6 through 4.10.

   4. Take H from the input coordinate resulting in $\phi_{WGS\ 84}$, $\lambda_{WGS\ 84}$, H.

c. <u>Approach 3 - Using a Molodensky datum shift</u>

   1. Compute $(\phi,\lambda)_{WGS\ 84}$ from $(\phi,\lambda,H)_{LGS}$ using Molodensky equation set 6.5, equation set 6.6, and shift parameters from Table B.1.

   2. Take H from the input coordinate resulting in $\phi_{WGS\ 84}$, $\lambda_{WGS\ 84}$, H.

8.2 <u>Choosing a datum shift method</u>. Several factors should be considered in selecting a datum shift method. These factors include: shift parameter availability, timing requirements, accuracy requirements, software availability, hardware limitations, standardization considerations and potential future requirements. If future requirements are considered as part of the system development, future product improvements may be minimized. These requirements may include the need to provide data to other systems and to meet higher accuracy requirements. The discussion below applies to a wide range of Army applications.

Special, high accuracy applications are outside the scope of this Technical Report (see Section 1.2).

The three-parameter shift is recommended for mapping and charting and most other military application. For Western Europe (EUR-M), and Ordnance Survey of Great Britain 1936 (OGB-M), STANAG 2211 specifies the seven-parameter shifts. The use of MRE's is not recommended, although it should be noted that STANAG 2211 specifies the use of MRE's for geodetic and geophysical applications where "data is available in geodetic coordinates."

Parameter sets for ΔX, ΔY, ΔZ shifts are found in Appendix B, and seven-parameter geometric shifts are contained in Appendix E. Be sure to use the same method throughout a project, identifying the method chosen in the project documentation so that future users of the data can replicate or backtrace your results.

      8.2.1 <u>Parameter availability</u>. For many datums, only mean-value three-parameter ΔX, ΔY, and ΔZ shifts are given in this Technical Report. Based on system implementation tradeoffs, the system developer can: apply the shifts directly to Cartesian coordinates (see Sections 6.3.1 and 6.3.2); use the three-step method (see Section 6.3.1.1 and 6.3.2.1); or apply shifts to geodetic coordinates using the Molodensky approximations (see Section 6.3.3). For other datums, some combination of seven-parameter geometric shifts, and mean and regional ΔX, ΔY, and ΔZ shifts are available. Select a parameter set from those that meet or come closest to meeting the accuracy requirements of the output product. Base the final selection on standardization requirements and system implementation tradeoffs.

      8.2.2 <u>Accuracy</u>. Errors in the shift parameters are the dominant error source in datum shifting. Errors introduced by any of the datum shift algorithms given in this Technical Report are small compared to the errors in the shift parameters and can generally be neglected. However, take care that accuracy is not degraded by the software and hardware implementation of these algorithms.

        a.   For the large datums, the regional ΔX, ΔY, ΔZ shifts and seven-parameter shifts are generally considered the most accurate, and mean value ΔX, ΔY, ΔZ shifts have the poorest accuracy. The estimated accuracies of the shift parameters are given in Appendices B and E for each listed datum. The errors in geodetic coordinates may be approximated from the ΔX, ΔY, ΔZ errors listed in Appendix B, using the equation set 6.8.

        b.   The error estimates and general guidance above must be used judiciously. For a given datum, all types of shift

parameters were derived from essentially the same set of control data, and errors were estimated by using a comparison to the same data. For small datums, with poorly distributed control, the better fitting shift type may not be more accurate. The seven-parameter model cannot accurately represent datum rotations and scale changes unless the control is geometrically well distributed across the datum.

c. Regional ΔX, ΔY, ΔZ shifts were developed when there were sufficient data to identify regional variations over the datum as a whole. Regional ΔX, ΔY, ΔZ shifts are more accurate than mean-value ΔX, ΔY, ΔZ shifts.

8.2.3 <u>Implementation issues</u>. System implementation issues may influence the selection of transformation method(s).

a. Computational speed and efficiency are important for some applications, such as real-time navigation. The shift type and method might be selected from those that meet accuracy requirements to minimize computational time. Alternately, shifts might be computed in a background process and applied as fixed values for a period of time or over a given area. The system designer must analyze the impact of such design strategies on system accuracy and performance.

b. For some applications, system computational and data storage capability are considerations in selecting and implementing transformations. The system designer must determine that word precision, range, and accuracy of mathematical functions preserve the required accuracy.

c. Code complexity, modularity, and software maintenance are important system design concerns. While the datum shift algorithms themselves are relatively simple, selection logic often is complex.

d. Both future requirements and present needs should be considered. Factors for consideration include: timing requirements, current accuracy requirements, potential future accuracy requirements (such as requirements to provide data for future systems), effects of improved databases, and the effects of improvements in other inputs. In other words, the savings in future planned product improvements should be evaluated against today's costs.

e. It is possible that a software developer may choose to use a transformation technique that is more accurate than is required. In such a case, the results are not necessarily superior to those arising from a less

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accurate, but acceptable, technique. The data are simply degraded less by the more accurate method; the transformed data can be no better than the input data. For example, techniques that achieve geodetic accuracy must be used when geodetic accuracy is needed. Such techniques can be used with a 1:250,000-scale product, but the transformed data still have the accuracy of a 1:250,000 product.

8.2.3.1 <u>Preferred shift methods</u>. Preferred shift methods and shift parameters may change as additional information becomes available. Shift parameters and identification information should be treated as data apart from the algorithm code so that they may be updated without modifying and recompiling the code. Some examples of implementation tradeoffs are:

a. If Cartesian coordinate are already available, ΔX, ΔY, ΔZ shifts can be applied directly (see Sections 6.3.1 and 6.3.2) rather than through the Molodensky approximations (see Section 6.3.3).

b. If the program contains algorithms for converting between Cartesian and geographic coordinates (see Section 4.3.4), the three-step method (see Sections 6.3.1.1 and 6.3.2.1) can be used instead of Molodensky approximations (see Section 6.3.3) to apply ΔX, ΔY, ΔZ shifts.

c. Both the Molodensky approximations (Section 6.3.3) and Cartesian-to-geographic conversion equations (Section 4.3.4.2) become singular at the poles. Selection of one technique or the other will reduce the logic needed to guard against singularity. An alternate equation to compute geodetic height from Cartesian coordinates in polar regions is given as equation 4.12.

8.2.3.2 <u>Software testing</u>. When developing software based on methods discussed in this Technical Report, it is advisable to check the software against test points. Test points can be found in Section 8.5.

a. Issues relating to software development and system performance might force a software developer to use a special-purpose transformation technique that is not mentioned in this Technical Report. The results of the special-purpose software should be compared to those of the techniques in this Technical Report to insure that the differences are acceptable.

b. Another useful debugging aid is **circular testing**. This is done by using the output from a transformation as the input for the inverse function; in other words trying to verify the original numbers are returned within

acceptable accuracy. Since the code for a transformation
and its inverse might contain mistakes that cancel each
other, circular testing can not prove that code is
correct. At best, it can demonstrate that code is
faulty.

c. The following is a list of points that might prove
    troublesome to coordinate transformation software. Test
    the code at these points, if it is appropriate:

    Poles
    Polar regions
    Discontinuity in λ (transition between 0,360 degrees or
    -180, 180 degrees)
    Origin expressed in (ϕ, λ)
    ϕ = 0
    λ = 0
    Both hemispheres in ϕ (North and South)
    Both hemispheres in λ (East and West)
    x = 0
    y = 0
    All four x-y quadrants
    Discontinuities in x and y at zone boundaries, equator,
       etc.

    8.3 <u>Error analysis</u>. Error analysis is an important part of
the system design process. Error analysis will estimate the
accuracy of the system's output products. This is a prediction
of how well the system will meet its accuracy requirements. The
system analyst can use error analysis to perform system
implementation studies. The major steps in error analysis are:

a. Develop the system model.

b. Determine the estimated errors for each source in the
    model.

c. Reformulate the estimated errors to a common description.

d. Propagate the errors through the model.

Each of these topics is discussed below. An example is given in
8.3.5.

    8.3.1 <u>System model</u>. The system model simply is a
description of the data used in the process and all the steps
(processes) used to process the data. The data include the raw
input data and parameters, like datum shift parameters, used in
the processing of the data. Both hardware-driven and software-
driven processes should be included in the model. For example,
errors in digitizing and plotting are primarily hardware related;

coordinate conversion and datum shifting are primarily software processes. For the datum shifting and coordinate conversion, the system model will probably be a simple, serial process. That is, the output of one process will be the input for the next, with no feedback to prior processes. For trade-off studies, alternate system models can be developed to evaluate accuracy versus implementation.

8.3.2 <u>Error estimates</u>. Once the system error model has been described, error estimates must be assigned to each process and data source in the model. When this is done, the analyst will quickly determine which error sources can be neglected and which must be propagated through the model.

   a. Estimates of the expected errors in the input data may be provided with the data or in product specifications. Also, the original source or intended use of the data may provide an indication of its accuracy (e.g. digital data derived from a cartographic source will have an accuracy no better than the cartographic source).

   b. Without proper guidance, it is often difficult to estimate the error in the data. This is especially true for cartographically derived elevations, because map accuracy standards for elevation allow for error in both the horizontal location and the vertical direction. Therefore, the allowable elevation error is dependent on the topography and may vary within a map sheet.

   c. In addition to providing datum shift parameters, Appendices B and E give the estimated errors in the parameters for ΔX, ΔY, ΔZ; and seven-parameter shifts, respectively. (Note: Correlations between parameters, which may be significant, were not available. Therefore, they are not given in the appendices.)

   d. The accuracy and behavior of the equations presented in this Technical Report are well defined. Table 8.1 identifies the location of accuracy guidance for the coordinate conversion, datum shift, and map projection equations. Additional references describing equation accuracy and behavior are noted in the text and listed in section 2.

   e. The errors introduced by system implementation factors such as computer word size, accuracy and precision of mathematical functions, truncation of terms, etc. must be quantified.

   f. The errors introduced by hardware processes must be determined from equipment specifications, analysis, and/or testing.

Table 8.1. Accuracy guidance for equations

Subject	Accuracy guidance section(s)
Coordinate Conversion Equations	
Cartesian/geodetic conversion	4.3.4
Height conversion	4.4.3
Datum shift Equations	
Seven-parameter geometric transformation	6.2.4
$\Delta X, \Delta Y, \Delta Z$ shifts	6.3.1, 6.3.2
Direct WGS 72 to WGS 84 transformation	6.4.2.2
Vertical datum shifts	6.7, 6.7.1, 6.7.2
Map projection Equations	
Mercator projection	7.2.3
Transverse Mercator projection	7.3.3
Universal Transverse Mercator (UTM) grid	7.4.4
Lambert Conformal Conic projection	7.5.4
Polar Stereographic	7.6.2
Universal Polar Stereographic	7.7.4
Shift Parameter Accuracy values	
$\Delta X, \Delta Y, \Delta Z$ shifts	Appendix B
Seven-parameter geometric transformation	Appendix E
Direct WGS 72 to WGS 84 transformation	Section 6.4.2.2

8.3.2.1 <u>Error types</u>. Depending on the source, error estimates may be described in several ways. Error descriptions may have a bias component and a random (noise) component. A set of error estimates describing the accuracy in multiple dimensions, for example, $\sigma_{\Delta X}$, $\sigma_{\Delta Y}$, $\sigma_{\Delta Z}$, may have correlations between the parameters. Generally, with MC&G products, the error descriptions do not differentiate between biased, correlated, and independent random errors; the assumption is made that the error estimates are unbiased, independent and random. The resultant error analysis will be less accurate than if the true statistical natures of the error sources were known.

8.3.3 <u>Reformulation</u>. Once the error sources have been quantified, they must be converted to a common frame of reference. The form used for the system output error estimate is often a convenient frame of reference.

a. The reference for the error estimates must be established
    (i.e. the ellipsoid or the map sheet). It may be
    necessary to scale from one reference to another.

b. The error estimates should be converted to the same
    units.

c. The datum shift parameter error estimates, $\sigma_{\Delta X}$, $\sigma_{\Delta Y}$, $\sigma_{\Delta Z}$,
    given in Appendix B, can be reformulated to $\sigma_{\Delta\phi}$, $\sigma_{\Delta\lambda}$, $\sigma_{\Delta H}$
    using equation set 6.8.

d. Error estimates typically are linear, circular, or
    spherical statistical values with associated
    probability levels. Error estimates should be converted
    to the same type of statistic at the same probability
    level. Two independent linear standard deviations
    (68.27% probability), σ_x and σ_y, can be converted to a
    circular standard deviation (39.35% probability), using

$$\sigma_c \approx \frac{(\sigma_x + \sigma_y)}{2} \qquad (8.1)$$

    if the values of σ_x and σ_y are reasonably close.
    Factors to convert between different levels of
    probability are given in Table 8.2.

8.3.4 Error propagation. Assuming the error estimates are
statistically independent and have been converted to the same
units and statistical type, they may be propagated through a
simple, serial system model using the Root-Sum-Square (RSS)
technique.

$$RSS = \left[(N_1)^2 + (N_2)^2 + \dots + (N_n)^2 \right]^{\frac{1}{2}} \qquad (8.2)$$

If error estimates have known biases or correlations, then a more
sophisticated error propagation technique should be used.

Table 8.2. Probability Level Conversion Factors

Notation				
Error Form	Standard Error (One Sigma)	Probable Error	Map Accuracy Standard	Near Certainty Error
Linear	σ (68.27%)	PE (50%)	MAS (90%)	3σ (99.73%)
Circular	σ_c (39.35%)	CEP (50%)	CMAS (90%)	$3.5\sigma_c$ (99.78%)
Spherical	σ_s (19.9%)	SEP (50%)	**SAS (90%)	$4\sigma_s$ (99.89%)
**SAS = Spherical Accuracy Standard				

Linear Error (One dimensional)				
From \ To	50%	68.27%	90%	99.73%
50%	1.0000	1.4826	2.4387	4.4475
68.27	0.6745	1.0000	1.6449	3.0000
90	0.4101	0.6080	1.0000	1.8239
99.73	0.2248	0.3333	0.5483	1.0000

Circular Error (Two dimensional)				
From \ To	39.35%	50%	90%	99.78%
39.35%	1.0000	1.1774	2.1460	3.5000
50	0.8493	1.0000	1.8227	2.9726
90	0.4660	0.5486	1.0000	1.6309
99.78	0.2857	0.3364	0.6131	1.0000

Spherical Error (Three Dimensional)				
From \ To	19.9%	50%	90%	99.89%
19.9%	1.000	1.538	2.500	4.000
50	0.650	1.000	1.625	2.600
90	0.400	0.615	1.000	1.600
99.89	0.250	0.385	0.625	1.000

(Note: Errors are assumed to be random and uncorrelated.)

8.3.5 _Example_. The commander wants a printed product showing the transportation routes near Seoul, Korea with WGS 84, UTM coordinates, at a scale of 1:50,000. The data source is a Class A, 1:50,000-scale topographic map on the Tokyo datum. The product will be created by digitizing the transportation routes on the map sheet, shifting the coordinates from the Tokyo datum to WGS 84 using the three-step method, and plotting the shifted data.

a. The system model is shown in Figure 8.5. In this model, it is assumed that computer-related issues will not degrade coordinate conversion and datum shift calculations. Also, it is assumed that the representation of the digitized data will not change through the rest of the process. Often this assumption is not valid because digital data representation must be changed to a predetermined format. Such data modifications can introduce error.

b. The error estimates for each of the elements in the model are

Element	Error
Map accuracy (horizontal)	0.02" CMAS(90%)
Digitizer	±0.005"
UTM to Geodetic conversion	<1 cm
Geodetic to XYZ conversion	0.0
Datum shift equations	0.0
Datum shift parameters: $\sigma_{\Delta X}$	5 m
$\sigma_{\Delta Y}$	3 m
$\sigma_{\Delta Z}$	3 m
XYZ to Geodetic conversion	0.1 m
Geodetic to UTM conversion	<1 cm
Plotter	±0.01"

Only the map, digitizer, datum shift parameter, and plotter errors are significant and will be considered in the rest of the analysis.

c. The error estimates must be reformulated to the same units and statistical basis. Since the commander wants to know the estimated error on the ground, we will convert to circular error, in meters, on the ellipsoid, at 50 percent probability (CEP).

1. Map error must be scaled from the map to the ellipsoid and changed from CMAS to CEP.

Map Error = 0.02" CMAS x 0.0254 m/in x 50,000
= 25.4 m CMAS x 0.5486
= 13.9 m CEP

2. It is assumed that the digitizer error is a near certainty, circular error specification that must be scaled from the map to the ellipsoid and converted to CEP.

Digitizer error = 0.005" max x 0.0254 m/in x 50,000
= 6.4 m max x 0.3364
= 2.1 m CEP

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3. The datum shift parameter error estimates, $\sigma_{\Delta X}$, $\sigma_{\Delta Y}$, $\sigma_{\Delta Z}$, must be converted to $\sigma_{\Delta\phi}$ and $\sigma_{\Delta\lambda}$, then to CEP.

Using the above values for $\sigma_{\Delta X}$, $\sigma_{\Delta Y}$, and $\sigma_{\Delta Z}$, ϕ = 37.5° and λ = 127°, equation set 6.8 yields $\sigma_{\Delta\phi}$ = 3.34 m and $\sigma_{\Delta\lambda}$ = 4.38 m.

$$CEP = 1.1774 \times \sigma_c \approx 1.1774 \times (\sigma_{\Delta\phi} + \sigma_{\Delta\lambda})/2 = 4.5 \text{ m}$$

4. It is assumed that the plotter error is a near certainty, circular error (99.78% probability) specification that must be scaled from the map to the ellipsoid and converted to CEP.

$$
\begin{aligned}
\text{Plotter error} &= 0.01" \text{ max} \times 0.0254 \text{ m/in} \times 50,000 \\
&= 12.7 \text{ m max} \times 0.3364 \\
&= 4.3 \text{ m CEP}
\end{aligned}
$$

5. The RSS error for the complete system model is

$$RSS = \left[(13.9)^2 + (2.1)^2 + (4.5)^2 + (4.3)^2 \right]^{\frac{1}{2}}$$
$$= 15.4 \text{ m CEP}$$

d. From the above analysis, it is evident that the error in the source data is the prime contributor to error in the output product for this application. For other types of applications, such as surveying, this may not be the case.

Figure 8.5. System Model

```
           ╭─────────────────╮
          ╱    1:50,000       ╲
         │   Class A Map       │
          ╲                   ╱
           ╰─────────────────╯
                    │
                    ▼
          ┌─────────────────┐
          │    Digitize      │
          └─────────────────┘
                    │
                    ▼
          ┌─────────────────┐
          │  Convert UTM to  │
          │    Geodetic      │
          └─────────────────┘
                    │
                    ▼
          ┌─────────────────┐
          │ Convert Geodetic │
          │    to XYZ        │
          └─────────────────┘
                    │
   ╭────╮           ▼
  ╱ 3-Step╲   ┌─────────────────┐      ╭──────────────────╮
 │  Datum   │ │ Shift Tokyo Datum│◄─────│  Datum Shift     │
 │  Shift   │ │   to WGS 84      │      │  Parameters      │
  ╲        ╱  └─────────────────┘      │  ΔX, ΔY, ΔZ      │
   ╰────╯            │                 ╰──────────────────╯
                     ▼
          ┌─────────────────┐
          │  Convert XYZ     │
          │  to Geodetic     │
          └─────────────────┘
                    │
                    ▼
          ┌─────────────────┐
          │ Convert Geodetic │
          │    to UTM        │
          └─────────────────┘
                    │
                    ▼
          ┌─────────────────┐
          │      Plot        │
          └─────────────────┘
                    │
                    ▼
           ╭─────────────────╮
          ╱  Transportation   ╲
         │    Route Map        │
          ╲                   ╱
           ╰─────────────────╯
```
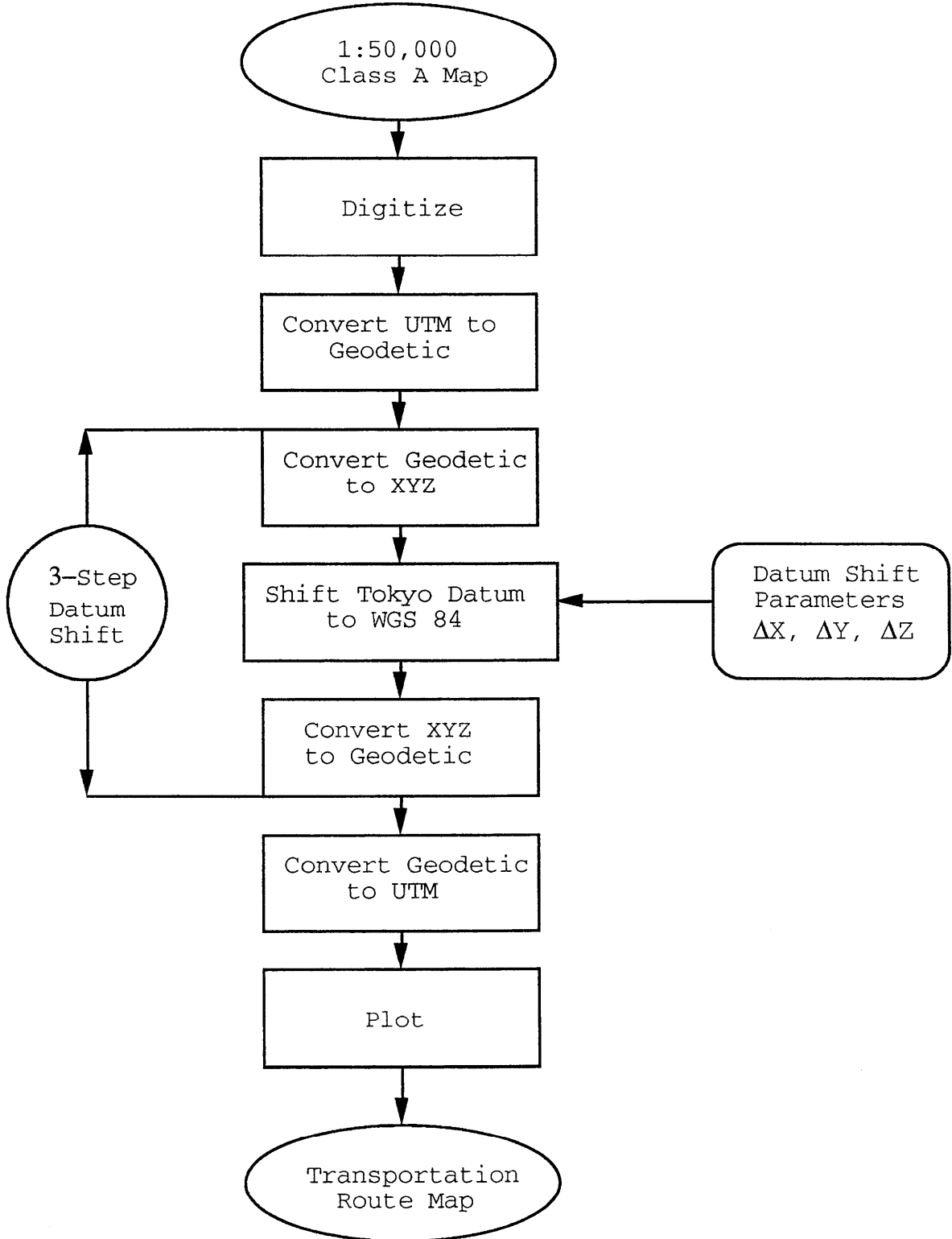
8.4 Numerical Examples.

8.4.1 Example 8.4. Convert WGS 84 geodetic coordinates $(\phi, \lambda, h)_{WGS\ 84}$ to Universal Transverse Mercator coordinates (x_{UTM}, y_{UTM}) in the NAD 27 reference system using a three-step ΔX, ΔY, ΔZ datum shift.

This example converts WGS 84 geodetic coordinates to Universal Transverse Mercator projection coordinates in two steps. First, the WGS 84 geodetic coordinates are converted to NAD 27 geodetic coordinates using the three-step ΔX, ΔY, ΔZ datum shift. Geodetic height is needed for this conversion. Next, the NAD 27 geodetic latitude and longitude are converted to UTM coordinates. Geodetic height is not needed for this step.

Beginning WGS 84 coordinates:

ϕ = 42° 56' 52.163"N = 0.749580918883376 radians

λ = 288° 22' 24.326"E = 1.891473118 radians

h = 203.380 m

Step 1: Convert WGS 84 geodetic coordinates to NAD 27 geodetic coordinates using the three-step ΔX, ΔY, ΔZ datum shift.

1.1 Convert $(\phi, \lambda, h)_{WGS\ 84}$ to $(X, Y, Z)_{WGS\ 84}$ using equation sets 4.4 and 4.5.

a = 6378137

f = 0.00335281066474

ε^2 = 2f - f^2 = 0.00669437999013

b = a(1 - f) = 6356752.3142

$$R_N = \frac{a}{\sqrt{1 - \varepsilon^2 \sin^2 \phi}} = 6388070.57383009$$

$X_{WGS\ 84} = \left(R_N + h\right) \cos\phi \cos\lambda = 1473933.5413$

$Y_{WGS\ 84} = \left(R_N + h\right) \cos\phi \sin\lambda = -4437679.0666$

$$Z_{WGS\ 84} = \left(\frac{b^2}{a^2}\ R_N + h\right)\ \sin\phi = 4323399.2717$$

1.2   Convert $(X,Y,Z)_{WGS\ 84}$ to $(X,Y,Z)_{NAD\ 27}$ using equation set 6.4 and $\Delta X, \Delta Y, \Delta Z$ datum shift values from Table B.1 NAS-C.

$$X_{NAD\ 27} = X_{WGS\ 84} - \Delta X = 1473933.5413 - (-8.0) = 1473941.5413\ m$$

$$Y_{NAD\ 27} = Y_{WGS\ 84} - \Delta Y = -4437679.0666 - 160.0 = -4437839.0666\ m$$

$$Z_{NAD\ 27} = Z_{WGS\ 84} - \Delta Z = 4323399.2717 - 176 = 4323223.2717\ m$$

1.3 Convert $(X,Y,Z)_{NAD\ 27}$ to $(\phi, \lambda)_{NAD\ 27}$ following equations given in Section 4.3.4.2.

$$\lambda_{NAD\ 27} = \arctan\left(\frac{Y_{NAD\ 27}}{X_{NAD\ 27}}\right) = 5.033056560575\ radians$$

Calculating ϕ is an iterative process.

1.3.1   Compute an initial approximation to a variable β, using equation 4.7.   Parameters a and b for the Clarke 1866 ellipsoid are found in Table A.

$$\beta_0 = \arctan\left[\frac{a\ Z}{b\ \sqrt{X^2 + Y^2}}\right] = 0.74788820090716142\ radians$$

1.3.2   Substitute β_O for β in equation 4.8 to get an approximation for ϕ.

$$\phi = \arctan\left[\frac{Z + \varepsilon'^2\ b\ \sin^3\ \beta}{\sqrt{X^2 + Y^2}\ -\ a\ \varepsilon^2\ \cos^3\ \beta}\right] = 0.74958142853674489\ radians$$

$$\varepsilon^2 = 2f - f^2 = 0.0067686579973$$

and

$$\varepsilon'^2 = \frac{e^2}{1 - e^2} = 0.0068147849459$$

For most applications, this first iteration is sufficient.

Step 2:   Convert the NAD 27 geodetic latitude and longitude to Universal Transverse Mercator projection coordinates.

2.1   Find the UTM zone z and the central meridian λ_O from equations 7.22 and 7.23.

$$z = \text{greatest integer} \le \left(\frac{180 \ \lambda}{6\pi} - 29 \right)$$

$$z = 19$$

$$\lambda_0 = (6z + 177) \frac{\pi}{180} = 5.078908123303 \text{ radians}$$

2.2 Calculate Transverse Mercator projection coordinates (X_{TM}, Y_{TM}) using equations 7.8 and 7.9.

2.2.1 Calculate terms in equation set 7.9 using ε^2 and ε'^2 as calculated in Step 1.3.2.

$$A_o = 1 - \frac{1}{4}\varepsilon^2 - \frac{3}{64}\varepsilon^4 - \frac{5}{256}\varepsilon^6 - \frac{175}{16384}\varepsilon^8 = 0.998305681856$$

$$A_2 = \frac{3}{8} \left(\varepsilon^2 + \frac{1}{4}\varepsilon^4 + \frac{15}{128}\varepsilon^6 - \frac{455}{4096}\varepsilon^8 \right) = 0.2542555420215 \times 10^{-2}$$

$$A_4 = \frac{15}{256} \left(\varepsilon^4 + \frac{3}{4}\varepsilon^6 - \frac{77}{128}\varepsilon^8 \right) = 0.2698010542695 \times 10^{-5}$$

$$A_6 = \frac{35}{3072} \left(\varepsilon^6 - \frac{41}{32}\varepsilon^8 \right) = 0.3502448582027 \times 10^{-8}$$

$$A_8 = - \frac{315}{131072} \varepsilon^8 = -0.5044416191579 \times 10^{-11}$$

$$S\phi = a \ [A_0\phi - A_2 \sin 2\phi + A_4 \sin 4\phi - A_6 \sin 6\phi + A_8 \sin 8\phi]$$

$$S\phi = 4756711.680282$$

$$N = R_N = \frac{a}{\sqrt{1 - \varepsilon^2 \sin^2 \phi}} = 6388250.562973$$

Note that $N = R_N$ differs from R_N calculated in Step 1.1, as Step 2 is in the NAD 27 reference frame and Step 1.1 is in the WGS 84 reference frame.

$$\eta = \varepsilon' \cos \phi = 0.6042571576496 \times 10^{-1}$$

$$t = \tan\phi = \tan(0.749580094822852639) = 0.9308149267480$$

$$\Lambda = \lambda - \lambda_0 = -0.04585156272818 \text{ radians}$$

2.2.2 Calculate the Transverse Mercator projection coordinates (x_{TM}, y_{TM}).

$$x_{TM} = N\ \Lambda\ \cos\phi + \frac{N\ \Lambda^3\ \cos^3\ \phi}{6}\ (1 - t^2 + \eta^2)$$

$$+ \frac{N\ \Lambda^5 \cos^5\phi}{120}\ (5 - 18t^2 + t^4 + 14\eta^2 - 58t^2\eta^2)$$

$$y_{TM} = S_\phi + \frac{N\ \Lambda^2}{2}\ \sin\ \phi\ \cos\ \phi$$

$$+ \frac{N\ \Lambda^4}{24}\ \sin\ \phi\ \cos^3\ \phi (5 - t^2 + 9\eta^2 + 4\eta^4)$$

$$+ \frac{N\ \Lambda^6}{720}\ \sin\ \phi\ \cos^5\ \phi (61 - 58t^2 + t^4 + 270\eta^2 - 330t^2\eta^2)$$

$$x_{TM} = -214408.9715411$$

$$y_{TM} = 4760061.989246$$

2.3 Convert the Transverse Mercator projection coordinates (x_{TM}, y_{TM}) to Universal Transverse Mercator projection coordinates (x_{UTM}, y_{UTM}) using equation set 7.24 (Northern hemisphere).

$$x_{UTM} = 0.9996\ x_{TM} + 500,000 = 0.9996\ (-214408.9715411) + 500,000$$

$$y_{UTM} = 0.9996\ y_{TM} = 0.9996\ (4760061.989246)$$

The resulting Universal Transverse Mercator coordinates (x_{UTM}, y_{UTM}) in the NAD 27 reference system are:

$$x_{UTM} = 285676.792$$

$$y_{UTM} = 4758157.964$$

8.4.2 <u>Example 8.5</u>. Convert WGS 84 geodetic coordin (ϕ, λ, h) to geodetic coordinates (ϕ, λ, h) in the NAD 27 re system using the standard Molodensky method.

This example converts WGS 84 geodetic coordinates to NAD 27 coordinates using the standard Molodensky method.

Beginning WGS 84 coordinates:

ϕ = 42° 56'52.163" N = 0.749580918883376 radians
λ = 108° 22'24.326" W = 1.891473118 radians
h = 203.380

Step 1: Compute Δa, Δf, W^2, M, N, ε^2, and ε'^2 using equation set 6.5. Obtain $a_{WGS\ 84}$, $A_{NAD\ 27}$, $f_{WGS\ 84}$, and f_{NAD27} from Table A.1.

$$\Delta a = a_{WGS\ 84} - a_{CLARKE\ 1866} = -69.4$$

$$\Delta f = f_{WGS\ 84} - f_{CLARKE\ 1866} = 3.726463863 \times 10^{-5}$$

$$\varepsilon^2 = 2f - f^2 = 0.006694379990$$

$$\varepsilon'^2 = \frac{\varepsilon^2}{1-\varepsilon^2} = 0.006739496742$$

$$W = \sqrt{1 - \varepsilon^2 \sin^2\phi} = 0.9984449805$$

$$N = R_N = \frac{a}{W} = 6388070.574$$

$$M = \frac{a(1 - \varepsilon^2)}{W^3} = 6365086.681$$

Step 2: Compute $\Delta\phi$ and $\Delta\lambda$, and Δh using equation set 6.5, and ΔX, ΔY, and ΔZ datum shifts from Appendix B for datum code NAS-A. In this example, the signs of ΔX, ΔY, and ΔZ are reversed from those of Appendix B due to the direction of the datum shift (WGS 84 to local coordinates).

$$\Delta\phi = \left[(-\sin\phi\cos\lambda\Delta x - \sin\phi\sin\lambda\Delta y + \cos\phi\Delta z + \frac{\varepsilon^2\sin\phi\cos\phi}{W}\Delta a + \right.$$
$$\left. \sin\phi\cos\phi(2N + \varepsilon'^2 M\sin^2\phi)(1 - f)\Delta f\right] / [M + h]$$

$$\Delta\phi = 0.131"$$

$$\Delta\lambda = \frac{[-\sin\lambda\Delta x + \cos\lambda\Delta y]}{\{(N + h)\cos\phi]} = 2.615"$$

$$\Delta h = \cos\phi\cos\lambda\Delta x + \cos\phi\sin\lambda\Delta y + \sin\phi\Delta z - W\Delta a + \frac{a(1 - f)}{W} \sin^2 f\ \Delta f$$

$$\Delta h = 28.65\ m$$

Step 3: Compute the NAD 27 geodetic coordinates (ϕ,λ,h) from equation set 6.7 that results in the following values1;

$$\phi_{NAD\ 27} = \phi_{WGS\ 84} + \Delta\phi = 42° \ 56'52.294"N$$

$$\lambda_{NAD\ 27} = \lambda_{WGS\ 84} + \Delta\lambda = 108° \ 22'21.711"W$$

$$h_{NAD\ 27} = H_{WGS\ 84} + \Delta H = 232.03 \ m$$

8.4.3 Example 8.6. Convert WGS 84 geodetic coordinates (ϕ, λ)$_{WGS\ 84}$ to Mercator projection coordinates (x,y).

This example converts WGS 84 geodetic coordinates to Mercator projection coordinates within the WGS 84 reference system. Geodetic height (h) is not used in this conversion. The Mercator projection central meridian λ_o is 288° = 5.026548246 radians

Beginning WGS 84 coordinates are

$$\phi_{WGS\ 84} = \ 42° \ 56' \ 52.163"N = 0.7495809189 \ radians$$

$$\lambda_{WGS\ 84} = 288° \ 22' \ 24.326"E = 5.033057346 \ radians$$

Step 1: Use equation 7.1 and WGS 84 datum parameters a and f from Table A.1 to calculate x.

$$x = a(\lambda - \lambda_o) = 6378137(5.033057346 - 5.026548246) = 41569.357$$

Step 2: Calculate ε from f using equation 4.9, then use equation 7.2 to calculate y.

$$\varepsilon = \sqrt{2f - f^2} \ = 0.0818191908426$$

$$y = a \ \ln \left[\tan \left(\frac{\pi}{4} + \frac{\phi}{2} \right) \cdot \left(\frac{1 - \varepsilon \sin \phi}{1 + \ \ \varepsilon \sin \phi} \right)^{\frac{\varepsilon}{2}} \right] = 5274911.868$$

The resulting Mercator projection coordinates (x,y) are

x = 41569.357

y = 5274911.868

8.4.4 Example 8.7. Converst Mercator projection coordinates (x,y) in the WGS 84 reference system to WGS 84 geodetic latitude and longitude (ϕ,λ).

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This example converts WGS 84 Mercator coordinates to WGS 84 latitude and longitude.

Beginning Mercator coordinates and central meridian (λ_O):

x = 41569.3572
y = 5274911.8684
λ_O = 288°E = 5.0265482457 radians

Step 1: Compute the longitude (λ) using equation 7.4 where a is the semi-major axis of the WGS 84 ellipsoid listed in Table A.1.

$$\lambda = \frac{x}{a} + \lambda_O = 5.0330657221 \text{ radians} = 288° \ 22'24.326"E \text{ or}$$
$$71° \ 37' \ 35.674"W$$

Step 2: Compute an intermediate value t from equation 7.5.

$$t = e^{(-\frac{y}{a})} = e^{(-\frac{5274911.8684}{6378137})} = .43734619376064$$

Step 3: Compute an initial value of the geodetic latitude (ϕ_O) from equation 7.6.

$$\phi_O = \frac{\pi}{2} - 2 \ \arctan(t) = 0.746233651 \text{ radians}$$

Step 4: Iterate equation 7.7 until the updated value of ϕ is sufficiently close to the previous value of ϕ.

$\varepsilon^2 = 2f - f^2 = 0.00669437999013$

$\varepsilon = 0.0818191908426$

$$\phi_{n+1} = \frac{\pi}{2} - 2 \ \arctan \left(t \left[\frac{1 - \varepsilon \ \sin \ \phi_n}{1 + \varepsilon \ \sin \ \phi_n} \right]^{\frac{\varepsilon}{2}} \right)$$

$\phi_1 = 0.7495688569418636$

The geodetic latitude will be given by the last value of ϕ calculated.

Final ϕ = 0.7495809188833704 radians = 42° 56' 52.163"N

The resulting WGS 84 geodetic latitude and longitude (ϕ,λ) are:

ϕ = 42° 56' 52.163"N

$\lambda =$ 288° 22' 24.326"E or 71° 37' 35.674"W

8.4.5 <u>Example 8.8</u>. Convert Universal Transverse Mercator coordinates (x_{UTM},y_{UTM}) in the NAD 27 reference system to NAD 27 geodetic latitude and longitude(ϕ, λ).

This example converts NAD 27 UTM coordinates to NAD 27 geodetic latitude and longitude. The UTM coordinates are converted to Transverse Mercator projection coordinates. The Transverse Mercator projection coordinates are then converted to NAD 27 latitude and longitude.

Beginning coordinates are
x_{UTM} = 285,677.332
y_{UTM} = 4,758,154.856
Zone = 19

Step 1: Convert Universal Transverse Mercator coordinates (x_{UTM},y_{UTM}) to Transverse Mercator coordinates (x_{TM},y_{TM}) using equation set 7.25.

$$x_{TM} = \frac{1}{0.9996} (x_{UTM} - 500,000) = -214,408.4314$$

$$y_{TM} = \frac{1}{0.9996} y_{UTM} = 4,760,058.8796$$

Step 2: Convert Transverse Mercator coordinates (x_{TM},y_{TM} to geodetic latitude and longitude(ϕ, λ).

2.1 Calculate the UTM zone central meridian (λ_0) using equation 7.23.

$$\lambda_0 = (6z + 177)\left(\frac{\pi}{180}\right) = 5.078908123303 \text{ radians}$$

2.2 Compute the eccentricity (ε) and second eccentricity (ε') of the Clarke 1866 ellipsoid, the reference ellipsoid for NAD 27, using equation set 7.9.

$$\varepsilon^2 = 2f - f^2 = 0.6768657997291 \times 10^{-2}$$

$$\varepsilon'^2 = \frac{\varepsilon^2}{1 - \varepsilon^2} = 0.6814784945915 \times 10^{-2}$$

2.3 Calculate the subscripted-A terms of equation set 7.9, using ε^2 calculated in Step 2.1. The subscripted-A terms are used in the calculation of the footpoint latitude.

$$A_0 = 1 - \frac{1}{4}\varepsilon^2 - \frac{3}{64}\varepsilon^4 - \frac{5}{256}\varepsilon^6 - \frac{175}{16384}\varepsilon^8 = 0.998305681856$$

$$A_2 = \frac{3}{8}\left(\varepsilon^2 + \frac{1}{4}\varepsilon^4 + \frac{15}{128}\varepsilon^6 - \frac{455}{4096}\varepsilon^8\right) = 0.2542555420215 \times 10^{-2}$$

$$A_4 = \frac{15}{256}\left(\varepsilon^4 + \frac{3}{4}\varepsilon^6 - \frac{77}{128}\varepsilon^8\right) = 0.2698010542695 \times 10^{-5}$$

$$A_6 = \frac{35}{3072}\left(\varepsilon^6 - \frac{41}{32}\varepsilon^8\right) = 0.3502448582027 \times 10^{-8}$$

$$A_8 = -\frac{315}{131072}\varepsilon^8 = -0.5044416191579 \times 10^{-11}$$

2.4 Compute the footprint latitude, ϕ_1, using equation 7.10 and a from Table A.1.

Make an initial estimate for ϕ_1:

$$\phi_1 = \frac{Y_{TM}}{a} = 0.7463005398433 \text{ radians}$$

2.5 Calculate S_ϕ using equation set 7.9 and the estimate of ϕ, ϕ_1.

$$S_\phi = a[A_0\phi - A_2 \sin 2\phi + A_4 \sin 4\phi - A_6 \sin 6\phi + A_8 \sin 8\phi]$$

$$S_\phi = 0.4735829138159 \times 10^7$$

2.6 Calculate S'_ϕ using equation set 7.11 and the estimate of ϕ, ϕ_1.

$$S'_\phi = a[A_0 - 2A_2\cos 2\phi + 4A_4 \cos 4\phi - 6A_6 \cos 6\phi + 8A_8 \cos 8\phi]$$

$$S'_\phi = 0.6364798134276 \times 10^7$$

2.7 Find an updated value of the footpoint latitude, ϕ_1 using equation 7.10.

$$\phi_1(\text{updated}) = \phi_1(\text{current}) - \frac{S_\phi - y}{S'_\phi} = 0.7501073756426 \text{ radians}$$

2.8 Iterate Steps 2.5 through 2.7 until successive values of ϕ_1 are as close as is desired.

For this example, the final value of ϕ_1 is:

Final ϕ_1 = 0.7501073020563 radians

     2.9 Calculate values used in equation set 7.12 using equation set 7.13.

$$t_1 = \tan \phi_1 = \tan (0.7501073020563)$$

$$N_1 = \frac{a}{\sqrt{1 - \varepsilon^2 \sin^2 \phi_1}} = 6{,}388{,}261.949762$$

$$R_1 = \frac{a (1 - \varepsilon^2)}{\left(\sqrt{1 - \varepsilon^2 \sin^2 \phi_1}\right)^3} = 6{,}365{,}044.225886$$

$$\eta_1^2 = \varepsilon'^2 \cos^2 \phi_1 = 0.3647692467262 \times 10^{-2}$$

$$B_3 = 1 + 2t_1^2 + \eta_1^2 = 2.740138642336$$

$$B_4 = 5 + 3t_1^2 + \eta_1^2 - 4\eta_1^4 - 9t_1^2\eta_1^2 = 7.579827062321$$

$$B_5 = 5 + 28t_1^2 + 24t_1^4 + 6\eta_1^2 + 8t_1^2\eta_1^2 = 47.45050110665$$

$$B_6 = 61 + 90t_1^2 + 46\eta_1^2 + 45t_1^4 - 252t_1^2\eta_1^2 - 3\eta_1^4 - 66t_1^2\eta_1^4 - 90t_1^4\eta_1^2 + 225t_1^4\eta_1^4 = 172.1890097418$$

     2.10 Calculate the geodetic latitude and longitude (ϕ, λ) using equation set 7.12.

$$\phi = \phi_1 - \frac{t_1 N_1}{R_1} \left[\frac{1}{2}\left(\frac{x}{N_1}\right)^2 - \frac{B_4}{24}\left(\frac{x}{N_1}\right)^4 + \frac{B_6}{720}\left(\frac{x}{N_1}\right)^6\right]$$

$$\phi = 0.74958094314009481 \text{ radians} = 42^O\ 56'\ 52.168"$$

$$\lambda = \lambda_0 + \sec \phi_1 \left[\frac{x}{N_1} - \frac{B_3}{6}\left(\frac{x}{N_1}\right)^3 + \frac{B_5}{120}\left(\frac{x}{N_1}\right)^5\right]$$

$$\lambda = 5.03305669671 \text{ radians} = 288^O\ 22'\ 22.464"$$

     8.4.6 <u>Example 8.9</u>. Convert WGS 84 latitude and longitude (ϕ, λ) to Lambert Conformal coordinates (x, y) using two standard parallels.

This example converts WGS 84 latitude and longitude to Lambert conformal projection coordinates in the WGS 84 reference system.

Beginning WGS 84 coordinates are

$\phi_{WGS\ 84}$ = 42° 56' 52.163"N = 0.749580918883376 radians

$\lambda_{WGS\ 84}$ = 288° 22' 24.326"E = 5.033065722110382 radians

Standard parallels:

ϕ_1 = 42° 30' 00.000" = 0.7417649320975901 radians

ϕ_2 = 43° 00' 00.000" = 0.7504915783575617 radians

Origin:

ϕ_0 = 42° 30' 00.000" = 0.7417649320975901 radians

λ_0 = 288° 00' 00.000" = 5.026548245743669 radians

Step 1: Calculate eccentricity ε from flattening f.

ε = $\sqrt{2f - f^2}$ = 0.0818191908426

Step 2: Calculate q at ϕ, ϕ_1, ϕ_2 and ϕ_0 using equation set 7.27.

$$q(\phi_i) = \ln\left[\tan\left(\frac{\pi}{4} + \frac{\phi_i}{2}\right) \cdot \left(\frac{1 - \varepsilon \sin \phi_i}{1 + \varepsilon \sin \phi_i}\right)^{\frac{\varepsilon}{2}} \right]$$

where i = null, 1, 2, 0

$q(\phi)$ = 0.8270301921136922

$q(\phi_1)$ = 0.8164293462094778

$q(\phi_2)$ = 0.8282703592274435

$q(\phi_0)$ = 0.8164293462094778

Step 3: Calculate N at ϕ_1 and ϕ_2 using equation set 7.27 and a from Table A-1.

$$N(\phi_i) = \frac{a}{\sqrt{1 - \varepsilon^2 \sin^2 \phi_i}}$$

$N(\phi_1) = 6387903.46785173$

$N(\phi_2) = 6388090.05758619$

Step 4: Calculate L from equation set 7.27.

$$L = \frac{\ln(N_1\cos\phi_1) - \ln(N_2\cos\phi_2)}{q_2 - q_1} = 0.6788029306559900$$

Step 5: Calculate K from equation set 7.27.

$$K = \frac{N_1 \cos \phi_1}{L \ \varepsilon^{-Lq_1}} = \frac{N_2 \cos \phi_2}{L \ \varepsilon^{-Lq_2}} = 12076169.25203636$$

Step 6: Calculate r and r_o from equation set 7.27.

$r = K\varepsilon^{-Lq(\phi)} = 6888432.18075871$

$r_0 = K\varepsilon^{-Lq(\phi_0)} = 6938179.32109766$

Step 7: Calculate the Lambert coordinates (x,y) using equation set 7.26.

$x = r \sin L(\lambda - \lambda_0) = 30474.8898082$

$y = r_0 - r \cos L(\lambda - \lambda_0) = 49814.5521555$

8.4.7 Example 8.10. Convery Lambert Conformal projection (with two standard parallels) coordinates (x,y) in the WGS 84 reference system to WGS 84 geodetic latitude and longitude (ϕ,λ).

This example converts WGS 84 Lambert Conformal coordinates to WGS 84 latitude and longitude.

Beginning Lambert Conformal coordinates, standard parallels (ϕ_1,ϕ_2) and origin (ϕ_O,λ_O) are

$x = 30474.890$
$y = 49814.552$
$\phi_1 = 42° 30'00" = 0.7417649320976$ radians
$\phi_2 = 43° 00'00" = 0.7504915783576$ radians
$\phi_o = 42° 30'00" = 0.7417649320976$ radians
$\lambda_o = 288° 00'00" = 5.0265482457437$ radians

Step 1: Compute ε, L, r_o, k, N_1, N_2, q_o, q_1, and q_2 using equation set 7.27. Obtain a from Table A.1.

$$\varepsilon = \sqrt{2f - f^2} = 0.0818191908426$$

$$q_i = \ln\left[\tan\left(\frac{\pi}{4} + \frac{\phi_i}{2}\right) \cdot \left(\frac{1 - \varepsilon \sin \phi_i}{1 + \varepsilon \sin \phi_i}\right)^{\frac{\varepsilon}{2}}\right]$$

$q_o = 0.8164293462094778$

$q_1 = 0.8164293462094778$

$q_2 = 0.8282703592274435$

$$N_i = \frac{a}{\sqrt{1 - \varepsilon^2 \sin^2 \phi_i}}$$

$N_1 = 6387903.467851727$

$N_2 = 6388090.057586191$

$$L = \frac{\ln(N_1\cos\phi_1) - \ln(N_2\cos\phi_2)}{q_2 - q_1} = 0.6788029306559713$$

$$K = \frac{N_1\cos\phi_1}{Le^{-Lq_1}} = \frac{N_2\cos\phi_2}{Le^{-Lq_2}} = 12076169.25203651$$

Step 2: Compute r_o using equation set 7.27.

$$r_O = Ke^{-Lq(\phi_0)} = 6938179.321097849$$

Step 3: Compute r and θ from equation set 7.29.

$$\theta = \arctan\frac{x}{r_o - y} = 4.4240820859501107 \times 10^{-3} \text{ radians}$$

$$r = \frac{r_o - y}{\cos\theta} = 6888432.18091526$$

Step 4: Compute geodetic longitude (λ) using equation set 7.28.

$$\lambda = \frac{\theta}{L} + \lambda_O = 5.0330657221511255 \text{ radians} = 288° \ 22' \ 24.326"$$

Step 5: Compute q using equation set 7.28.

$$q = \frac{\ln\left(\frac{k}{r}\right)}{L} = 0.8270301920802524$$

Step 6: Find an initial approximation (ϕ') for the geodetic latitude from equation set 7.30.

$$\phi' = 2 \arctan(e^q) - \frac{\pi}{2} = 0.746233650926621$$

Step 7: Compute $f(\phi')$ and $f'(\phi')$ from equation set 7.30.

$$f(\phi') = -q + \frac{1}{2}\ln\left[\left(\frac{1+\sin\phi'}{1-\sin\phi'}\right) \cdot \left(\frac{1 - \varepsilon\sin\phi'}{1 + \varepsilon\sin\phi'}\right)^\varepsilon\right]$$

$$f(\phi') = -4.5493507336696604 \times 10^{-3}$$

$$f'(\phi') = \frac{1 - \varepsilon^2}{(1 - \varepsilon^2\sin^2\phi')\cos\phi'} = 1.357001593456580$$

Step 8: Find an updated approximation (ϕ'') for the geodetic latitude using equation 7.31.

$$\phi'' = \phi' - \frac{f(\phi')}{f'(\phi')} = 0.7495861530575$$

Step 9: Iterate steps 7 and 8 until successive values of ϕ'' are sufficiently close after many iterations.

Final (ϕ) = 0.74958091885881052 radians = 42° 56' 52.163".

The resulting WGS 84 coordinates are:

ϕ = 42° 56' 52.163"

λ = 288° 22' 24.326"

   8.4.8 Example 8.11. Convert geodetic latitude and longitude (ϕ,λ) on the International Ellipsoid to Universal Polar Stereographic coordinates (x,y).

   This example converts geodetic latitude and longitude to UPS projection coordinates. Coordinates are computed on the International Ellipsoid.

Beginning coordinates are

ϕ = 87° 17' 14.400" S = -1.523451361952464 radians
λ = 132° 14' 52.303" E = 2.308160619653466 radians

Step 1: Compute the eccentricity (ε) of the International Ellipsoid using equation set 7.38 and a and b from Table A.1.

ε^2 = 2f - f^2 = 0.0067226700623
ε = 0.08199189022

Step 2: Compute k using equation set 7.38.

$$K = \frac{2a^2}{b} \left[\frac{(1 - \varepsilon)}{(1 + \varepsilon)}\right]^{\frac{\varepsilon}{2}} = 12713920.17558161$$

Step 3: Compute r using equation set 7.38.

$$r = K\tan\left(\frac{\pi}{4} - \frac{|\phi|}{2}\right) \cdot \left(\frac{1 + \varepsilon \sin |\phi|}{1 - \varepsilon \sin |\phi|}\right)^{\frac{\varepsilon}{2}} = 303059.087879879$$

Step 4: Compute UPS coordinates using equations 7.49 and 7.50.

x = 2,000,000 + 0.994 r sin λ = 2222991.410

y = 2,000,000 - 0.994 r cos λ = 1797464.051

8.4.9 Example 8.12. Convert Universal Polar Stereographic coordinates (X,Y) to geodetic latitude and longitude (ϕ,λ).

This example converts UPS coordinates calculated on the International Ellipsoid to geodetic latitude and longitude on the International Ellipsoid.

Beginning UPS coordinates in the Southern hemisphere are

X = 2222991.410

Y = 1797464.051

Step 1: Convert Universal Polar Stereographic coordinates (X_{UPS},Y_{UPS}) to Polar Stereographic coordinates (X_{PS},Y_{PS}).

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$$X_{PS} = \frac{X_{UPS} - 2000000}{0.994} = 224337.434205231$$

$$Y_{PS} = \frac{Y_{UPS} - 2000000}{0.994} = -203758.500100604$$

Step 2:   Compute the longitude (λ) from equation 7.39.

$$\lambda = \pi + \arctan \frac{X_{PS}}{Y_{PS}}$$

$$\lambda = 2.308160619653466 \text{ radians} = 132° \ 14'52.303" \ E$$

Step 3:   Compute k using equation 7.41, and a and b from Table A.1.

$$K = \frac{2a^2}{b} \left[\frac{(1 - \varepsilon)}{(1 + \varepsilon)} \right]^{\frac{\varepsilon}{2}} = 12713920.17558161$$

Step 4:   Compute q using equation set 7.40.

$$e^{-q} = \frac{|Y|_{PS}}{K \ |\cos \lambda|}$$

$$q = 3.73652495299514$$

Step 5:   Compute an initial estimate (ϕ_o) of the geodetic latitude using equation set 7.42.

$$\phi_o = 2 \arctan e^q - \frac{\pi}{2} = 1.52313176628106 \text{ radians}$$

Step 6:   Compute $f(\phi_o)$ and $f'(\phi_o)$ using equation set 7.42.

$$f(\phi) = -q + \frac{1}{2} \ln \left[\left(\frac{1 + \sin \phi}{1 - \sin \phi} \right) \left(\frac{1 - \varepsilon \sin \phi}{1 + \varepsilon \sin \phi} \right)^\varepsilon \right]$$

$$f(\phi_o) = -0.006730109061328$$

$$f'(\phi) = \frac{1 - \varepsilon^2}{(1 - \varepsilon^2 \sin^2\phi) \ \cos\phi} = 20.9875717033244$$

Step 7:   Compute an updated geodetic latitude ϕ_1, using equation set 7.43.

$$\phi_n = \phi_{n-1} - \frac{f(\phi_{n-1})}{f'(\phi_{n-1})} = 1.523452437445710$$

Step 8: Iterate steps 6 and 7 until successive values of the geodetic latitude are sufficiently close.

Final (ϕ) = 1.523451361952464 radians = 87° 17' 14.4"S.

The resulting geodetic latitude and longitude (ϕ,λ) coordinates are:

ϕ = 87° 17' 14.4"S

λ = 132° 14'52.303"E

    8.5 <u>List of test points</u>. The test points listed here are taken from DMA TM 8358.2 (1989) and Snyder (1987).

Mercator

      Central Meridian = 180°
      Ellipsoid   = Clarke 1866

      ϕ = 35° N
      λ = 75° W   (285° E)

      x = 11,688,673.7 m
      y = 4,139,145.6 m

      k = 1.2194146

Universal Transverse Mercator

      Central Meridian = 285° E (75° W)
      Ellipsoid   = Clarke 1866

      ϕ = 40° 30' N
      λ = 73° 30' W (286° 30' E)

      x =   627,106.5 m
      y = 4,484,124.4 m

      k = 0.9997989

Lambert Conformal Conic

      Ellipsoid   = Clarke 1866
      Standard parallels = 33° N
                       = 45° N

     Origin: ϕ_o = 23° N
             λ_o = 96° W (264° E)

ϕ = 35° N
λ = 75° W (285° E)

x = 1,894,410.9 m
y = 1,564,649.5 m

k = 0.9970171

Universal Polar Stereographic

   Ellipsoid = International

ϕ = 87° 17' 14.400"S
λ = 132° 14' 52.303"E

y = 1,797,464.051 m
x = 2,222,991.410 m

9. NOTES

9.1 <u>Intended use</u>. This Technical Report provides guidance on coordinate conversion and datum shifting for a wide range of Army mapping, charting, and positioning applications. Guidance is provided to aid the Army and others in developing software for applications of datum shifts and other transformations and in implementing this software in support of operational units.

9.2 <u>Subject (key word) listing</u>

Cartography
Charts
Coordinate Systems
Datums
Elevation
Geodetic Satellites
Geodesy
Geodetic Coordinates
Geographic Coordinates
Grids
Hydrographic Surveying
Latitude
Longitude
Map Projection
Mapping
Maps
MC&G
Mean Sea Level
Military Standards
Photogrammetry
Projections
Surveying
Topographic Maps
Topography
UTM

APPENDIX A

REFERENCE ELLIPSOID PARAMETERS

10.   GENERAL

10.1 Scope. This appendix identifies reference ellipsoids with associated geometric parameters.

20.   APPLICABLE DOCUMENTS.

The ellipsoids and associated parameters are taken from DMA TR 8350.2 (Edition 2, Insert 1.) Note: Use semi-major axis a and flattening f to calculate E and E'. Do not use semi-minor axis b, which is a derived geometric constant.

30.   GENERAL REQUIREMENTS.

30.1 Application. The ellipsoid furnishes a simple, consistent, and uniform reference system for geodetic coordinates.

TABLE A.1
REFERENCE ELLIPSOIDS

Reference Ellipsoid	ID Code	Semi-Major Axis a	Semi-Minor Axis b	Flattening f
Airy (1930)	AA	6377563.396	6356256.9090	1/299.3249646
Australian National	AN	6378160	6356774.7190	1/298.25
Bessel 1841				
Ethiopia, Indonesia, Japan, Korea	BR	6377397.155	6356078.9630	1/299.1528128
Namibia	BN	6377483.865	6356165.383	1/299.1528128
Clarke 1866	CC	6378206.4	6356583.800	1/294.9786982
Clarke 1880	CD	6378249.145	6356514.870	1/293.465
Everest				
Brunei & E. Malasia (Sabah & Sarawak)	EB	6377298.556	6356097.550	1/300.8017
India 1830	EA	6377276.345	6356075.413	1/300.8017
India 1956*	EC	6377301.243	6356100.228	1/300.8017
Pakistan*	EF	6377309.613	6356109.571	1/300.8017
W. Malasia and Singapore 1948	EE	6377304.063	6356103.039	1/300.8017
W. Malasia 1969*	ED	6377295.664	6356094.668	1/300.8017
Geodetic Reference System 1980 (GRS 80)	RF	6378137	6356752.3141	1/298.257222101
Helmert 1906	HE	6378200	6356818.170	1/298.3
Hough 1960	HO	6378270	6356794.343	1/297
Indonesian 1974	ID	6378160	6356774.504	1/298.247
International 1924	IN	6378388	6356911.946	1/297
Krassovsky 1940	KA	6378245	6356863.019	1/298.3
Modified Airy	AM	6377340.189	6356034.4480	1/299.3249646
Modified Fischer 1960 (South Asia)	FA	6378155	6356773.320	1/298.3
South American 1969	SA	6378160	6356774.719	1/298.25
World Geodetic System 1972 (WGS 72)	WD	6378135	6356750.520	1/298.26
World Geodetic System 1984 (WGS 84)	WE	6378137	6356752.3142	1/298.257223563

Source: DMA TR 8350.2 (second edition, Insert 1).
NOTE: All measurements are expressed in meters.
* Through adoption of a new yard-to-meter conversion factor in the referenced country.

APPENDIX B

ΔX, ΔY, and ΔZ DATUM SHIFT PARAMETERS

10. GENERAL

10.1 <u>Scope</u>. This appendix identifies the ellipsoid center shift transformation parameters used to transform a local geodetic system to WGS 84. An identifier, the reference ellipsoid, ΔX, ΔY, ΔZ shifts, and the estimated errors (σ) in the shifts are listed for each local geodetic system.

20. APPLICABLE DOCUMENTS.

The datum shift parameters and associated information for each local geodetic system are taken from DMA TR 8350.2 (Second Edition, Insert 1.)

30. GENERAL REQUIREMENTS

30.1 <u>Application</u>. The ΔX, ΔY, and ΔZ shifts are used for datum transformations using the three-step method discussed in section 6.3, and shown by equation sets 6.3 and 6.4.

The following symbology is used in Table B.1.

*   WGS 84 minus local geodetic system.

**   Derived from non-Dobbler sources, accuracy unknown.

***   Derived from a single Doppler station, accuracy not verified.

+   See Appendix F to transform Old Hawaiian Datum coordinates on the International Ellipsoid.

++   Also known as Hito XVIII, 1963.

+++   Use A = 6,377,483.865 meters for the Bessel 1941 Wllipsoid in Namibia.

&   Derived from non-satellite sources, accuracy unknown.

TABLE B.1

Ellipsoid Center Shift Transformation Parameters
- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*					
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$
ADINDAN Mean Value: Ethiopia and Sudan	ADI ADI-M	Clarke 1880	-166	5	-15	5	204	3
Burkina Faso***	ADI-E		-118	25	-14	25	218	25
Cameroon***	ADI-F		-134	25	-2	25	210	25
Ethiopia	ADI-A		-165	3	-11	3	206	3
Mali***	ADI-C		-123	25	-20	25	220	25
Senegal	ADI-D		-128	25	-18	25	224	25
Sudan	ADI-B		-161	3	-14	5	205	3
AFGOOYE*** Somalia	AFG	Krassovsky 1940	-43	25	-163	25	45	25
AIN EL ABD 1970	AIN	International 1924						
Bahrain Island	AIN-A		-150	25	-250	25	-1	25
Saudi Arabia	AIN-B		-143	10	-236	10	7	10
AMERICAN SAMOA 1962 American Somoa Is.	AMA	Clarke 1866	-115	25	118	25	426	25
ANNA 1 ASTRO 1965 *** Cocos Islands	ANO	Australian National	-491	25	-22	25	435	25
ANTIGUA ISLAND ASTRO 1943 Antigua, Leeward Is.	AIA	Clarke 1880	-270	25	13	25	62	25

TABLE B.1   (cont'd)

Ellipsoid Center Shift Transformation Parameters
- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*					
			ΔX	$\sigma_{\Delta x}$	ΔY	$\sigma_{\Delta y}$	ΔZ	$\sigma_{\Delta z}$
ARC 1950 Mean Value: Botswana, Lesotho, Malawi, Swaziland, Zaire, Zambia, and Zimbabwe	ARF ARF-M	Clarke 1880	-143	20	-90	33	-294	20
Botswana	ARF-A		-138	3	-105	5	-289	3
Burundi	ARF-H		-153	20	-5	20	-292	20
Lesotho	ARF-B		-125	3	-108	3	-295	8
Malawi	ARF-C		-161	9	-73	24	-317	8
Swaziland	ARF-D		-134	15	-105	15	-295	15
Zaire	ARF-E		-169	25	-19	25	-278	25
Zambia	ARF-F		-147	21	-74	21	-283	27
Zimbabwe	ARF-G		-142	5	-96	8	-293	11
ARC 1960 Mean Value: Kenya, Tanzania	ARS ARS-M	Clarke 1880	-160	20	-6	20	-302	20
ASCENSION ISLAND 1958 Ascension Island	ASC	International 1924	-205	25	107	25	53	25
ASTRO BEACON "E" 1845 Iwo Jima Island***	ATF	International 1924	145	25	75	25	-272	25

Local Geodetic System Identifier Ellipsoid   **Ellipsoid Center Shift (M)**

			Δx	σx	Δy	σy	Δz	σz
ASTRO DOS 71/4 SHB St. Helena Island***		International 1924	-320	25	550	25	-494	25
ASTRONOMIC STATION 1952 ASQ Marcus Island ***		International 1924	124	25	-234	25	-25	25
ASTRO/TERN ISLAND TRN (FRIG) 1961 *** Tern Island		International 1924	114	25	-116	25	-333	25
AUSTRALIAN GEODETIC 1966 AUA Australia and Tasmania Island		Australian National	-133	3	-48	3	148	3 `
AUSTRALIAN GEODETIC 1984 AUG Australia and Tasmania Island		Australian National	-134	2	-48	2	149	2
AYBELLA LIGHTHOUSE PHA Djibouti***		Clarke 1880	-79	25	-129	25	145	25
BELLEVUE (IGN) IBE Efate and Erromango Islands		International 1924	-127	20	-769	20	472	20
BERMUDA 1957 BER Bermuda Islands		Clarke 1866	-73	20	213	20	296	20

TABLE B.1   (cont'd)

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*							
			ΔX	$\sigma_{\Delta x}$	ΔY	$\sigma_{\Delta y}$	ΔZ	$\sigma_{\Delta z}$		
BISSAU Guinea-Bissau	BID	International 1924	-173	25	253	25	27	25		
BOGOTA OBSERVATORY Colombia	BOO	International 1924	307	6	304	5	-318	6		
BUKIT RIMPAH (BANKA) Bangka and Belitung Islands (Indonesia)	BUR	Bessel 1841	-384	**	664	**	-48	**		
CAMP AREA ASTRO Camp McMurdo Area, Antarctica	CAZ	International 1924	-104	**	-129	**	239	**		
CAMPO INCHAUSPE 1969 Argentina	CAI	International 1924	-148	5	136	5	90	5		
CANTON ASTRO 1966 Phoenix Islands	CAO	International 1924	298	15	-304	15	-375	15		
CAPE South Africa	CAP	Clarke 1880	-136	3	-108	6	-292	6		
CAPE CANAVERAL Mean Value: Florida and Bahama Islands	CAC	Clarke 1866	-2	3	151	3	181	3		

TABLE B-1   (cont'd)

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*							
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$
CARTHAGE Tunisia	CGE	Clarke 1880	-263	6	6	9	431	8
CHATHAM ISLAND ASTRO 1971 Chatham Island (New Zealand)	CHI	International 1924	175	15	-38	15	113	15
CHUA ASTRO Paraguay	CHU	International 1924	-134	6	229	9	-29	5
CORREGO ALEGRE Brazil	COA	International 1924	-206	5	172	3	-6	5
DJAKARTA (BATAVIA) Sumatra Island (Indonesia)	BAT	Bessel 1841	-377	3	681	3	-50	3
DABOLA Guinea	DAL	Clarke 1880	-83	15	37	15	124	15
DECEPTION ISLAND Deception Island, Antarctica	DID	Clarke 1880	260	20	12	20	-147	20
DOS 1968 *** Gizo Island (New Georgia Islands)	GIZ	International 1924	230	25	-199	25	-752	25
EASTER ISLAND 1967 *** Easter Island	EAS	International 1924	211	25	147	25	111	25

TABLE B-1   (cont'd)

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*					
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$
EUROPEAN 1950 Mean Value: Austria, Belgium, Denmark, Finland, France, FRG (Federal Republic of Germany) (before 1990), Gibraltar, Greece Italy, Luxembourg, Netherlands, Norway, Portugal, Spain, Sweden, and Switzerland	EUR EUR-M	International 1924	-87	3	-98	8	-121	5
Western Europe (Limited to Austria, Denmark, France, Federal Republic of Germany(before 1990), Netherlands, and Switzerland)	EUR-A		-87	3	-96	3	-120	3
Cyprus	EUR-E		-104	15	-101	15	-140	15
Egypt	EUR-F		-130	6	-117	8	-151	8
England, Ireland, Scotland, and Shetland Islands	EUR-K		-86	3	-96	3	-120	3
England, Channel Is. Scotland, and Shetland Islands	EUR-G		-86	3	-96	3	-120	3

TABLE B-1   (cont'd)

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*					
			ΔX	σΔX	ΔY	σΔY	ΔZ	σΔZ
EUROPEAN 1950 (Cont'd)								
Greece	EUR-B		-84	25	-95	25	-130	25
Iran	EUR-H		-117	9	-132	12	-164	11
Iraq, Israel, Jordan, Kuwait, Lebanon, Saudi Arabia, and Syria	EUR-S		-103	**	-106	**	-141	**
Sardinia (Italy)	EUR-I		-97	25	-103	25	-120	25
Sicily (Italy)	EUR-J		-97	20	-88	20	-135	20
Norway and Finland	EUR-C		-87	3	-95	5	-120	3
Portugal and Spain	EUR-D		-84	5	-107	6	-120	3
Malta***	EUR-L		-107	25	-88	25	-149	25
Tunisia	EUR-T		-112	25	-77	25	-145	25
EUROPEAN 1979 Mean Value: Austria, Finland, Netherlands, Norway, Spain, Sweden, and Switzerland	EUS	International 1924	-86	3	-98	3	-119	3
FORT THOMAS 1955 Nevis, St. Kitts, Leeward Islands	FOR	Clarke 1880	-7	25	215	25	225	25
GAN DATUM 1970 Republic of Maldives	GAA	International 1924	-133	25	-321	25	50	25
GEODETIC DATUM 1949 New Zealand	GEO	International 1924	84	5	-22	3	209	5

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*								
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$
GRACIOSA BASE SW 1948 (Southwest Base) Faial, Graciosa, Pico, Sao Jorge, and Terceira Islands (Azores)	GRA	International 1924	-104	3	167	3	-38	3
GUAM 1963 Guam Island	GUA	Clarke 1866	-100	3	-248	3	259	3
GUNUNG SEGARA Kalimantan Island (Indonesia)	GSE	Bessel 1841	-403	**	684	**	41	**
GUX 1 ASTRO *** Guadalcanal Island	DOB	International 1924	252	25	-209	25	-751	25
HERAT NORTH Afghanistan	HEN	International 1924	-333	**	-222	**	114	**
HJORSEY 1955 Iceland	HJO	International 1924	-73	3	46	3	-86	6
HONG KONG 1963 Hong Kong	HKD	International 1924	-156	25	-271	25	-189	25
HU-TZU-SHAN Taiwan	HTN	International 1924	-637	15	-549	15	-203	15

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*						
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$
INDIAN	IND	Everest						
Bangladesh	IND-B	India 1830	282	10	726	8	254	15
India and Nepal	IND-I	India 1956	295	12	736	10	257	
INDIAN 1954	INF	Everest						
Thailand	INF-A	India 1830	217	15	823	6	299	12
INDIAN 1960	ING	Everest						
Vietnam		India 1830						
(near 16 deg. N)	ING-A		198	25	881	25	317	25
Con Son Island								
(Vietanam)***	ING-B		182	25	915	25	344	25
INDIAN 1975	INH	Everest						
Thailand	INH-A	India 1830	209	12	818	10	290	12
INDONESIAN 1974	IDN	Indonesian	-24	25	-15	25	5	25
Indonesia***		1974						
IRELAND 1965	IRL	Modified Airy	506	3	-122	3	611	3
Ireland and Northern Ireland								
ISTS 061 ASTRO 1968	ISG	International	-794	25	119	25	-298	25
South Georgia*** Islands		1924						
ISTS 073 ASTRO 1969	IST	International	208	25	-435	25	-229	25
Diego Garcia		1924						

TABLE B-1 (cont'd)

Ellipsoid Center Shift Transformation Parameters

– Local Geodetic System to WGS 84 –

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*							
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$		
JOHNSTON ISLAND 1961 Johnston Island	JOH	International 1924	189	25	-79	25	-202	25		
KANDAWALA Sri Lanka	KAN	Everest 1830	-97	20	787	20	86	20		
KERGUELEN ISLAND 1949 Kerguelen Island ***	KEG	International 1924	145	25	-187	25	103	25		
KERTAU 1948 (Malaysian Revised) West Malaysia and Singapore	KEA	Everest 1948	-11	10	851	8	5	6		
KUSAIE ASTRO 1951 Caroline Islands, Fed. States of Micronesia***	KUS	International 1924	647	25	1777	25	-1124	25		
L.C. 5 ASTRO 1961 *** Cayman Brac Island	LCF	Clarke 1866	42	25	124	25	147	25		
LIBERIA 1964 Liberia	LIB	Clarke 1880	-90	15	40	15	88	15		
LEIGON Ghana	LEH	Clarke 1880	-130	2	29	3	364	2		

Ellipsoid Center Shift Transformation Parameters

– Local Geodetic System to WGS 84 –

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*					
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$
LUZON Philippines (Excluding Mindanao Island	LUZ LUZ-A	Clarke 1866	-133	8	-77	11	-51	9
Mindanao Island ***	LUZ-B		-133	25	-79	25	-72	25
MAHE 1971 *** Mahe Island	MIK	Clarke 1880	41	25	-220	25	-134	25
MASSAWA *** Eritrea (Ethiopia)	MAS	Bessel 1841	639	25	405	25	60	25
MERCHICH Morocco	MER	Clarke 1880	31	5	146	3	47	3
MIDWAY ASTRO 1961 *** Midway Island	MID	International 1924	912	25	-58	25	1227	25
MINNA Cameroon Nigeria	MIN MIN-A MIN-B	Clarke 1880	-81 -92	25 3	-84 -93	25 6	115 122	25 5
MONTSERRAT ISLAND ASTRO 1958 Montserrat, Leeward Islands***	ASM	Clarke 1880	174	25	359	25	365	25

TABLE B-1   (cont'd)

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*					
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$
M'PORALOKO Gabon***	MPO	Clarke 1880	-74	25	-130	25	42	25
NAHRWAN	NAH	Clarke 1880						
Masirah Island(Oman)	NAH-A		-247	25	-148	25	369	25
United Arab Emirates	NAH-B		-249	25	-156	25	381	25
Saudi Arabia	NAH-C		-243	20	-192	20	477	20
NAPARIMA, BWI Trinidad and Tobago	NAP	International 1924	-10	15	375	15	165	15
NORTH AMERICAN 1927 Mean value: (CONUS)	NAS NAS-C	Clarke 1866	-8	5	160	5	176	6
Western United States (Arizona, Arkansas, California, Colorado, Idaho, Iowa, Kansas, Montana, Nebraska, Nevada, New Mexico, North Dakota, Oklahoma, Oregon, South Dakota, Texas, Utah, Washington, and Wyoming)	NAS-B		-8	5	159	3	175	3

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*					
		ΔX	σ_ΔX	ΔY	σ_ΔY	ΔZ	σ_ΔZ
NORTH AMERICAN 1927 (CONT'D)	NAS						
Eastern United States (Alabama, Connecticut, Delaware, District of Columbia, Florida, Georgia, Illinois, Indiana, Kentucky, Louisiana, Maine, Maryland, Massachusetts, Michigan, Minnesota, Mississippi, Missouri, New Hampshire, New Jersey, New York, North Carolina, Ohio, Pennsylvania, Rhode Island, South Carolina, Tennessee, Vermont, Virginia, West Virginia, and Wisconsin	NAS-A	-9	5	161	5	179	8
Alaska (Excluding Aleutian Is.)	NAS-D	-5	5	135	9	172	5
Aleutian Islands East of 180 degrees W	NAS-V	-2	6	152	8	149	10
West of 180 degrees W	NAS-W	2	10	204	10	105	10

TABLE B-1   (cont'd)

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*					
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$
NORTH AMERICAN 1927 (cont'd)		Clarke 1866						
Canada (Mean value, including Newfoundland Island)	NAS-E		-10	15	158	11	187	6
Alberta and British Columbia	NAS-F		-7	8	162	8	188	6
East Canada (Newfoundland, New Brunswick, Nova Scotia, and Quebec)	NAS-G		-22	6	160	6	190	3
Manitoba and Ontario	NAS-H		-9	9	157	5	184	5
Northwest Territories and Saskatchewan	NAS-I		4	5	159	5	188	3
Yukon	NAS-J		-7	5	139	8	181	3
Canal Zone	NAS-O		0	20	125	20	201	20
Bahamas (Excluding San Salvador Island)	NAS-Q		-4	5	154	3	178	5
San Salvador Island ***	NAS-R		1	25	140	25	165	25
Central America (Belize, Costa Rica, El Salvador, Guatemala, Honduras, and Nicaragua)	NAS-N		0	8	125	3	194	5

TABLE B-1   (cont'd)

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*					
			ΔX	σ$_{ΔX}$	ΔY	σ$_{ΔY}$	ΔZ	σ$_{ΔZ}$
NORTH AMERICAN 1927 (CONT'D) Caribbean (Antigua Islands, Barbados, Barbuda, Cuba, Caicos Islands, Dominican Republic, Grand Cayman, Jamaica, and Turks Islands)	NAS-P		-3	3	142	9	183	12
Cuba ***	NAS-T		-9	25	152	25	178	25
Greenland (Hayes Peninsula)	NAS-U		11	25	114	25	195	25
Mexico	NAS-L		-12	8	130	6	190	6
NORTH AMERICAN 1983	NAR	GRS 80						
Alaska (Excluding Aleutian Is.)	NAR-A		0	2	0	2	0	2
Aleutian Islands	NAR-E		-2	5	0	2	4	5
Canada	NAR-B		0	2	0	2	0	2
CONUS	NAR-C		0	2	0	2	0	2
Hawaii	NAR-H		1	2	1	2	-1	2
Mexico and Central America	NAR-D		0	2	0	2	0	2
NORTH SAHARA 1959 Algeria	NSD	Clarke 1880	-186	25	-93	25	310	25

TABLE B-1   (cont'd)

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters) *					
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$
OBSERVATORIO METEROLOGICO 1939 Corvo and Flores Islands (Azores)	FLO	International 1924	-425	20	-169	20	81	20
OLD EGYPTIAN 1907 Egypt	OEG	Helmert 1906	-130	3	110	6	-13	8
OLD HAWAIIAN	OHA	Clarke 1866 +						
Mean Value	OHA-M		61	25	-285	20	-181	20
Hawaii	OHA-A		89	25	-279	25	-183	25
Kauai	OHA-B		45	20	-290	20	-172	20
Maui	OHA-C		65	25	-290	25	-190	25
Oahu	OHA-D		58	10	-283	6	-182	
OMAN Oman	FAH	Clarke 1880	-346	3	-1	3	224	9
ORDNANCE SURVEY OF GREAT BRITAIN 1936	OGB	Airy						
Mean Value: England, Isle of Man, Scotland, Shetland Is., and Wales	OGB-M		375	10	-111	10	431	15
England	OGB-A		371	5	-112	5	434	6
England, Isle of Man, and Wales	OGB-B		371	10	-111	10	434	15
Scotland and Shetland Island	OGB-C		384	10	-111	10	425	10

TABLE B-1   (cont'd)

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*					
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$
REUNION *** Mascarene Island	REU	International 1924	94	25	-948	25	-1262	25
ROME 1940 *** (Roma Monte Maria) Sardinia Island	MOD	International 1924	-225	25	-65	25	9	25
S-42 (PULKOVO 1942) Hungary	SPK	Krassovsky 1940	28	2	-121	2	-77	2
S-JTSK Czech Rep., Slovakia	CCD	Bessel 1841	589	4	76	2	480	3
SANTO (DOS) 1965 *** Espirito Santo Island	SAE	International 1924	170	25	42	25	84	25
SAO BRAZ Sao Miguel, Santa Maria Islands (Azores)	SAO	International 1924	-203	25	141	25	53	25
SAPPER HILL 1943 East Falkland Island	SAP	International 1924	-355	1	21	1	72	1
SCHWARZECK Namibia	SCK	Bessel 1841++	616	20	97	20	-251	20
SELVAGEM GRADE 1938 Salvage Islands***	SGM	International 1924	-289	25	-124	25	60	25

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters) *					
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$
PROVISIONAL SOUTH AMERICAN 1956	PRP	International 1924						
Mean Value: Bolivia, Chile, Colombia, Ecuador, Guyana, Peru, and Venezuela	PRP-M		-288	17	175	27	-376	27
Bolivia	PRP-A		-270	5	188	11	-388	14
Northern Chile *** (near 19°S)	PRP-B		-270	25	183	25	-390	25
Southern Chile (near 43°S)	PRP-C		-305	20	243	20	-442	20
Colombia	PRP-D		-282	15	169	15	-371	15
Ecuador	PRP-E		-278	3	171	5	-367	3
Guyana	PRP-F		-298	6	159	14	-369	5
Peru	PRP-G		-279	6	175	8	-379	12
Venezuela	PRP-H		-295	9	173	14	-371	15
PUERTO RICO Puerto Rico and Virgin Islands	PUR	Clarke 1866	11	3	72	3	-101	3
QATAR NATIONAL Qatar (Persian Gulf)	QAT	International 1924	-128	20	-283	20	22	20
QORNOQ South Greenland	QUO	International 1924	164	25	138	25	-189	32

TABLE B-1   (cont'd)

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters) *							
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$
REUNION *** Mascarene Island	REU	International 1924	94	25	-948	25	-1262	25
ROME 1940 *** (Roma Monte Maria) Sardinia Island	MOD	International 1924	-225	25	-65	25	9	25
S-42 (PULKOVO 1942) Hungary	SPK	Krassovsky 1940	28	2	-121	2	-77	2
S-JTSK Czech Rep., Slovakia	CCD	Bessel 1841	589	4	76	2	480	3
SANTO (DOS) 1965 *** Espirito Santo Island	SAE	International 1924	170	25	42	25	84	25
SAO BRAZ Sao Miguel, Santa Maria Islands (Azores)	SAO	International 1924	-203	25	141	25	53	25
SAPPER HILL 1943 East Falkland Island	SAP	International 1924	-355	1	21	1	72	1
SCHWARZECK Namibia	SCK	Bessel 1841++	616	20	97	20	-251	20
SELVAGEM GRADE 1938 Salvage Islands***	SGM	International 1924	-289	25	-124	25	60	25

TABLE B-1   (cont'd)

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*					
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$
SOUTH AMERICAN 1969 Mean Value: Argentina, Bolivia, Brazil, Chile, Colombia, Ecuador, Guyana, Paraguay, Peru, Trinidad and Tobago, and Venezuela	SAN SAN-M	South American 1969	-57	15	1	6	-41	9
Argentina	SAN-A		-62	5	-1	5	-37	5
Baltra, Galapagos Is.***	SAN-J		-47	25	26	25	-42	25
Bolivia	SAN-B		-61	15	2	15	-48	15
Brazil	SAN-C		-60	3	-2	5	-41	5
Chile	SAN-D		-75	15	-1	8	-44	1
Colombia	SAN-E		-44	6	6	6	-36	5
Ecuador (excluding Galapagos Islands)	SAN-F		-48	3	3	3	-44	3
Guyana	SAN-G		-53	9	3	5	-47	5
Paraguay	SAN-H		-61	15	2	15	-33	15
Peru	SAN-I		-58	5	0	5	-44	5
Trinidad and Tobago ***	SAN-K		-45	25	12	25	-33	25
Venezuela	SAN-L		-45	3	8	6	-33	3
SOUTH ASIA (Southeast Asia) Singapore ***	SOA	Modified Fischer 1960	7	25	-10	25	-26	25
TANANARIVE OBSERVATORY 1925 Madagascar	TAN	International 1924	-189	&	-242	&	-91	&

Ellipsoid Center Shift Transformation Parameters

- Local Geodetic System to WGS 84 -

LOCAL GEODETIC SYSTEM*	IDENTIFIER	ELLIPSOID	ELLIPSOID CENTER SHIFT (meters)*					
			ΔX	$\sigma_{\Delta X}$	ΔY	$\sigma_{\Delta Y}$	ΔZ	$\sigma_{\Delta Z}$
TIMBALAI 1948 Brunei and East Malaysia (Sarawak and Sabah)	TIL	Everest	-679	10	669	10	-48	12
TOKYO Mean Value: Japan, S. Korea, and Okinawa	TOY TOY-M	Bessel 1841	-148	20	507	5	685	20
Japan	TOY-A		-148	8	507	5	685	8
South Korea	TOY-B		-146	8	507	5	687	8
Okinawa	TOY-C		-158	20	507	5	676	20
TRISTAN ASTRO 1968 *** Tristan da Cunha	TDC	International 1924	-632	25	438	25	-609	25
VOIROL 1960	VOR	Clarke 1880	-123	25	-206	25	219	25
VOITI LEVU 1916 *** Viti Levu Island (Fiji Islands)	MVS	Clarke 1880	51	25	391	25	-36	25
WAKE-ENIWETOK 1960 Marshall Islands	ENW	Hough	102	3	52	3	-38	3
WAKE ISLAND ASTRO 1952 Wake Atoll	WAK	International 1924	276	25	-57	25	149	25
YACARE Uruguay	YAC	International 1924	-155	**	171	**	37	**
ZANDERIJ Surinam	ZAN	International 1924	-265	5	120	5	-358	8

APPENDIX C

DATUM LIST

10. GENERAL

10.1 Scope. This appendix consists of a list of countries and geographic areas and their associated datums.

20. APPLICABLE DOCUMENTS.

This section is not applicable to this appendix.

30. GENERAL REQUIREMENTS

30.1 Application. The intended use of this appendix is to establish a table of countries or geographic areas for quick reference to their respective datum.

TABLE C-1
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Afganistan	Herat North
Alaska	North American 1927, 1983
Alberta	North American 1927, 1983
Antigua	Antigua Island Astro 1943 North American 1927
Argentina	Campo Inchauspe, South American 1969
Ascension Island	Ascension Island 1958
Australia	Australian Geodetic 1966, 1984
Austria	European 1950, 1979
Bahama Islands	Cape Canaveral, North American 1927
Bahrain Island	Ain el ABD 1970
Baltra, Galapagos Islands	South American 1969
Bangladesh	Indian
Barbados	North American 1927
Barbuda	North American 1927

TABLE C-1   (cont'd)
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Belgium	European 1950
Belize	North American 1927
Bolivia	Provisional South American 1956, South American 1969
Botswana	ARC 1950
Brazil	Corrego Alegre, South American 1969
British Columbia	North American 1927, 1983
Bruñei	Timbalai 1948
Burkina Faso	Adindan Point 58
Burundi	ARC 1950
Caicos Islands	North American 1927
Cameroon	Adindan Minna
Canada	North American 1927, 1983
Canal Zone	North American 1927
Canary Islands	Pico de Las Nieves

TABLE C-1   (cont'd)
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Caribbean	North American 1927
Caroline Islands	Kusaie Astro 1951
Cayman Brac Island	L.C. 5 Astro 1961
Central America	North American 1927, 1983
Chatham Island	Chatham Island Astro 1971
Chile	Provisional South American 1956, South American 1969, PS Chile 1963
Cocos Island	Anna 1 Astro 1965
Colombia	Bogota Observatory, South American 1969
Congo	Point Noire 1948
CONUS	North American 1927, 1983
Corvo Island (Azores)	Observatorio Meteorologico 1939
Costa Rica	North American 1927
Cuba	North American 1927
Cyprus	European 1950

TABLE C-1   (cont'd)
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Denmark	European 1950
Diego Garcia	ISTS 073 Astro 1969
Djibouti	Aybella Lighthouse
Dominican Republic	North American 1927, 1983
East Canada	North American 1927, 1983
East Falkland Island	Sapper Hill 1943
Easter Island	Easter Island 1967
Eastern United States	North American 1927, 1983
Ecuador	Provisional South American 1956 South American 1969
Eftate Island	Bellevue (IGN)
Egypt	European 1950 Old Egyptian
El Salvador	North American 1927, 1983
England	European 1950, Ordnance Survey of Great Britain 1936
Erromango Island	Bellevue (IGN)

TABLE C-1   (cont'd)
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Eritrea	Massawa
Espirito Santo Island	Santo (DOS) 1965
Ethiopia	Adindan
Faial Island (Azores)	Graciosa Base SW 1948
Federal Republic of Germany (before 1990)	European 1950
Federal States of Micronesia	Kusaie 1951
Finland	European 1950, 1979
Flores Island (Azores)	Observatorio Meteorologico 1939
Florida	Cape Canaveral North American 1927, 1983
France	European 1950
Gabon	M'poraloko
Ghana	Leigon
Gibraltar	European 1950

TABLE C-1   (cont'd)
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Gizo Island (New Georia Islands)	DOS 1968
Graciosa Island	Graciosa Base SW
Grand Canyon	North American 1950
Greece	European 1950
Greenland (Hayes Peninsula)	North American 1927
Guadacanal Island	Gux 1 Astro
Guam	Guam 1963
Guatemala	North American 1927, 1983
Guinea-Bissau	Bissau
Guinea	Dabola
Guyana	Provisional South American 1956 South American 1969
Hawaii	Old Hawaiian North American 1983
Honduras	North American 1927, 1983

TABLE C-1   (cont'd)
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Hong Kong	Hong Kong 1963
Iceland	Hjorsey 1955
India	Indian
Iran	European 1950
Iraq	European 1950
Ireland	European 1950, Ireland 1965
Isle of Man	Ordnance Survey of Great Britain 1936
Israel	European 1950
Italy	European 1950
Iwo Jima	Astro Beacon "E"
Jamaica	North American 1927
Japan	Tokyo
Johnston Island	Johnston Island 1961
Jordan	European 1950
Kalimantan Island (Indonesia)	Gunung Segara

TABLE C-1   (cont'd)
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Kauai	Old Hawaiian North American 1983
Kenya	Arc 1960
Kerguelen Island	Kerguelen Island 1949
Korea	Tokyo
Kuwait	European 1950
Lebanon	European 1950
Leeward Islands	Fort Thomas 1955 Antigua Island Astro 1943 Montserrat Island Astro 1958
Liberia	Liberia 1964
Luxembourg	European 1950
Madagascar	Tananarive Observatory
Madeira Islands	Porto Santo 1936
Mahe Island	Mahe 1971
Malawi	Arc 1950

TABLE C-1   (cont'd)
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Malaysia	Timbalai 1948
Maldives, Republic of	Gan
Mali	Adindan
Malta	European 1950
Manitoba	North American 1927, 1983
Marcus Islands	Astronomic Station 1952
Marshall Islands	Wake Eniwetok 1960
Mascarene Island	Reunion
Masirah Island (Oman)	Nahrwan
Maui	Old Hawaiian North American 1983
Mexico	North American 1927, 1983
Micronesia	Kusaie 1951
Midway Island	Midway Astro 1961
Mindanao Island	Luzon

TABLE C-1   (cont'd)
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Montserrat	Montserrat Island Astro 1958
Morocco	Merchich
Namibia	Schwarzeck
Nepal	Indian
Netherlands	European 1950, 1979
Nevis	Fort Thomas 1955
New Brunswick	North American 1927, 1983
New Foundland	North American 1927, 1983
New Zealand	Geodetic Datum 1949
Nicaragua	North American 1927
Niger	Point 58
Nigeria	Minna
Northern Ireland	Ireland 1965
Northwest Territories	North American 1927, 1983
Norway	European 1950, 1979

TABLE C-1   (cont'd)
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Nova Scotia	North American 1927, 1983
Oahu	Old Hawaiian North American 1983
Okinawa	Tokyo
Oman	Oman
Ontario	North American 1927, 1983
Paraguay	Chua Astro, South American 1969
Peru	Provisional South American 1956, South American 1969
Philippines	Luzon
Phoenix Islands	Canton Astro 1966
Pico Island (Azores)	Graciosa Base SW
Pitcairn Island	Pitcairn Astro 1967
Porto Santo Island	Porto Santo 1936
Portugal	European 1950
Puerto Rico	Puerto Rico

TABLE C-1   (cont'd)
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Qatar	Qatar National
Quebec	North American 1927, 1983
Salvage Islands	Selvagem Grade 1938
San Salvador Island	North American 1927
Santa Maria Islands (Azores)	Sao Braz
São Jorge (Azores)	Graciosa Base SW
São Miguel Island	São Braz
Sarawak and Sabah	Timbalai 1948
Sardinia	European 1950, Rome 1940
Saskatchewan	North American 1927, 1983
Saudi Arabia	Ain El Abd 1970, European 1950, Nahrwan
Scotland	European 1950, Ordnance Survey of Great Britain 1936
Senegal	Adindan
Shetland Islands	European 1950, Ordnance Survey of Great Britain 1936

TABLE C-1   (cont'd)
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Sicily	European 1950
Singapore	Kertau 1948, South Asia
Somalia	Afgooye
South Africa	Cape
South Chile	Provisional South Chilean 1963
South Geogia Islands	ISTS Astro 1968
South Greenland	Qornoq
Spain	European 1950, 1979
Sri Lanka	Kandawala
St. Helena Island	Astro Dos 71/4
St. Kitts	Fort Thomas 1955
Sudan	Adindan
Sumatra Island (Indonesia)	Djakarta (Botavia)
Surinam	Zanderij
Swaziland	Arc 1950

TABLE C-1   (cont'd)
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Sweden	European 1950, 1979
Switzerland	European 1950, 1979
Syria	European 1950, 1979
Taiwan	Hu-Tzu-Shan
Tanzania	Arc 1960
Tasmania Island	Australian Geodetic 1966, 1984
Terceira Islands (Azores)	Graciosa Base SW
Tern Island	Astro/Tern Island (Frig) 1961
Thailand	Indian
Trinidad and Tobago	Naparima, BWI; South American 1969
Tristan da Cunha	Tristan Astro
Tunisia	Carthage
Turk and Caicos Islands	North American 1927
United Arab Emirates	Nahrwan
Uruguay	Yacare

TABLE C-1   (cont'd)
Countries and Their Associated Datums

COUNTRY/GEOGRAPHIC AREA	DATUM
Venezuela	Provisional South American 1956, South American 1969
Vietnam	Indian
Virgin Islands	Puerto Rico
Viti Levu Island (Fiji Island)	Indian
Wales	Ordnance Survey of Great Britain 1936
Wake Atoll	Wake Island Astro 1952
Western Europe	European 1950
West Malaysia	Kertau 1948
Western United States	North American 1927, 1983
Yukon	North American 1927, 1983
Zaire	Arc 1950
Zambia	Arc 1950
Zimbabwe	Arc 1950

WFS 84 Geoid Separation Computation

10.0   GENERAL.

10.1 Scope. This appendix provides the bilinear method for
interpolating geoid separation and WGS 84 Geoid Separations on a
10° x 10° grid. These separations are expressed in meters.

20.   APPLICABLE DOCUMENTS.

This section is not applicable to this appendix.

30.   GENERAL REQUIREMENTS.

30.1   The Bilinear Interpolation Method of Calculating WGS 84
Geoid Separation ($N_{WGS\ 84}$). The bilinear interpolation method uses
interpolation in two directions (ϕ, λ) to approximate the geoid
separation $N_{WGS\ 84}$, from four known WGS 84 geoid separations given
in a look-up table. The bi-linear interpolation method has the
form given in equation set D.1. Figure D.1 shows the relationship
between the positions of the known $N_{WGS\ 84}$ values and the
interpolation point.

$$N_p(\phi, \lambda) = a_0 + a_1 X + a_2 Y + a_3 XY \qquad (D.1)$$

Where

$N_p(\phi, \lambda)$ = Geoid Separation (N) to be interpolated at Point
$P(\phi, \lambda)$.

$a_0 = N_1$

$a_1 = N_2 - N_1$

$a_2 = N_4 - N_1$

$a_3 = N_1 + N_3 - N_2 - N_4$

$$X = \frac{(\lambda - \lambda_1)}{(\lambda_2 - \lambda_1)}$$

$$Y = \frac{(\phi - \phi_1)}{(\phi_2 - \phi_1)}$$

ϕ = Geodetic latitude of Point P
λ = Geodetic longitude of Point P
N_1, N_2, N_3, N_4 = Known geoid heights at grid points used in the interpolation process.
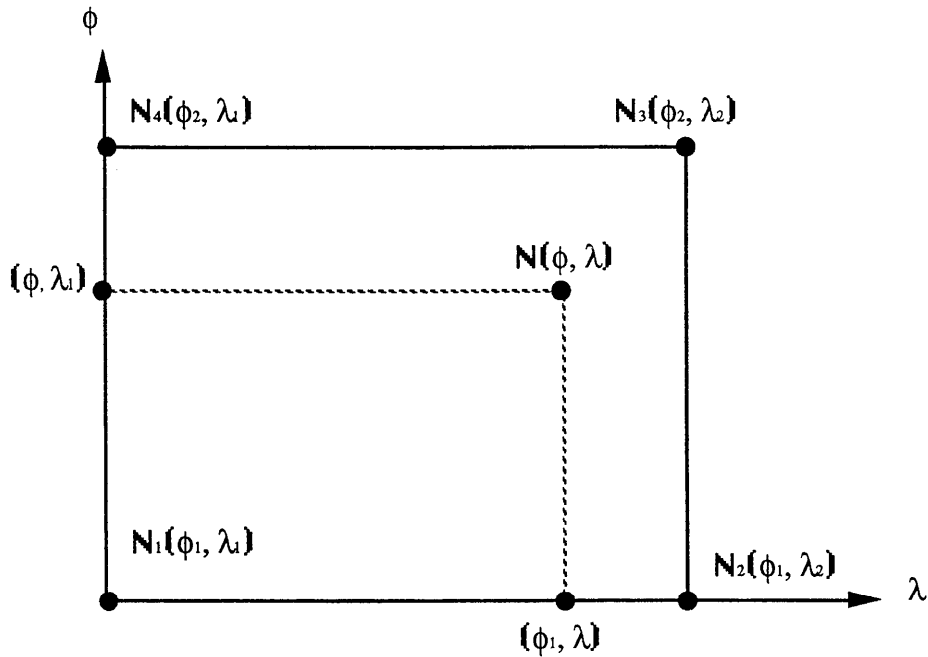
Figure F-1 Coordinate System Associated with
Geoid Seperation Bi-Linear Interpolation Method

30.2 N_WGS 84 Tables. N_WGS 84 values are tabulated on a 10° x 10° grid in Table D.1. The values in Table D.1 are suitable for use in moderate accuracy mapping and charting applications. Qualified users may obtain digitally-stored tables of N_WGS 84 values on 30' x 30' and 1° x 1° grids for survey and high accuracy mapping and charting applications.

TABLE D.1

WGS 84 Geoid Separations
(N = M = 180, 10° x 10° Grid, Units = meters)

Latitude (degrees) × Longitude (degrees)

Lat\Lon	0	10	20	30	40	50	60	70	80	90	100	110	120	130	140	150	160	170
90	13	13	13	13	13	13	13	13	13	13	13	13	13	13	13	13	13	13
80	33	34	28	23	17	13	9	4	4	1	-2	-2	0	2	3	2	1	1
70	51	43	29	20	12	5	-2	-10	-14	-12	-10	-14	-12	-6	-2	3	6	4
60	47	41	21	18	14	7	-3	-22	-29	-32	-32	-26	-15	-2	13	17	19	6
50	47	48	42	28	12	-10	-19	-33	-43	-42	-43	-29	-2	17	23	22	6	2
40	52	48	35	40	33	-9	-28	-39	-48	-59	-50	-28	3	23	37	18	-1	-11
30	36	28	29	17	12	-20	-15	-40	-33	-34	-34	-28	7	29	43	20	4	-6
20	31	26	15	6	1	-29	-44	-61	-67	-59	-36	-11	21	39	49	39	22	10
10	22	23	2	-3	-7	-36	-59	-90	-95	-63	-24	12	53	60	58	46	36	26
0	18	12	-13	-9	-28	-49	-62	-89	-102	-63	-9	33	58	73	74	63	50	32
-10	12	13	-2	-14	-25	-32	-38	-60	-75	-63	-26	0	35	52	68	76	64	52
-20	17	23	21	8	-9	-10	-11	-20	-40	-47	-45	-25	5	23	45	58	57	63
-30	22	27	34	29	14	15	15	7	-9	-25	-37	-39	-23	-14	15	33	34	45
-40	18	26	31	33	39	41	30	24	13	-2	-20	-32	-33	-27	-14	-2	5	20
-50	25	26	34	39	45	45	38	39	28	13	-1	-15	-22	-22	-18	-15	-14	-10
-60	16	19	25	30	35	35	33	30	27	10	-2	-14	-23	-30	-33	-29	-35	-43
-70	16	16	17	21	20	26	26	22	16	10	-1	-16	-29	-36	-46	-55	-54	-59
-80	-4	-1	1	4	4	6	5	4	2	-6	-15	-24	-33	-40	-48	-50	-53	-52
-90	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30

TABLE D.1 (continued)

WGS 84 Geoid Separations
(N = M = 180, 10° x 10° Grid, Units = meters)

Latitude (degrees)

Lat \ Long	180	190	200	210	220	230	240	250	260	270	280	290	300	310	320	330	340	350
90	13	13	13	13	13	13	13	13	13	13	13	13	13	13	13	13	13	13
80	3	1	-2	-3	-3	-3	-1	3	1	5	9	11	19	27	31	34	33	34
70	2	2	1	-1	-3	-7	-14	-24	-27	-25	-19	3	24	37	47	60	61	58
60	2	9	17	10	13	1	-14	-30	-39	-46	-42	-21	6	29	49	65	60	57
50	-8	8	8	1	-11	-19	-16	-18	-22	-35	-40	-26	-12	24	45	63	62	59
40	-12	-10	-13	-20	-31	-34	-21	-16	-26	-34	-33	-35	-26	2	33	59	52	51
30	-7	-5	-8	-15	-28	-40	-42	-29	-22	-26	-32	-51	-40	-17	17	31	34	44
20	5	10	7	-7	-23	-39	-47	-34	-9	-10	-20	-45	-48	-32	-9	17	25	31
10	13	12	11	2	-11	-28	-38	-29	-10	3	1	-11	-41	-42	-16	3	17	33
0	22	16	17	13	1	-12	-23	-20	-14	-3	14	10	-15	-27	-18	3	12	20
-10	36	22	11	6	-1	-8	-10	-8	-11	-9	1	32	4	-18	-13	-9	4	14
-20	51	27	10	0	-9	-11	-5	-2	-3	-1	9	35	20	-5	-6	-5	0	13
-30	46	22	5	-2	-8	-13	-10	-7	-4	1	9	32	16	4	-8	4	12	15
-40	21	6	1	-7	-12	-12	-12	-10	-7	-1	8	23	15	-2	-6	6	21	24
-50	-15	-18	-18	-16	-17	-15	-10	-10	-8	-2	6	14	13	3	3	10	20	27
-60	-45	-43	-37	-32	-30	-26	-23	-22	-16	-10	-2	10	20	20	21	24	22	17
-70	-61	-60	-61	-55	-49	-44	-38	-31	-25	-16	-6	1	4	5	4	2	6	12
-80	-53	-54	-55	-52	-48	-42	-38	-38	-29	-26	-26	-24	-23	-21	-19	-16	-12	-8
-90	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30	-30

Longitude (degrees)

APPENDIX E

Seven-Parameter Geometric Datum Shifts

10. GENERAL

10.1 <u>Scope</u>. This appendix provides the data needed to perform seven-parameter geometric datum shifts between two local geodetic systems and WGS 84.

20. APPLICABLE DOCUMENTS.

This section is not applicable to this appendix.

30. GENERAL REQUIREMENTS

30.1 <u>Application</u>. Table E-1 lists the coverage area, ellipsoid applicable to the local geodetic system, ellipsoid center shifts (ΔX, ΔY, ΔZ), axes rotations (ε, ψ, ω), and scale factor change (ΔS) for use in seven-parameter geometric datum shifts between two local geodetic systems and WGS 84. The root-mean square differences between Doppler-derived WGS 84 geodetic coordinates and WGS 84 coordinates, computed from local system coordinates using the seven-parameter model, are provided as an indication of the datum shift consistency.

TABLE E.1

Seven-Parameter Local Geodetic System to WGS 84

Datum Transformations

DATUM: EUROPEAN 1950 (EUR-M)

COVERAGE AREA: Western Europe

ELLIPSOID: International

$\Delta X = -102$ m $\Delta Y = -102$ m $\Delta Z = -129$ m

$\varepsilon = 0.413"$ $\psi = -0.184"$ $\omega = 0.385"$ $\Delta S = 2.4664 \times 10^{-6}$

RMS DIFFERENCES : WGS 84 surveyed, minus computed coordinates

$\Delta\phi = 2$ m $\Delta\lambda = 3$ m $\Delta H = 2$ m

DATUM: ORDNANCE SURVEY OF GREAT BRITAIN 1936

COVERAGE AREA: England, Isle of Man, Scotland, Shetland Islands,

and Wales

ELLIPSOID: Airy

$\Delta X = 446$ m $\Delta Y = -99$ m $\Delta Z = 544$ m

$\varepsilon = -0.945"$ $\psi = -0.261"$ $\omega = -0.435"$ $\Delta S = -20.8927 \times 10^{-6}$

RMS DIFFERENCES : WGS 84 surveyed minus computed coordinates

$\Delta\phi = 2$ m $\Delta\lambda = 2$ m $\Delta H = 1$ m

APPENDIX F

Old Hawaiian Datum with International Ellipsoid

10. GENERAL

10.1 _Scope_. This appendix provides the information needed to shift coordinates between the Clarke 1866 and International in the Old Hawaiian datum.

20. APPLICABLE DOCUMENTS

DMA SGG MEMORANDOM FOR RECORD
SUBJECT: Old Hawaiian Datum with
International Ellipsoid
13 June 1990

Army Map Service GEODETIC MEMO
No. 687, SUBJECT: Hawaiian Islands,
UTM Coordinates and International Spheroid
Geographic Positions. P.O.# 23832-002,
29 November 1950.

30. GENERAL REQUIREMENTS

30.1 _Background_. Positions on the Old Hawaiian datum may be on either the Clarke 1866 ellipsoid or the International ellipsoid. Users of Old Hawaiian datum positions must determine which ellipsoid applies. Datum shift parameters from Old Hawaiian datum to WGS 84 have been determined only for the Clarke 1866 ellipsoid. Old Hawaiian datum coordinates on the international ellipsoid must be converted to the Clarke 1866 ellipsoid before being shifted to WGS 84. The conversion between International and Clarke 1866 ellipsoids has been defined in terms of UTM coordinates. Geodetic latitude and longitude must be converted to UTM before changing ellipsoids. Height conversions should not be attempted as no conversion method is available.

30.2 _Application_. To convert Old Hawaiian datum UTM coordinates on the International ellipsoid (E_{INT}, N_{INT}) to UTM coordinates on the Clarke 1866 ellipsoid (E_{C66}, N_{C66}) use:

In UTM Zone 4

$$E_{C66} = E_{INT} - 3$$
$$N_{C66} = N_{INT} - 169$$
(I.1)

In UTM Zone 5

$$E_{C66} = E_{INT} + 13$$
$$N_{C66} = N_{INT} - 169$$
(I.2)

To convert Old Hawaiian datum UTM coordinates on the Clarke 1866 ellipsoid to the International ellipsoid use:

In UTM Zone 4

$$E_{INT} = E_{C66} + 3$$
$$N_{INT} = N_{C66} + 169$$

(I.3)

In UTM Zone 5

$$E_{INT} = E_{C66} - 13$$
$$N_{INT} = N_{C66} + 169$$

(I.4)

30.3 <u>Example</u>. Convert geodetic survey coordinates on the Old Hawaiian datum, International ellipsoid, to WGS 84 geodetic coordinates.

$$\phi_{INT} = 21° \ 19' \ 37.431"N$$
$$\lambda_{INT} = 157° \ 58' \ 25.610"W$$

This position is on the island of Oahu.

a. Convert ϕ_{INT}, λ_{INT} to UTM:

Zone = 4

$$E_{INT} = 606,428$$
$$N_{INT} = 2,358,722$$

b. Convert E_{INT}, N_{INT} to the Clarke 1866 ellipsoid using equation set I.1, for zone 4.

$$E_{C66} = 606,425$$
$$N_{C66} = 2,358,553$$

c. Convert E_{C66}, N_{C66} to geodetic coordinates on the Clarke 1866 ellipsoid.

$$\phi_{C66} = 21° \ 19' \ 37.425"N$$
$$\lambda_{C66} = 157° \ 58' \ 25.631"W$$

d. ΔX, ΔY, ΔZ shift parameters for datum identifier code OHA-D should be used for shifting geodetic survey coordinates on Oahu.

e. Shift ϕ_{C66}, λ_{C66} to WGS 84 using, for example, the three-step method.

$$\phi_{WGS\ 84} = 21°\ 19'\ 26.069"N$$
$$\lambda_{WGS\ 84} = 157°\ 58'\ 15.766"W$$